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## Impact of fire and frost drought events on vegetation greenness and albedo in Subarctic Scandinavia

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Master thesis, 30 credits, in Physical Geography and Ecosystem Science

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#### Abstract

The frequency in climate extremes such as drought, heavy rainfalls, and extreme temperatures is predicted to increase with climate change. These extremes can potentially stress and damage vegetation over large areas, leading to altered carbon cycle, vegetation greenness, albedo, and shifted species composition. This study focuses on two different types of extreme events; frost drought, caused by winter warming, and fires. The first aim was to study how these events affect vegetation greenness and albedo in the Subarctic region of Norway and Sweden. A second aim was to test whether the effects differ between land cover categories. A before-after-impact approach was applied to determine the effect of the events.

When studying all sites, the results show no impact for neither albedo nor vegetation greenness as caused by the fire events. However, when studying the impact for the different land cover categories, the category *Mosaic tree shrubs, herbaceous cover* shows a decrease in albedo after the fire events. No impact was found for the categories *Broadleaved forest* and *Needleleaved forest*. For the frost drought events, an increase in albedo was found when analysing all sites together. When studying the sites according to the land cover categories, no impact on albedo nor vegetation greenness, as caused by the frost drought events, was found. This result indicates that the impact of frost drought events on albedo is poorly documented. Additionally, further studies should be conducted due to important sources of error caused by uncertainties related to the analysis and limitations of the 30 m Landsat data. This study was focused on the effect of single fire and frost drought events. However, how vegetation will be affected by multiple extreme events of the same and different types interacting, as well as increased frequency in extreme events, must be further studied.

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#### 1. Introduction

With climate change, extreme climate events such as drought, heavy rainfalls, and extreme temperatures, are predicted to be more frequent (AMAP 2017; Bokhorst et al 2009; Box et al. 2019). Extreme climate events can have a large impact on infrastructure, food security, tourism, economics, ecosystems, among others (IPCC 2012). Related to the impacts on ecosystems, extremes can potentially stress and damage vegetation over large areas, resulting in reduced uptake and loss of carbon (Bokhorst et al 2009; Parmentier et al. 2018; Phoenix and Bjerke 2016; Treharne et al. 2019), altered surface reflectance, i.e., albedo (Bright et al. 2013; Huang et al. 2014; Mack et al. 2011; Rocha and Shaver 2011), and shifts in ecosystem composition (AMAP 2017; Bret-Harte 2013; Heim et al. 2020; Parmentier et al. 2018). The knowledge of how extreme climate events affect vegetation is therefore important in multiple fields. Parmentier et al. (2018) states that this knowledge should be considered in research aiming to simulate models and predict future climate. Also, this knowledge is important for the agricultural sector, which is highly vulnerable to climate extremes, due to the dependence of water (IPCC 2012), as well as for grazing animals (Riseth et al. 2011).

This study is focused on two kinds of extreme events, frost drought and fire, especially in the Subarctic region of Sweden and Norway. Studying the distribution of fires in Siberia, Rason et al (2003) concluded that many fires are linked to human activities, and thus human-caused. However, with climate change, that leads to increased maximum air temperatures and drier conditions, the risk of fires is predicted to increase (Box et al. 2019; Jolly et al. 2015). As a result of warmer temperatures, an overall greening trend (increase in plant biomass over time) has been observed in the Arctic over the past 30 years (AMAP 2017; Box et al. 2019; Elmendorf et. al. 2012; Phoenix and Bjerke 2016). However, lately a browning trend (decline in plant biomass over time) caused by winter warming events and frost drought, has been identified as well (Bjerke et al 2014; Parmentier et al. 2018; Phoenix and Bjerke 2016). Winter warming can lead to thaw of the snow cover that insulates and protects vegetation from winter weather conditions. After a few thaw days, ground vegetation gets exposed, and hibernation is being interrupted. When temperature later drops to normal winter conditions, exposed vegetation risks to get freezing damages - a frost drought event occurs (Bokhorst et al. 2008).

Understanding how extreme climate events affect vegetation is important as it gives us a hint on how vegetation will respond to climate change and an increased frequency in extremes (Treharne et al. 2019). Studying different kinds of extremes can provide information on which vegetation is the most vulnerable to, or if different types of events affect vegetation in different ways. The effect of frost drought events is a relatively unexplored topic. A reason for this is that extreme winter warming events are given less attention than the longer summer warming events (Bokhorst et al. 2009). Especially, little is known on how albedo is affected by frost drought, while multiple studies investigate damage ratio, i.e., portion of damaged vegetation, and the Normalized Difference Vegetation Index (NDVI) – a measure of vegetation greenness (Bjerke et al. 2017; Bokhorst et al. 2009; Bokhorst et al. 2012). In general, the impact of fire on vegetation is well documented through studies of NDVI, albedo, and burn severity (Heim et al. 2020; Rocha and Shaver 2011; Sizov et al. 2021). However, studies conducted on fires are often focused on the short-term effects rather than long-term vegetation recovery (Heim et al. 2020; Jones et al 2009).

The aim of this study is to analyse the effect of fire and frost drought events on vegetation greenness and albedo in the Subarctic region of Scandinavia between year 2000 and 2020, to answer the following questions:

- *a) Did the fire and frost drought events have an impact on vegetation greenness and albedo?*
- b) Does the impact on vegetation greenness and albedo of fire and frost drought events differ between land cover categories?

Against the background of earlier studies, a decrease in both albedo and NDVI is expected directly after a fire event (Chambers et al. 2005; French et al. 2016; Rocha and Shaver 2011). However, after a reduction in NDVI, values higher than the values before the event are expected (Heim 2020). For the frost drought events, a reduction in NDVI values is expected. Regarding albedo, an increase in albedo values is expected after frost drought, except for lichen-dominated vegetation where a reduction is expected (Aartsma et al. 2020; Bjerke et al. 2014; Parmentier et al. 2018). Land coverage such as lichens, mosses, and evergreen shrubs are expected to show the greatest vulnerability to frost drought (Bokhorst et al. 2012; Bjerke et al. 2017; Bokhorst et al., 2009). Regarding fire, lichens and mosses are expected to be the most vulnerable (Jandit et al. 2008; Sizov et al. 2021).

#### 2. Theoretical Background

#### 2.1 Impact of Frost Drought on Vegetation Greenness and Albedo

Damage and dieback of vegetation caused by frost drought events can alter NDVI in the Arctics (Bjerke et al. 2014; Bokhorst et al. 2009). Bokhorst et al. (2009) investigated the impacts following an extreme warming event during the winter season 2007/2008. Comparing summer NDVI values the year before and after the event, a reduction of 26% was observed. The study area was the Subarctic region between Abisko (68° 21' N, 18° 49' E) in northern Sweden and Narvik (68° 25' N, 17° 33' E) at the Norwegian coast, covering 1424 km2. Furthermore, Borkhorst et al. (2008) conducted a warming experiment and found that frost drought events can harm plants in terms of development, growth, and reproduction (production of flowers and berries) in the following growing season. Contrastingly, no effect on photosynthesis was found.

Bokhorst et al. (2012) and Parmentier et al. (2018) reported an NDVI recovery time of 2 years after a frost drought event. Bokhorst et al. (2012) suggested that this rapid recovery could be caused by the recovery of non-visibly damaged, dominated species, such as *Betula nana* (dwarf birch) and *Betula pubescens* (downy birch), masking the longer recovery time of *Empetrum nigrum* (crowberry). However, Parmentier et al. (2018) noted that the damage caused by frost drought vary depending on its timing, meaning that events that occur in the absence of sunlight are not as harmful to vegetation, and associated with a relatively short recovery time of the vegetation. The fact that vegetation is capable to recover rapidly from frost drought events means that as long as these events occur infrequently, major vegetation shifts are unlikely (Bokhorst et al. 2012). However, Bokhost et al. (2008; 2012) suggests that if the frequency of these events increases, lasting damages, and shifts in vegetation may occur.

The impacts of frost drought events seem to differ between species. Evergreen dwarf shrubs, and lichens, have been reported to be more vulnerable, in favour of more resilient mosses and sedges (Bjerke 2011; Bjerke et al. 2014; Bjerke et al. 2017; Bokhorst et al., 2009; Bokhorst et al. 2012; Parmentier et al. 2018). Bokhorst et al. (2008) found that the deciduous shrub species *Vaccinium myrtillus* (blueberry) was more vulnerable to warming events than evergreen species. Bokhorst et al. suggest that the greater sensitivity of *Vaccinium myrtillus* could be related to the early bud burst and flowering of this species, and therefore show a greater vulnerability to winter warming events that occur in the beginning of the bud development. Moreover, studying *Calluna vulgaris* (heather) and *Empetrum nigrum*, Bjerke et al. (2017) found that plant damage was correlated with plant height (0-40 cm) with an increased damage ratio for increased plant height.

It is likely that frost drought events can affect other surface properties, such as albedo. In the literature, no previous study on frost drought and albedo was found. However, Parmentier et al. (2018) suggest that a vegetation shift caused by a decline in the more vulnerable species possibly could alter albedo. One study conducted on boreal forest after Bark beetle outbreak found no changes in the summer albedo (Bright et al. 2013), while another study, also in the region of boreal forest, found an increase in albedo during a drought year, and after a decade of recovering, the albedo values were back to prior drought values (Huang et al. 2014). However, lichens have been identified to have a high reflectance and thus a reduction in lichen coverage, potentially caused by frost drought, can cause a decline in albedo (Aartsma et al. 2020).

#### 2.2 Impact of Fire Event on Vegetation Greenness and Albedo

A decrease in both NDVI and albedo has been reported for the years following fire events across the Arctic region (Barrett et al. 2012; Chambers et al. 2005; French et al. 2016; Heim et al. 2020). Heim et al. (2020) studied a fire in Western Siberia at a site with a mix of forested tundra and open treeless shrub tundra and observed a decline in NDVI directly after the fire. Also, after recovering to prior fire event values, Chambers et al. (2005) observed that NDVI continued to increase. A similar trend has been identified by Barrett et al (2012) and Sizov et al (2021). Heim et al. (2020) associated this continuous increase to an observed spread of shrubs after the fire event, and Sizov et al. (2021) suggests that the recovered vegetation is characterised by a larger amount of biomass, likely related to a shift in species. However, studies have identified burn severity as an important regulator for the vegetation composition seen after a fire event (Euskirchen et al. 2009), where the largest change in composition is associated with greater burn severity (Bernhardt et al. 2011; Barrett et al. 2012). Increased burn severity has also been identified to cause greater reduction in NDVI after the fire event (Díaz-Delgado et al. 2003; Rocha and Shaver 2011). For the NDVI recovery time, a range of 2-28 years has been reported (Barrett et al 2012; Heim et al. 2020; Sizov et al 2021).

Studies have reported a decline in albedo after fire events (Chambers et al. 2005; French et al. 2016; Rocha and Shaver 2011). Also, Chambers et al. (2005) found that the decline in albedo was greater at tundra sites compared to a spruce forest. However, Payette and Delwaide (2018) studied lichen woodland in Subarctic North America and suggested an increase in albedo after a fire event due to a shift in vegetation composition from closed-crowned forest stand to lichen dominated vegetation. Alexander et al. (2012) conducted a study on burned sites 20-50 years after the fire events in Alaska and found a shift from evergreen tree species to deciduous tree species, and an associated increase in albedo due to the higher albedo of deciduous forest than evergreen forest.

Like NDVI, the reduction in albedo has been found to be greater with increased burn severity (Rocha and Shaver 2011). Studies report a relatively rapid recovery of albedo for tundra vegetation (French et al. 2016; Rocha and Shaver 2011). However, the albedo recovery time seems to differ in relation to burn severity (French et al 2016; Rocha and Shaver 2011). The change and recovery after a fire event mentioned above is presented in Table 1 for NDVI, and Table 2 for albedo.

Studying tundra fires and the recovery process, lichens have been reported to have a poor recovery, with a recovery time of decades to centuries (Jandit et al. 2008; Sizov et al. 2021; Racine et al 2004). Additionally, Racine et al (2004), also found low recovery of mosses and 20 to 30 years after the fire a large spread of shrubs was observed. However, generally the damage on moss and plant coverage increases with increased burn severity (Rocha and Shaver 2011).

Vegetation type	Change	Recovery (years)	Reference						
Tundra	-40%	8	Heim et al. (2020)						
Tundra	-	3	Barrett et al (2012)						
Tundra and woodlands	-	2-28	Sizov et al (2021)						

Table 1. Change in NDVI and recovery time after fire event.

ruble 2. Change in abedd and recovery time after the event.								
Vegetation type	Change	Recovery (years)	Reference					
Tundra	-75%	-	Chambers et al. (2005)					
Spruce forest	-24%	-	Chambers et al. (2005)					
Tundra	-	3	Rocha and Shaver (2011)					
Tundra	-	4-5	French et al. (2016)					

Table 2. Change in albedo and recovery time after fire event.

#### 3. Material and Method

#### 3.1 Study Area

Events compared were fire and frost drought events that occurred between year 2000 and 2020. The study area was limited to the Subarctic region of Sweden and Norway. The study sites were located in the northern part of Norway and Sweden, and the central and south-western part of Norway. Geographical placement of the study area and distribution of the affected sites and the control sites for the frost drought and fire events are visualised in Fig. 1 and Fig. 2, respectively. The climate conditions at the northern study area (Lofoten, Tromsø Nothern Sweden; Fig. 1 and Fig. 2) are snowy, humid with cool summers and cold winters (Köppen-Geiger climate classification: Dfc), or tundra climate (ET). The study region in south-western part of Norway (Rogaland county; Fig. 1) and central Norway (Municipality of Flatanger and Frøya) has a warm temperate, humid climate with cool summers and cold winters (Cfc) (Kottek et. al. 2006).

#### **3.2 Data Collection**





Fig. 1. Methodological flow of data collection and data analysis.

#### 3.2.1 Frost Drought Events and their Control Sites

All sites used for analysis of the impact of frost drought events were located in Norway. The study sites were given by Bjerke et al. (2017), Parmentier et al. (2018), and Stavanger Aftenblad (2013). Sites based on Bjerke et al. (2017) and Parmentier et al. (2018) were defined by the coordinates used in the publications. The sites studied by Bjerke et al. (2017) are all from the frost drought event that occurred along the Norwegian coast in 2014. This has been documented as a severe event that led to widespread damage on vegetation. The event studied by Parmentier et al. (2018), and also included in the paper by Bjerke et al. (2017), did severely damage shrubs and lead to a reduction in both NDVI and CO<sub>2</sub> uptake. The frost drought event from Stavanger Aftenblad (2013) did severely damaged vegetation, especially Vaccinium myrtillus and juniper trees, in many parts of the county Rogaland in south Norway. No information on intensity of the frost drought events was found. If some coordinates given by the publication were located within 100 m from each other, they were combined into a new point in between the coordinates. Additionally, 7 regions affected by frost drought events were given by Stavanger Aftenblad (2013). Sites derived from these regions were randomly generated within the mentioned region.

Land coverage was defined by the Land Cover Map (version v2.0.7) for year 2015 provided by the European Space Agency (ESA) (2015), with a spatial resolution of 300 m. For each site, the land cover category was defined based on the midpoint of the sites, with the assumption that this applies for the entire site. In case study sites were located within water bodies or urban areas, the point was removed from the data set (1 point was removed because of this). In total 36 points were used. The majority was in north-western Norway, at Lofoten, 1 site was close to Tromsø, and 7 sites were in the south-western Norway, in Rogaland county.

Snow covered control sites, paired with the affected sites, were selected based on snow depth data obtained from the Norwegian meteorological stations Botnhamn, Kvitfossen i Vågan, Kongsmarka, Gausvik, Breivikeidet, and Bangdalen (Norsk klimaservicesenter n.d.). These stations indicated a steady snow coverage the entire winter season for when the specific frost drought event occurred. Using the meteorological stations as starting points, control sites with similar elevation and land coverage as the frost damaged sites were selected. Topography data was derived from the European Digital Elevation Model (version 1.1), with a spatial resolution of 25 m and provided by the European Environment Agency (Copernicus programme) (2016). The land cover was defined by the Land Cover Map from ESA (2015). The distance between the study points and the control point were in the range 20-80 km, except for the sites in south west of Norway. These sites were located ~800 km from the meteorological station (Fig. 2). This large distance is due to the lack of station data, and that other stations further south did not have steady snow coverage the entire winter that the specific event occurred. The large distance introduces a risk that the sites do not have the same climate conditions. However, both regions have a warm

temperate, humid climate with cool summers and cold winters (Cfc) according to the Köppen-Geiger climate classification (Kottek et. al. 2006). For more details on affected sites and control sites, see Appendix I.



Fig 2. Spatial distribution for the studied frost drought events (study point: orange and control point: green) across Scandinavia.

#### 3.2.2 Fire Events and their Control Sites

Sites selected to study the fire events were based on data collected by the Swedish rescue service Myndigheten för samhällsskydd och beredskap (MSB), Hamarøy municipality, and Direktoratet for Samfunnssikkerhet og Beredskap (DSB). From the MSB data set (2021), 29 fire affected sites were used, all located in the municipalities Jokkmokk, Kiruna or Gällivare, and defined by coordinates and date. Hamarøy municipality (2020) provided information on 7 affected sites, also defined by coordinates and date. Lastly, DSB (2014) reports 2 fire events. Study coordinates for these 2 events were randomly selected within the burned area mapped by Brakstad (2014) and Svarstad (2014). No information was found on damage area, intensity of the fire, and burn severity of the sites given by Hamarøy municipality (2020). The burned area of the sites given by MSB (2021) vary largely. 16 of the sites have an area smaller than 1000 m<sup>2</sup>, 12 sites have an area between 1000 and 30000 m<sup>2</sup>, and 4 sites and area greater than 30000 m<sup>2</sup>. The intensity and burn severity of these fire

events is unknown. The 2 fire events given by DSB (2014) covered a large area, both approximately 8km<sup>2</sup>. These fires are described as great and prolonged, but the burn severity is unknown, as well as the intensity of the fire.

Like for the frost drought events, the land cover category was defined based on the midpoint of the sites, and in case an affected site was located within the land cover classes water or urban (ESA 2015), the site was removed from the data set (3 sites were excluded because of this). In total 38 points were selected for the fire events.

To define the control sites, paired with affected sites, Red Green Blue (RGB) colour composites based on Landsat 5, 7, and 8 satellite data, covering the affected sites and their surrounding area within a radius of 100 km, were used. The burning year and the years before and after the fire events were studied visually, with the aim to define areas of unaffected vegetation. This information, together with information on land coverage from the Land Coverage Map by ESA (2015), elevation data from the European Digital Elevation Model by EEA (2016), and the location of the affected sites, was used to select control sites. For more details on the affected sites and control sites see Appendix II.



Fig 3. Spatial distribution for the studied fire events (study point: orange and control point: green) across Scandinavia.

#### 3.3 Remote Sensing Data: Collection and Pre-processing

The period 2000 to 2020 was selected because of the significant improvements in the quality of satellite data that were made associated with the launch of Landsat 7 (Storey et al. 2000). Data used for this study were mainly collected by the satellites Landsat 7 and 8. In case Landsat 7 or 8 data were not available, Landsat 5 was used. Due to seasonal dynamics in NDVI and albedo only data from the period July the 1<sup>st</sup> to August the 31<sup>st</sup> was used. This specific period was defined to assure that albedo values were collected in snow-free conditions and NDVI values when the vegetation is at peak (Barrett et al. 2012; French et al. 2016; Heim et al 2020). The data set was downloaded from Google Earth Engine (GEE). GEE is a platform that allows the users to run geospatial analysis on the infrastructure of Google through the code editor. The code editor is a web-based integrated development environment for writing and running scripts using either JavaScript or Python (Google Earth Engine n.d.). For each year, a mosaic of available images for the period July the 1<sup>st</sup> to August the 31<sup>st</sup>, was created using the median pixel of the period. The images used had a spatial resolution of 30 m. However, if the fire or frost drought event occurred during the growing season, images collected before the event were removed from the mosaic. The images have been radiometrically corrected for atmospheric noise (NASA) and clouds have been masked out using the embedded pixel quality bitmask. Occasionally, the embedded quality mask (CFMask: Foga et al. 2017) did not detect obvious cloud coverage (Appendix III). Hence, RGB colour compositions were visually inspected, and sites with pixels identified as affected were manually excluded from the data set (39 site-years of the fire events and 45 site-years of the frost drought events were excluded because of this).

Based on the Landsat satellite data, NDVI and albedo were calculated. NDVI is a measure of vegetation greenness or amount of photosynthetically active vegetation calculated as:

$$NDVI = (NIR-RED)/(NIR+RED),$$

where NIR is the near-infrared reflectance (0.7–1.1  $\mu$ m) and RED is the visible red (0.6–0.7  $\mu$ m) (Rouse et al. 1974; Tucker 1979).

Albedo is the ratio reflected to incoming solar radiation. It is computed by the integrated amount of shortwave solar reflectance across the spectral range 0.35-2.5  $\mu$ m (Chuevieco 2016; Tsuyuzaki et al. 2009). From Landsat images, converted to top of atmosphere reflectance, albedo can be computed as:

$$A_{short} = ((0.356*B1) + (0.130*B3) + (0.373*B4) + (0.085*B5) + (0.072*B7) - 0.018) / (0.356 + 0.130 + 0.373 + 0.085 + 0.072),$$

where B1, B3, B4, B5 and B7 represent Landsat band 1, 3, 4, 5 and 7 respectively (Liang 2001; Smith 2010).

For the affected and control sites a window of 100 m radius around each study point was created. Water bodies and urban areas were masked out and the windows were then used to extract albedo and NDVI.

#### 3.4 Data Analysis

# **3.4.1** a) Did fire and frost drought events have an impact on vegetation greenness and albedo?

For each site and each year, mean NDVI and albedo of all pixels in the created window (100 m radius) was calculated. To quantify the impact of the frost drought and fire events on vegetation greenness and albedo, mean NDVI and albedo of the control sites were subtracted from the values of the paired affected sites. This was done for each site and each year of the full time-series. From now on this difference between the affected and control sites is referred to as  $\Delta$ albedo and  $\Delta$ NDVI. The data of all study sites were then temporally aligned by defining the season following the event as year 0. If the event occurred during the growing season, the current year was defined as year 0.

Thereafter, for each site, the mean and the ordinary least square linear regression of albedo and NDVI of the affected site and the control site, as well as  $\Delta$ albedo and  $\Delta$ NDVI, for the 5 years before and after the event, was calculated. All analysis and statistical tests were limited to the 5 years before and after the events. To reduce the risk that long-term variations that naturally occur in ecosystems over time could obscure the effect of the studied event. The slope coefficient of the linear trend of  $\Delta$ NDVI and  $\Delta$ albedo, mean  $\Delta$ NDVI and mean  $\Delta$ albedo after the event were compared against the values before the event using paired sampled t-tests. The paired sampled t-test is used to determine differences between the means of two samples, for instance, when there are before and after observations (Bernhardt et al. 2011; Saxena et al. 2020)). Finally, the mean and the standard deviation (SD) of the mean and the linear trend of  $\Delta$ NDVI and  $\Delta$ albedo for all sites the years before and after the event, was calculated.

# **3.4.2** b) Does the impact on vegetation greenness and albedo of fire and frost drought events differ between land cover categories?

To study the differences between land cover categories, the classification by the Land Coverage Map by ESA (2015) was used. The original classes of the map were modified in order to allow comparisons between the frost drought events and the fire events and make the categories representable because of lack of data in some classes.

Table 3 presents the categories used and the corresponding classes defined by ESA (2015).

To determine the variations between the land cover categories, the analysis conducted for research question a) was repeated once for each of the land cover categories. Separated according to the land cover categories (Table 3) the mean and the slope coefficient of the linear trend of albedo and NDVI of the affected site and the control site as well as  $\Delta$ NDVI and  $\Delta$ albedo before and after the event was compared, using the paired sampled t-test. Finally, the mean and SD of the mean and the slope coefficient of the linear trend of  $\Delta$ NDVI and  $\Delta$ albedo for each of the land cover categories the years before and after the event, was calculated.

Table 3. Land cover categories used and derived from the Land Coverage Map by ESA (2015).

Land coverage defined by Land Coverage Map by ESA (2015)	Categories used
Tree cover, broadleaved, deciduous, closed to open (>15%)	Broadleaved forest
Tree cover, needleleaved, evergreen, closed to open (>15%)	Needleleaved forest
Sparse vegetation (tree, shrub, herbaceous cover) (<15%)	
Sparse shrub (<15%)	
Mosaic tree and shrub (>50%) / herbaceous cover (<50%)	
Mosaic herbaceous cover (>50%) / tree and shrub (<50%)	Mosaic tree shruhs
Cropland, rainfed	herbaceous cover
Mosaic natural vegetation (tree, shrub, herbaceous cover) (>50%) / cropland (<50%)	
Shrub or herbaceous cover, flooded, fresh/saline/brakish water	

#### 4. Results

#### 4.1 a) Did fire and frost drought events have an impact on NDVI and albedo?

#### 4.1.1 Analysis of NDVI and Albedo at Affected and Control sites

The paired sampled t-test conducted to test the difference between the slope coefficient and the mean of mean albedo and NDVI of the affected and control sites, before and after the events, show no difference for the fire events and albedo, neither for the affected sites nor the control sites. For NDVI a difference was found for the control site both in the mean and the trend, indicating an increase in the NDVI values (Fig. 5, Table 4, and Table 5). For the frost drought events, an increase in NDVI was shown both when testing the difference in the mean and the slope coefficient of the linear trend at the affected sites and the control sites. For albedo, a difference indicating an increase in albedo was found for the affected site when testing the difference in the slope coefficient, also indicating an increase in albedo values (Fig. 4, Table 4, and Table 5).



Fig. 4. Time series of albedo and NDVI mean values of the affected and control sites for the frost drought events.



Fig. 5. Time series of albedo and NDVI mean values of the affected and control sites for the fire events.

Table 4. Results of the paired sampled t-test testing for a difference in mean albedo and NDVI of the affected and control sites for fire respective frost drought before and after the event. The mean values are the mean of mean NDVI and albedo of all sites for the fire respective frost drought events and SD is the standard deviation.

				Paired	samples			
			Me	an	t-t	t-test		
			Before	After	t-value	p-value	size	
	Albada	Affected	$0.123\pm0.028$	$0.124\pm0.030$	-0.271	0.788	38	
Fire	Albedo	Control	$0.128\pm0.024$	$0.127\pm0.023$	0.607	0.548	38	
гпе	NDVI	Affected	$0.693 \pm 0.100$	$0.697\pm0.117$	-0.578	0.566	38	
	NDVI	Control	$0.733\pm0.073$	$0.746\pm0.067$	-2.559	0.015	38	
	Albada	Affected	$0.147\pm0.032$	$0.152\pm0.028$	-2.547	0.015	36	
Frost drought	Albedo	Control	$0.141\pm0.025$	$0.143 \pm 0.025$	-1.358	0.183	36	
	NDVI	Affected	$0.701\pm0.076$	$0.733 \pm 0.074$	-6.029	< 0.001	36	
		Control	$0.725\pm0.072$	$0.752\pm0.067$	-5.040	< 0.001	36	

Fire

Table 5. Results of the paired sampled t-test testing for difference in slope coefficient of the linear trend of mean albedo and NDVI of the affected and control sites for fire respective frost drought before and after the event. The mean values are the mean of the slope coefficient of mean NDVI and albedo of all sites for the fire respective frost drought events and SD is the standard deviation.

				Paired s	samples			
			Me	an	t-t	t-test		
			Before	After	t-value	p-value	size	
A 11	Albada	Affected	$\textbf{-0.001} \pm 0.007$	$-0.001 \pm 0.006$	0.115	0.909	38	
Fina	Albedo	Control	$0.000\pm0.005$	$0.002\pm0.011$	-0.935	0.356	38	
Fire	NDVI	Affected	$0.003\pm0.017$	$0.014\pm0.014$	-2.650	0.012	38	
	NDVI	Control	$0.003\pm0.011$	$0.013 \pm 0.021$	-2.392	0.022	38	
	Albada	Affected	$\textbf{-0.002} \pm 0.014$	$0.001\pm0.003$	-1.301	0.202	36	
Frost	Albedo	Control	$\textbf{-0.001} \pm 0.005$	$0.002\pm0.003$	-3.558	0.001	36	
drought	NDVI	Affected	$0.004\pm0.020$	$0.013\pm0.010$	-2.678	0.011	36	
	NUVI	Control	$0.000\pm0.010$	$0.010\pm0.011$	-3.640	0.001	36	

#### 4.1.2 Analysis of $\Delta$ NDVI and $\Delta$ albedo

Analysing the difference between the slope coefficient of the linear trend and the mean of  $\Delta$ albedo and  $\Delta$ NDVI, before and after the events, using the paired sampled t-test, resulted in no difference for the fire events (Fig. 7, Table 6, and Table 7). For the frost drought events, an increase in  $\Delta$ albedo was found, but no difference was found for  $\Delta$ NDVI. No difference in slope coefficient before and after the event was found, neither for  $\Delta$ albedo, nor  $\Delta$ NDVI (Fig. 6, Table 6, and Table 7).



*Fig.* 6. *Time series of a*)  $\Delta$ *albedo; and b*)  $\Delta$ *NDVI for the frost drought events.* 



Fig. 7. Time series of a)  $\triangle$  albedo; and b)  $\triangle$ NDVI for the fire events.

Table 6. Results of the paired sampled t-test testing for a difference in mean  $\Delta$ albedo and  $\Delta$ NDVI of fire respective frost drought before and after the event. The mean values are the mean of mean  $\Delta$ NDVI and  $\Delta$ albedo of all sites for the fire respective frost drought events and SD is the standard deviation.

		Mean	± SD	Paired san	nples t-test	Sample
		Before	After	t-value	p-value	size
Fine	Albedo	$-0.004 \pm 0.027$	$-0.003 \pm 0.028$	-0.146	0.885	38
гпе	NDVI	$-0.040 \pm 0.096$	$-0.051 \pm 0.098$	1.493	0.144	38
Frost	Albedo	$0.004\pm0.042$	$0.009\pm0.038$	-2.038	0.049	36
drought	NDVI	$-0.023 \pm 0.086$	$-0.019 \pm 0.093$	-0.603	0.550	36

Table 7. Results of the paired sampled t-test testing for a difference in the slope coefficient of the linear trend of  $\Delta$ albedo and  $\Delta$ NDVI, before and after the event. The mean values are the mean of the trend of  $\Delta$ NDVI and  $\Delta$ albedo of all sites for the fire respective frost drought events and SD is the standard deviation.

		Mean	Paired san	Sample		
		Before	After	t-value	p-value	size
Eine	Albedo	$-0.001 \pm 0.006$	$-0.001 \pm 0.005$	0.410	0.684	38
Fire	NDVI	$-0.002 \pm 0.018$	$0.003\pm0.016$	-1.104	0.277	38
Frost	Albedo	$0.000\pm0.014$	$0.000\pm0.005$	0.186	0.853	36
drought	NDVI	$0.000\pm0.026$	$0.004\pm0.014$	-0.624	0.537	36

**4.2 b)** Does the impact on vegetation greenness and albedo of fire and frost drought events differ between land cover categories?

# **4.2.1** Analysis of NDVI and Albedo at Affected and Control sites separated in Land Cover Categories

The paired sampled t-test conducted for each of the land cover categories, show an increase in NDVI at the affected sites after frost drought for the land cover category *Needleleaved forest*, both when analysing the mean and the slope coefficient (Fig. 10, Table 8, Table 9). The land cover category *Mosaic tree, shrubs, herbaceous cover*, a difference was found only for NDVI when analysing the difference in mean values, indicating an increase in the mean of mean NDVI after frost drought. This was shown

both at the affected and control sites. Regarding the land cover category *Broadleaved forest*, an increase in mean NDVI after frost drought, both for the affected sites and the control sites, was found. This difference was reflected also in the slope coefficient for the control sites, indicating a shift from negative to positive mean slope coefficient, but no difference was found for the affected sites (Fig. 10, Table 8, Table 9). For albedo, a difference was found only for the land cover category *Broadleaved forest* when analysing the slope coefficient, with a shift from negative to positive mean slope coefficient, both for the affected sites and the control sites (Fig. 11, Table 8, Table 9).

Regarding the fire events, no impact was found when analysing the difference in mean values, neither for albedo nor NDVI (Fig. 8, Fig. 9, Table 8). When analysing the slope coefficient of NDVI before and after the events an increase in the mean slope coefficient was found for the land cover category *Mosaic tree, shrubs, herbaceous cover* at the control sites. Also, a difference in NDVI, with a shift from a negative to a positive mean slope coefficient, was found at the affected sites for the land cover category *Needleleaved forest* (Fig. 8, Table 9).



Fig. 8. NDVI for the fire events at the affected and control sites for the land cover categories a) Mosaic tree, shrubs, herbaceous cover, b) Broadleaved forest, and c) Needleleaved forest.



Fig. 9. Albedo for the fire events at the affected and control sites for the land cover categories a) Mosaic tree, shrubs, herbaceous cover, b) Broadleaved forest, and c) Needleleaved forest.



Fig. 10. NDVI for the frost drought events at the affected and control sites for the land cover categories a) Mosaic tree, shrubs, herbaceous cover, b) Broadleaved forest, and c) Needleleaved forest.



Fig. 11. Albedo for the frost drought events at the affected and control sites for the land cover categories a) Mosaic tree, shrubs, herbaceous cover, b) Broadleaved forest, and c) Needleleaved forest.

#### Frost Drought: Albedo

Table 8: Results of the paired sampled t-test testing the difference in the mean albedo and NDVI of fire respective frost drought before and after the event, at the affected sites and the control sites. The mean values are the mean of mean NDVI and albedo of all sites for each of the land cover categories and for the affected and control sites of the fire respective frost drought events, and SD is the standard deviation.

				Paired s	Paired samples		
			Mean			t-test	
			Before	After	t-value	p-value	size
		1	Mosaic tree, shrubs	, herbaceous cove	r		
	Albado	Affected	$0.122\pm0.018$	$0.112\pm0.016$	2.732	0.052	7
Fire	Albedo	Control	$0.126\pm0.014$	$0.134\pm0.011$	-1.876	0.134	7
THE	NDVI	Affected	$0.641\pm0.103$	$0.600\pm0.015$	1.292	0.266	7
	IND VI	Control	$0.645\pm0.078$	$0.657\pm0.074$	-1.366	0.244	7
	Albada	Affected	$0.149 \pm 0.033$	$0.155\pm0.030$	-1.841	0.081	19
Frost	Albedo	Control	$0.143 \pm 0.023$	$0.144 \pm 0.024$	-0.612	0.548	19
drought	NDVI	Affected	$0.685\pm0.078$	$0.720\pm0.072$	-4.510	< 0.001	19
	NDVI	Control	$0.699\pm0.074$	$0.734\pm0.065$	-5.240	< 0.001	19
			Broadleav	red forest			
	A 111 -	Affected	$0.139 \pm 0.030$	$0.146 \pm 0.029$	-1.873	0.088	12
Fine	Albedo	Control	$0.147 \pm 0.024$	$0.141 \pm 0.019$	1.268	0.231	12
Fire	NDVI	Affected	$0.738 \pm 0.109$	$0.744 \pm 0.120$	-0.575	0.577	12
		Control	$0.773\pm0.073$	$0.788 \pm 0.055$	-1.680	0.121	12
	Albada	Affected	$0.156\pm0.020$	$0.158 \pm 0.019$	-0.809	0.437	6
Frost	Albedo	Control	$0.143 \pm 0.024$	$0.146\pm0.023$	-1.266	0.234	6
drought	NDVI	Affected	$0.738 \pm 0.031$	$0.760\pm0.037$	-2.938	0.015	6
	NDVI	Control	$0.761\pm0.060$	$0.791 \pm 0.057$	-4.963	0.001	6
			Needleleav	ved forest			
	A 11 1 .	Affected	$0.110\pm0.020$	$0.111\pm0.023$	-0.283	0.780	19
Fine	Albedo	Control	$0.114\pm0.018$	$0.113 \pm 0.020$	0.670	0.511	19
Fire	NIDVI	Affected	$0.678\pm0.088$	$0.693 \pm 0.090$	-2.000	0.061	19
	NDVI	Control	$0.742\pm0.034$	$0.753 \pm 0.032$	-1.454	0.163	19
	A 11 1 -	Affected	$0.125\pm0.033$	$0.134 \pm 0.029$	-2.195	0.080	11
Frost	Albedo	Control	$0.132\pm0.029$	$0.135\pm0.029$	-0.845	0.437	11
drought	NIDVI	Affected	$0.681 \pm 0.092$	$0.715\pm0.104$	-2.578	0.050	11
		Control	$0.751\pm0.034$	$0.744 \pm 0.060$	0.429	0.685	11

Table 9: Results of the paired sampled t-test testing the difference in the slope coefficient of albedo and NDVI of fire respective frost drought before and after the event, at the affected sites and the control sites. The mean values are the mean of the slope coefficient of NDVI and albedo of all sites for each of the land cover categories and for the affected and control sites of the fire respective frost drought events, and SD is the standard deviation.

				Paired s			
			Me	an	t-te	est	Sample
			Before	After	t-value	p-value	size
		]	Mosaic tree, shrubs	, herbaceous cove	er		
	Albedo	Affected	$-0.003 \pm 0.003$	$-0.001 \pm 0.006$	-0.472	0.662	7
Fire	Albedo	Control	$0.002\pm0.009$	$0.002\pm0.003$	0.034	0.975	7
THE	NDVI	Affected	$0.008 \pm 0.010$	$0.022\pm0.013$	-1.978	0.119	7
	IND VI	Control	$0.002\pm0.010$	$0.015\pm0.007$	-3.025	0.039	7
	Albada	Affected	$0.003 \pm 0.007$	$0.001\pm0.003$	0.800	0.435	19
Frost	Albedo	Control	$0.000\pm0.005$	$0.001\pm0.004$	-1.237	0.231	19
drought	NDVI	Affected	$0.008 \pm 0.017$	$0.012\pm0.011$	-1.195	0.249	19
	NDVI	Control	$0.002\pm0.010$	$0.006\pm0.010$	-1.500	0.150	19
			Broadleav	red forest			
	Albada	Affected	$0.003\pm0.009$	$0.002\pm0.004$	0.315	0.760	12
<b>F</b> 's s	Albedo	Control	$0.002\pm0.003$	$0.007\pm0.018$	-0.980	0.350	12
Fire	NDVI	Affected	$0.008 \pm 0.011$	$0.007\pm0.009$	0.388	0.707	12
		Control	$0.002\pm0.012$	$0.016\pm0.029$	-1.500	0.164	12
	Albada	Affected	$-0.001 \pm 0.003$	$0.002\pm0.003$	-2.256	0.048	6
Frost	Albedo	Control	$-0.003 \pm 0.003$	$0.002\pm0.002$	-4.377	0.001	6
drought	NDVI	Affected	$0.008 \pm 0.015$	$0.013 \pm 0.008$	-1.296	0.224	6
	NDVI	Control	$\textbf{-0.002} \pm 0.008$	$0.011\pm0.008$	-3.304	0.008	6
			Needleleav	ved forest			
	Alleado	Affected	$-0.002 \pm 0.006$	$-0.002 \pm 0.004$	-0.297	0.770	19
<b>F</b> 's s	Albedo	Control	$-0.001 \pm 0.004$	$-0.001 \pm 0.004$	0.118	0.907	19
Fire	NIDVI	Affected	$-0.003 \pm 0.020$	$0.015\pm0.011$	-2.777	0.012	19
	NDVI	Control	$0.005\pm0.011$	$0.012\pm0.018$	-1.317	0.204	19
	A 11 1 -	Affected	$-0.016 \pm 0.026$	$0.001\pm0.004$	-1.631	0.164	11
Frost	Albedo	Control	$-0.004 \pm 0.004$	$0.002\pm0.002$	-2.106	0.089	11
drought	NDVI	Affected	$\textbf{-0.016} \pm 0.020$	$0.015\pm0.007$	-2.645	0.046	11
		Control	$-0.005 \pm 0.012$	$0.019\pm0.011$	-2.476	0.056	11

#### 4.2.2 Analysis of ΔNDVI and Δalbedo separated in Land Cover Categories

It is apparent that none of the land cover categories show a difference in the slope coefficient of the linear trend of  $\Delta$ albedo and  $\Delta$ NDVI for the fire or frost drought events (Fig. 12, Fig. 13, Table 11). For the land cover categories, *Mosaic tree, shrubs, herbaceous cover* did show a difference comparing mean  $\Delta$ albedo before and after the fire drought events, with a decrease in the mean of mean  $\Delta$ albedo from -0.002 ± 0.020 to -0.023 ± 0.012. None of the other two land cover categories shown a difference in

mean  $\Delta$ albedo. Also, no land cover category shows a difference in mean  $\Delta$ NDVI comparing before and after event values (Fig. 12, Fig. 13, Table 10).



Fig. 12. ∆albedo and ∆NDVI of frost drought events for the land cover categories
a) Mosaic tree, shrubs, herbaceous cover, b) Broadleaved forest, and
c) Needleleaved forest.



Fig. 13.  $\triangle$ albedo and  $\triangle$ NDVI of frost drought events for the land cover categories a) Mosaic tree, shrubs, herbaceous cover, b) Broadleaved forest, and c) Needleleaved forest.

Table 10. Results of the paired sampled t-test testing the difference in the mean  $\Delta$ albedo and  $\Delta$ NDVI of fire respective frost drought before and after the event. The mean values are the mean of mean  $\Delta$ NDVI and  $\Delta$ albedo of all sites for each of the land cover categories and for the fire respective frost drought events and SD is the standard deviation.

		Mean	± SD	Paired sar	Sample	
		Before	After	t-value	p-value	size
		Mosaic tre	ee, shrubs, herbace	eous cover		
Eiro	Albedo	$-0.002 \pm 0.020$	$-0.023 \pm 0.012$	3.305	0.030	7
гпе	NDVI	$-0.011 \pm 0.085$	$-0.057 \pm 0.119$	1.530	0.201	7
Frost	Albedo	$0.005\pm0.037$	$0.011\pm0.035$	-1.630	0.120	19
drought	NDVI	$\textbf{-0.013} \pm 0.083$	$-0.014 \pm 0.088$	0.079	0.938	19
			Broadleaved fores	t		
Eine	Albedo	$-0.006 \pm 0.034$	$0.003\pm0.300$	-2.131	0.056	12
гпе	NDVI	$-0.034 \pm 0.083$	$-0.046 \pm 0.096$	1.233	0.243	12
Frost	Albedo	$0.011\pm0.039$	$0.013 \pm 0.033$	-0.719	0.488	6
drought	NDVI	$-0.021 \pm 0.075$	$-0.030 \pm 0.072$	0.822	0.430	6
		1	Needleleaved fores	st		
Eine	Albedo	$-0.004 \pm 0.022$	$-0.002 \pm 0.026$	-0.843	0.410	19
Fire	NDVI	$-0.063 \pm 0.101$	$-0.063 \pm 0.091$	-0.037	0.971	19
Frost	Albedo	$-0.008 \pm 0.054$	$0.000\pm0.049$	-0.953	0.384	11
drought	NDVI	$\textbf{-0.068} \pm 0.098$	$-0.031 \pm 0.134$	-1.984	0.104	11

Table 11. Results of the paired sampled t-test testing the difference in the slope coefficient of the linear trend of  $\Delta$ albedo and  $\Delta$ NDVI, before and after the event. The mean values are the mean of the slope coefficient of  $\Delta$ NDVI and  $\Delta$ albedo of all sites for each of the land cover categories and for the fire respective frost drought events and SD is the standard deviation.

		Mean	Paired san	Sample		
		Before	After	t-value	p-value	size
		Mosaic tre	ee, shrubs, herbace	eous cover		
Fire	Albedo	$-0.004 \pm 0.006$	$-0.002 \pm 0.005$	-0.326	0.761	7
гпе	NDVI	$0.004\pm0.015$	$0.008 \pm 0.011$	-0.325	0.762	7
Frost	Albedo	$0.003\pm0.010$	$0.000\pm0.004$	1.728	0.102	19
drought	NDVI	$-0.003 \pm 0.023$	$0.006\pm0.013$	-1.162	0.261	19
		]	Broadleaved forest	t		
Fina	Albedo	$0.000\pm0.007$	$0.001 \pm 0.004$	-0.055	0.957	12
гпе	NDVI	$0.006\pm0.013$	$-0.002 \pm 0.011$	1.737	0.116	12
Frost	Albedo	$0.001\pm0.005$	$-0.001 \pm 0.004$	1.313	0.218	6
drought	NDVI	$0.010\pm0.014$	$0.001\pm0.009$	1.780	0.105	6
		١	Needleleaved fores	st		
Fina	Albedo	$-0.001 \pm 0.006$	$-0.001 \pm 0.005$	0.002	0.999	19
гпе	NDVI	$-0.009 \pm 0.019$	$0.003 \pm 0.018$	-1.678	0.111	19
Frost	Albedo	$-0.011 \pm 0.024$	$0.003\pm0.008$	-1.072	0.333	11
drought	NDVI	$-0.013 \pm 0.039$	$-0.001 \pm 0.019$	-0.492	0.644	11

#### 5. Discussion

# 5.1 a) Did fire and frost drought events have an impact on vegetation greenness and albedo?

In this work, a before-after impact approach was applied to study the impact of fire and frost drought events on vegetation greenness and albedo, using Landsat satellite data. The statistical analysis of all sites together for the fire events shows no difference in mean  $\Delta$ albedo and  $\Delta$ NDVI or the slope coefficient of the linear trend of  $\Delta$ albedo and  $\Delta$ NDVI, before and after the events (Fig. 7, Table 6, Table 7). Also, there was no difference in mean  $\Delta$ NDVI or the trend of  $\Delta$ NDVI following frost drought events (Fig. 6, Table 6, Table 7). For mean  $\Delta$ albedo, however, there was an increase in the mean from 0.004 ± 0.042 to 0.009 ± 0.038 (Table 6). Despite this, no difference was found for the slope coefficient.

Analysing the difference in the mean and slope coefficient before and after the fire events for the affected and control sites separately, a difference was shown in the slope coefficient of NDVI both at the affected site and the control site. For mean NDVI a difference was found only at the control site (Table 4 and Table 5). For the frost drought events a difference was shown in both the mean and the slope coefficient of NDVI, both at the affected and control site. For albedo, a difference was shown at the affected site when analysing the difference in mean, while a difference in the control site was found when analysing the difference in the slope coefficient (Table 4 and Table 5). However, the impacts seen at the affected sites could be caused by other parameters than the fire and frost drought events and therefore  $\Delta$ NDVI and  $\Delta$ albedo was used. Using  $\Delta$ NDVI and  $\Delta$ albedo, on the other hand, introduces a risk that variability in the control sites could lead to that the values of the affected sites are less separable from values of the control sites, with the consequence that the impact of the fire and frost drought events would not be shown in  $\Delta$ albedo and  $\Delta$ NDVI.

The results of this report contradict the results of earlier studies on fire and frost drought events that indicate a reduction in NDVI after the events (Bokhorst et al. 2009; Bjerke et al. 2014; Bjerke et al. 2017; Heim et al. 2020). Also, earlier studies conducted on fire events found that after a reduction, NDVI tends to reach higher values, presumably caused by a shift in species towards a land coverage characterised by a larger amount of green biomass (Barrett et al; 2012; Heim et al. 2020; Sizov et al 2021). Such a trend was not identified in this study. However, given that Bernhardt et al. (2011) and Barrett et al. (2012) suggest that the largest change in composition is associated with burn severity, and because greater burn severity causes greater reduction in NDVI directly after the fire (Díaz-Delgado et al. 2003), it could be that the severity of the fire events studied are of low character, and thus does not show an effect in NDVI.

For the frost drought event an explanation for the conflicting results is the different methods used. Sites in the region of Troms and Lofoten, obtained by Bjerke et al. (2017), were examined during fieldwork and the extent of damages and NDVI values were measured and observed. In this work, the same sites were studied but NDVI and albedo values were obtained using satellite-based remote sensing. Thus, it could be that these effects are difficult to detect with the 30 m spatial resolution of Landsat, even though they were clearly visible in the field. It could also be that the effect is being obscure when using a mosaic of the median. Here the median was used to extract Landsat data for each growing season (July the 1<sup>st</sup> to August the 31<sup>st</sup>). The median can in some cases be a better measure of the "normal" values than the average. This as it does not allow extreme values to impact the results, which is especially beneficial when the sample size is small. However, Bokhorst et al. (2009) did use remote sensing to study the impact of frost drought on satellite-based NDVI and identified a reduction of 26%. The study area was in the same region as of this work but investigated a winter warming event that occurred during the season 2007/2008. Instead of having multiple study sites, Bokhorst et al. (2009) covered a large area (1424 km<sup>2</sup>) and observed how the NDVI values over the entire area were altered. It could be that the frost drought events occur over a large area with a spatially irregular distribution, making the method used by Bokhorst et al. (2009) more suitable as it accounts for this irregularity.

For albedo, earlier studies report a decrease in albedo after a fire event (Chambers et al. 2005; Rocha and Shaver 2011). In the literature, no earlier studies conducted on albedo and frost drought were found. Parmentier et al. (2018) did not specifically study the albedo but suggest that a vegetation shift followed by a frost drought event could possibly alter it. The results of this report show an increase when comparing mean  $\Delta$ albedo before and after the frost drought event (Table 6). This is an indication that frost drought do have an impact on vegetation in terms of albedo. However, this needs to be further investigated.

# **5.2 b)** Does the impact on vegetation greenness and albedo of fire and frost drought events differ between land cover categories?

The second research question aimed to investigate the impact frost drought and fire events have on different land cover categories. No difference was observed when analysing the mean and the slope coefficient of the linear trend of  $\Delta$ albedo and  $\Delta$ NDVI for the frost drought events for each of the land cover categories (Fig. 12, Fig. 13, Table 10, Table 11). This result conflict with the results of earlier studies. Earlier studies have pointed out shrubs and lichens as especially vulnerable to frost drought (Bjerke 2011; Bjerke et al. 2014; Bjerke et al. 2017; Bokhorst et al., 2009; Bokhorst et al. 2012), while mosses and sedges have been identified as less vulnerable (Parmentier et al. 2018). One reason for the obtained result is that all of these belong to the land cover category *Mosaic tree, shrubs, herbaceous cover*. Thus, because the land coverage is a mosaic of different vegetation types where the different species are

not equally vulnerable to the extreme event, it could be difficult to detect the effect of the frost drought event using these broad land cover categories. For instance, Bokhorst et al. (2012) suggested that the rapid (within 2 years) recovery of NDVI after a frost drought event could be due to non-visibly damaged, dominated species, such *Betula nana* and *B. pubescens*, masking the longer recovery time of *Empetrum nigrum*.

The observed reduction in mean  $\Delta$ albedo after fire for the land cover category *Mosaic tree, shrubs, herbaceous cover*, is partly in line with the work by Chambers et al. (2005) who report a larger reduction for tundra sites than for spruce forest (-75% versus -24%). Also, lichens, which belong to the land cover category *Mosaic tree, shrubs, herbaceous cover*, have been identified as especially vulnerable with long recovery rates (Jandit et al. 2008; Sizov et al. 2021; Racine et al 2004). Earlier studies have seen altered albedo both at open tundra and in forest ecosystems (Chambers 2005; French et al. 2016; Heim et al. 2020; Payette and Delwaide 2018), while the result of this study show no impact on albedo for the categories *Broadleaved forest* and *Needleleaved forest*. However, burn severity seems to generally be an important factor for the effect of fire on albedo (French et al 2016; Rocha and Shaver 2011).

#### 5.3 Limitations and Source of Error

An important limitation of the study is the lack of knowledge on the physical properties of the affected sites and the control sites, such as more detailed information on vegetation coverage, area of affected vegetation, intensity of event, and damage ratio. Even though the extent of affected area was unknown for some of the frost drought and the fire events, an area with radius 100 m was created for the affected and control sites, assuming this to be an area suitable to collect the signal of the effect of the fire or frost drought event. However, the known area of the fire events studied were of a large variety, and having an affected area that is smaller than the defined study site could lead to that the effect of the event is being masked, which is a source of error caused by this assumption. The intensity of the events and the damage ratio of the vegetation is unknown, which means that the effect of the events could be obscures so that it is not detectable in the spectral signal. For instance, non-affected vegetation could possibly expand and thus compensate for the negative effect the fire or frost drought event possibly had on the more vulnerable vegetation. Moreover, it was difficult to remotely define control sites due to the lack of detailed information on vegetation coverage at the affected sites and the control sites defined. This means that how well the control sites represent the affected sites is unknown. Additionally, as earlier mentioned, variability in the control sites could lead to that non-existing impacts are shown, or that the values of the affected sites are less separable from values of the control sites. Consequently, the effects of the fire and frost drought events would not be shown in  $\Delta$ albedo and  $\Delta$ NDVI, for all sites and the different land cover categories.

Moreover, for the burned sites, it could have been relevant to use other remotely sensed estimate of vegetation such as burn severity index to test the portion of vegetation and soil organic matter consumed in the fire and thus the damage ratio of the vegetation at the affected studied, as well as check the damage ratio at the control sites. For the frost drought events, it could be that these events are not limited to a well-defined area, such as fire events, but may occur spatially irregularly. Therefore, even though snow depth data from meteorological stations was used, there is a risk that also the selected control sites have been affected. If field studies would have been conducted information on physical properties, and the extent of the damaged area could have been collected and verified.

The limitation of only using the 5 growing seasons before and after the events was made to avoid that variations that naturally occur in ecosystems over time would obscure the effect of the event or result in a difference that was not caused by the event. Still, this limitation was made with the assumption that no trends that strongly vary from the mean occur during these years. If so, that trend could largely impact the mean that the after-event values are being compared with. Moreover, averaging the 5 growing seasons before and after the events also holds a risk that very short-lasting effect are being obscure and not shown in the results.

The CFMask algorithm did not mask out clouds as expected, which is an important uncertainty of the results. With the aim of compensating for this error, data affected by unfiltered clouds were removed manually. All mentioned uncertainties indicate the complexity of ecosystems and why both field studies and studies using remote sensing are important. Field works allow us to collect observational data, while the great advantage of remote sensing analysis is that it allows us to go back in time and analyse historical data. Furthermore, when conducting multiple tests simultaneously there is an increased risk that obtained results are caused by the chance. Therefore, it would have been relevant to apply a correction method in order to compensate for the large number of statistical tests conducted. Also, the data set used for analysis was relatively small which was problematic when assigning the sites to land coverage categories, and therefore only 3 categories were created.

Finally, this work studied single events, but increased frequency in fire and frost drought events could possibly cause pronounced effects on vegetation (Parmentier et al. 2018; Bokhorst el al. 2012; Jolly et al. 2015), as well as that extreme event can possibly interact, making the vegetation more vulnerable to other type of extremes and causing greater severity (Bjerke et al 2017). Thus, these interactions must be studied for a complete understanding.

#### Conclusion

The results presented in this report show no impact on vegetation greenness after single fire events. A decrease in the  $\Delta$ albedo (mean of mean  $\Delta$ albedo: from -0.002 ± 0.020 to -0.023 ± 0.012) was found only for the land cover category *Mosaic tree*, *shrubs, herbaceous cover* and albedo. Indicating that the land cover class *Mosaic tree*, *shrubs, herbaceous cover* is more vulnerable to the impacts of fire events than the other two studied land cover categories; *Broadleaved forest* and *Needleleaved forest*. For the frost drought events no impacts was found on vegetation greenness when analysing all sites. However, an increase in  $\Delta$ albedo (mean of mean  $\Delta$ albedo: from 0.004 ± 0.042 to 0.009 ± 0.038) was found. Although a weak difference (p=0.049), this is an indication that how frost drought impacts albedo values should be further investigated, and especially since no earlier studies investigating the impact of frost drought on albedo were found.

However, uncertainties and limitations of the 30 m Landsat data and variability in values of control sites, as well as unknown intensity of the events, size of damage areas and damage ratio of vegetation, are important sources of error of this study. Also, this work studied single events, but increased frequency in fire and frost drought events and interactions between different types of extremes could possibly cause pronounced effects on vegetation. Thus, these effects and interactions must be investigated in a more comprehensive study, including field observations, to further determine the effects of fire and frost drought event on vegetation.

#### References

- Aartsma, P., Asplund, J., Odland, A., Reinhardt, S., and Renssen, H. 2020. Surface albedo of alpine lichen heaths and shrub vegetation. *Arctic, Antarctic, and Alpine Research* 52:312–322. https://doi.org/10.1080/15230430.2020.1778890
- Alexander, H. D., Mack, M. C., Goetz, S., Beck, P. S. A., and Belshe, E. F. 2012. Implications of increased deciduous cover on stand structure and aboveground carbon pools of Alaskan boreal forests. *Ecosphere* 3: 45. https://doi.org/10.1890/es11-00364.1
- AMAP. 2017. Snow, Water, Ice and Permafrost in the Arctic (SWIPA) 2017. Arctic Monitoring and Assessment Programme (AMAP), Oslo, Norway.
- Barrett, K., Rocha, A. V., Van De Weg, M. J., Shaver, G., and Janet, M. 2012.
  Remote Sensing Letters Vegetation shifts observed in arctic tundra 17 years after fire Vegetation shifts observed in arctic tundra 17 years after fire. *Remote Sensing Letters* 3: 729–736. https://doi.org/10.1080/2150704X.2012.676741
- Bernhardt, E. L., Hollingsworth, T. N., and Chapin, F. S. 2011. Fire severity mediates climate-driven shifts in understorey community composition of black spruce stands of interior Alaska. *Journal of Vegetation Science* 22: 32–44. https://doi.org/10.1111/j.1654-1103.2010.01231.x
- Bjerke, J. W. 2011. Winter climate change: Ice encapsulation at mild subfreezing temperatures kills freeze-tolerant lichens. *Environmental and Experimental Botany* 72: 404–408. https://doi.org/10.1016/j.envexpbot.2010.05.014
- Bjerke, J. W., Karlsen, S. R., Høgda, K. A., Malnes, E., Jepsen, J. U., Lovibond, S., Vikhamar-Schuler, D., and Tømmervik, H. 2014. Record-low primary productivity and high plant damage in the Nordic Arctic Region in 2012 caused by multiple weather events and pest outbreaks. *Environmental Research Letters* 9: 084006. https://doi.org/10.1088/1748-9326/9/8/084006
- Bjerke, J. W., Treharne, R., Vikhamar-Schuler, D., Karlsen, S. R., Ravolainen, V., Bokhorst, S., Phoenix, G. K., Bochenek, Z., et al. 2017. Understanding the drivers of extensive plant damage in boreal and Arctic ecosystems: Insights from field surveys in the aftermath of damage. *Science of the Total Environment* 599– 600: 1965–1976. https://doi.org/10.1016/j.scitotenv.2017.05.050
- Bokhorst, S., Bjerke, J. W., Bowles, F. W., Melillo, J., Callaghan, T. V., and Phoenix, G. K. 2008. Impacts of extreme winter warming in the sub-Arctic: Growing season responses of dwarf shrub heathland. *Global Change Biology*, 14: 2603– 2612. https://doi.org/10.1111/j.1365-2486.2008.01689.x

- Bokhorst, S. F., Bjerke, J. W., Tømmervik, H., Callaghan, T. V., and Phoenix, G. K. 2009. Winter warming events damage sub-Arctic vegetation: Consistent evidence from an experimental manipulation and a natural event. *Journal of Ecology* 97: 1408–1415. https://doi.org/10.1111/j.1365-2745.2009.01554.x
- Bokhorst, S., Tømmervik, H., Callaghan, T.V., Phoenixa, G.K., and Bjerke, J. W. 2012. Vegetation recovery following extreme winter warming events in the sub-Arctic estimated using NDVI from remote sensing and handheld passive proximal sensors. *Environmental and Experimental Botany* 81: 18–25. https://doi.org/10.1016/j.envexpbot.2012.02.011
- Box, J. E., Colgan, W. T., Christensen, T. R., Schmidt, N. M., Lund, M., Parmentier, F. J. W., Brown, R., Bhatt, U. S., et al. 2019. Key indicators of Arctic climate change: 1971-2017. *Environmental Research Letters* 14: 045010. https://doi.org/10.1088/1748-9326/aafc1b
- Brakstad, Tommy. 2014, January 28. Brannen i Flatanger beveger seg fortsatt, men brannvesenet er nå mer optimistiske enn tidligere. *Nettavisen*. Retrieved March, 2021, from https://www.nettavisen.no/nyheter/ti-bygninger-i-fare/s/12-95-3748362
- Bret-Harte, M. S., Mack, M. C., Shaver, G. R., Huebner, D.C., Johnston, M., Mojica, C. A., Pizano, C., and J. A. Reiskind. 2013. The response of Arctic vegetation and soils following an unusually severe tundra fire. *Philosophical Transactions* of the Royal Society 368: 20120490. https://doi.org/10.1098/rstb.2012.0490
- Bright, B. C., Hicke, J. A., and Meddens, A. J. H. 2013. Effects of bark beetle-caused tree mortality on biogeochemical and biogeophysical MODIS products. *Journal* of Geophysical Research: Biogeosciences 118: 974–982. https://doi.org/10.1002/jgrg.20078
- Calabria, L. M., Petersen, K., Hamman, S. T., and Smith, R. J. 2016. Prescribed Fire Decreases Lichen and Bryophyte Biomass and Alters Functional Group Composition in Pacific Northwest Prairies. *Northwest Science* 90: 470–483. https://doi.org/10.3955/046.090.0407
- Chambers, S. D., Beringer, J., Randerson, J. T., and Chapin, I. S. 2005. Fire effects on net radiation and energy partitioning: Contrasting responses of tundra and boreal forest ecosystems. *Journal of Geophysical Research D: Atmospheres* 110: 1–9. https://doi.org/10.1029/2004JD005299
- Chuevieco, E. 2016. *Fundamentals of Satellite Remote Sensing*. Boca Raton: CRC Press.

- Cox, P. M., Betts, R. A., Bunton, C. B., Ezzery, R. L. H., Rowntree, P. R., Smith. J., 1999. The impact of new land surface physics on the GCM simulation of climate and climate sensitivity. *Climate Dynamics* 15: 183-203. https://doi.org/10.1007/s003820050276
- Díaz-Delgado, R., Lloret, F., and Pons, X. 2003. Influence of fire severity on plant regeneration by means of remote sensing imagery. *International Journal of Remote Sensing* 24: 1751–1763. https://doi.org/10.1080/01431160210144732
- DSB. 2014. Brannene i Lærdal, Flatanger og på Frøya vinteren 2014, 978-82-7768-342-3.
- Elmendorf, S. C., Henry, G. H. R., Hollister, R. D., Björk, R. G., Bjorkman, A. D., Callaghan, T. V., Collier, L. S., Cooper, E. J., et al. 2012. Global assessment of experimental climate warming on tundra vegetation: Heterogeneity over space and time. *Ecology Letters* 15: 164–175. https://doi.org/10.1111/j.1461-0248.2011.01716.x
- EOS. 2019. NDVI FAQ: All You Need To Know About Index. Retrieved 28 April, 2021, from https://eos.com/b log/ndvi-faq-all-you-need-to-know-about ndvi/#:~:text=In%20most %20cases% 2C%20NDVI%20values, possible%20density%20of%20green%20leaves.
- European Environment Agency (Copernicus programme). 2016. European Digital Elevation Model version 1.1. Retrieved 25 March, 2021, from https://land.copernicus.eu/imagery-in-situ/eu-dem/eu-dem-v1.1?tab=mapview
- Euskirchen, E. S., Mcguire, A. D., Rupp, T. S. Chapin III., and Walsh, J. E. 2009. Projected changes in atmospheric heating due to changes in fire disturbance and the snow season in the western Arctic, 2003-2100. *Journal of Geophysical Research* 114:1–15. https://doi.org/10.1029/2009JG001095
- ESA. 2015. Land Cover Maps v2.0.7. Retrieved 25 March, 2021, from http://maps.elie.ucl.ac.be/CCI/viewer/download.php
- Foga, S., Scaramuzza, P. L., Gou, S., Zhu, Z., Dilley, R. D., Beckmann, T., Schmidt, G. L., Dwyer. J. L., et al. 2017. Cloud detection algorithm comparison and validation for operational Landsat data products. *Remote Sensing of Environment* 194: 379-390. http://doi.org/10.1016/j.rse.2017.03.026
- French, N. H. F., Whitley, M. A., and Jenkins, L. K. 2016. Fire disturbance effects on land surface albedo in Alaskan tundra. *Journal of Geophysical Research: Biogeosciences*, 121: 841–854. https://doi.org/10.1002/2015JG003177

- Google Earth Engine. n.d. Platform. Retrieved 18 May, 2021, form https://earthengine.google.com/platform/
- Government of Canada. n.d. Download data. Retrieved 7 February, 2021, from https://cwfis.cfs.nrcan.gc.ca/datamart/download/nfdbpoly?token=caf55e87a64b8 9fec6e24c65c7c43929
- Hamarøy kommune (2020). Oversikt over branner i naturmark (skog/lyng) for Hamarøy kommune fra 1.1.2020. Svein Morten Sandnes.
- Heim, R. J., Bucharova, A., Brodt, L., Kamp, J., Rieker, D., Soromotin, A. V., Yurtaev, A., and Hölzel, N. 2020. Post-fire vegetation succession in the Siberian subarctic tundra over 45 years. *Science of the Total Environment* 760: 143425. https://doi.org/10.1016/j.scitotenv.2020.143425
- Huang, C. Y., and Anderegg, W. R. L. 2014. Vegetation, land surface brightness, and temperature dynamics after aspen forest die-off. *Journal of Geophysical Research G: Biogeosciences* 119: 1297–1308. https://doi.org/10.1002/2013JG002489
- Hollinger, D. Y., Ollingerw, S. V., Richardsonw, A. D., Meyersz, T. P., Dail, D. B., Martinw, M. E., Scott, N. A., Arkebauerk, S. B., et al. 2010. Albedo estimates for land surface models and support for a new paradigm based on foliage nitrogen concentration. *Global Change Biology* 16: 696–710. https://doi.org/10.1111/j.1365-2486.2009.02028.x
- IPCC. 2012. Managing the Risks of Extreme Events and Disasters to Advance Climate Change Adaptation [Field, C.B., Barros, V., Stocker, T.F., Qin, D., Dokken, D.J., Ebi, K.L., Mastrandrea, M.D., Mach, K.J., Plattner, G.-K., Allen, S.K., Tignor, M., and Midgley, P.M. (eds.)] Cambridge University Press, Cambridge, UK
- Jandt, R., Joly, K., Meyers, C. R., & Racine, C. (2008). Slow recovery of lichen on burned caribou winter range in Alaska tundra: Potential influences of climate warming and other disturbance factors. *Arctic, Antarctic, and Alpine Research*, 40(1), 89–95. https://doi.org/10.1657/1523-0430(06-122)[JANDT]2.0.CO;2
- Jenkins, L. K., Bourgeau-Chavez, L. L., French, N. H. F., Loboda, T. V., and Thelen, B. J. 1992. remote sensing Development of Methods for Detection and Monitoring of Fire Disturbance in the Alaskan Tundra Using a Two-Decade Long Record of Synthetic Aperture Radar Satellite Images. *Remote Sens* 6: 6347–6364. https://doi.org/10.3390/rs6076347
- Jones, B. M., Kolden, C. A., Jandt, R., Abatzoglou, J. T., Urban, F., and Arp, C. D. 2009. Fire Behavior, Weather, and Burn Severity of the 2007 Anaktuvuk River

Tundra Fire, North Slope, Alaska. *Arctic, Antarctic, and Alpine Research* 41: 309–316. https://doi.org/10.1657/1938-4246-41.3.309

- Jolly, W. M., Cochrane, M. A., Freeborn, P. H., Holden, Z. A., Brown, T. J., Williamson, G. J., and Bowman, D. M. J. S. 2015. Climate-induced variations in global wildfire danger from 1979 to 2013. *Nature Communications* 6:7537. https://doi.org/10.1038/ncomms8537
- Key, C. H., and Benson, N. 2006. Landscape Assessment: Ground measure of severity, the Composite Burn Index; and Remote sensing of severity, the Normalized Burn Ratio. In *Fire Effects Monitoring and Inventory System*, ed. Lutes, D. C., Keane, R. E., Caratti, J. F., Key, C. H., Benson, N. C., Sutherland, S., and L. J. Gangi, LA 1-51. Ogden: USDA Forest Service, Rocky Mountain Research Station.
- Kottek, M., Grieser, J., Beck, C., Rudolf, B., and Rubel, F. 2006. World map of the Köppen-Geiger climate classification updated. *Meteorologische Zeitschrift* 15: 259–263. https://doi.org/10.1127/0941-2948/2006/0130
- Liang, S. 2001. Narrowband to broadband conversions of land surface albedo I Algorithms. *Remote Sensing of Environment* 76: 213-238. https://doi.org/10.1016/S0034-4257(00)00205-4
- Mack, M. C., Bret-Harte, M. S., Hollingsworth, T. N., Jandt, R. R., Schuur, E. A. G., Shaver, G. R., and Verbyla, D. L. 2011. Carbon loss from an unprecedented Arctic tundra wildfire. *Nature* 475: 489- 492. https://doi.org/10.1038/nature10283
- MSB (2021). Utdrag ur Räddningstjänstens insatser gällande bränder i skog och mark 2000-2019. Joakim Ekberg. Version (MSB 2021-02203-1).
- Norsk klimaservicesenter. n.d. Retrieved 10 March, 2021, from https://seklima.met.no/observations/
- Oke, T.R., 1987. Boundary layer climates. The Journal of Applied Ecology. 17 (2): 517. https://doi.org/10.2307/2402350
- Parmentier, F. W., Rasse, D. P., Lund, M., Bjerke, J. W., Drake, B.G., Weldon, S., Tømmervik, H., and Hansen, G. H. 2018. Vulnerability and resilience of the carbon exchange of a subarctic peatland to an extreme winter event. *Environmental Research Letters* 13: 065009. https://doi.org/10.1088/1748-9326/aabff3

- Payette, S., and Delwaide, A. 2018. Tamm review: The North-American lichen woodland. *Forest Ecology and Management* 417: 167–183. https://doi.org/10.1016/j.foreco.2018.02.043
- Phoenix, G. K., and Bjerke, J. W. 2016. Arctic browning: extreme events and trends reversing arctic greening. *Global Change Biology* 22: 2960–2962. https://doi.org/10.1111/gcb.13261
- Racine, C., Jandt, R., Meyers, C., and J. D. (2004). Tundra Fire and Vegetation Change along a Hillslope on the Seward Peninsula , Alaska, U.S.A. Arctic, Antarctic, and Alpine Researc, 36(1), 1–10.
- Ranson, K. J., Sun, G., Kovacs, K., and Kharuk, V. I. 2003. MODIS NDVI Response Following Fires in Siberia. *IEEE* 5: 3290–3292. https://doi.org/10.1109/igarss.2003.1294759
- Riseth, J. Å., Tømmervik, H., Helander-Renvall, E., Labba, N., Johansson, C., Malnes, E., Bjerke, J., Jonsson, W. C., et al. 2011. Sámi traditional ecological knowledge as a guide to science: Snow, ice and reindeer pasture facing climate change. *Polar Record* 47: 202–217. https://doi.org/10.1017/S0032247410000434
- Ritz, E., Bjerke, J. W., and Tømmervik, H. 2020. Monitoring Winter Stress Vulnerability of High-Latitude Understory Vegetation Using Intraspecific Trait Variability and Remote Sensing Approaches. *Sensors* 20: 1-16. https://doi.org/10.3390/s20072102
- Rocha, A. V., and Shaver, G. R. 2011. Postfire energy exchange in arctic tundra: The importance and climatic implications of burn severity. *Global Change Biology* 17:2831–2841. https://doi.org/10.1111/j.1365-2486.2011.02441.x
- Rouse, J.W., Haas, R. H., Schell, J. A., Deering, D.W., and Harlan, J. C. 1974.
  Monitoring the Vernal Advancement of Retrogradation of Natural Vegetation.
  Texas A&M University Remote Sensing Center. Report RSC 1978-4, Greenbelt, USA.
- Saxena, S., Rabha, A., Tahlani, P, and Ray, S.S. 2020. Crop Situation in India, Before, During and After COVID-19 Lockdown, as Seen from the Satellite Data of Resourcesat-2 AWiFS. *Journal of the Indian Society of Remote Sensing* 49: 356-376. https://doi.org/10.1007/s12524-020-01213-5
- Smith, R. B. 2010. The heat budget of the earth's surface deduced from space. Retrieved 6 February, 2021, from https://yceo.yale.edu/sites/default/files/files/Surface\_Heat\_Budget\_From\_Space. pdf

Sizov, O., Ezhova, E., Tsymbarovich, P., Soromotin, A., Prihod'Ko, N., Petäjä, T., Zilitinkevich, S., Kulmala, M., Bäck, J., & Köster, K. (2021). Fire and vegetation dynamics in northwest Siberia during the last 60 years based on high-resolution remote sensing. *Biogeosciences* 18: 207–228. https://doi.org/10.5194/bg-18-207-2021

Stavanger Aftenblad. 29 May, 2013. Frostskader gir dårlig bærår.

- Storey, J. C., and Choate, M. 2000. Landsat 7 on-orbit geometric calibration and performance. Algorithms for Multispectral, Hyperspectral, and Ultraspectral Imagery VI 4049: 143-154. https://doi.org/10.1117/12.410335
- Svarstad, J. 2014, January 29. Har kontroll med brannen på Frøya. *Aftenposten*. Retrieved March, 2021, from https://www.aftenposten.no/norge/i/rLql3/harkontroll-med-brannen-paa-froeya
- Treharne, R., Bjerke, J. W., Tømmervik, H., Stendardi, L., and Phoenix, G. K. 2019. Arctic browning: Impacts of extreme climatic events on heathland ecosystem CO 2 fluxes. *Global Change Biology* 25: 489–503. https://doi.org/10.1111/gcb.14500
- Tucker, C. J. 1979. Red and photographic infrared linear combinations for monitoring vegetation. *Remote sensing of Environment* 8: 127–150. https://doi.org/10.1016/0034-4257(79)90013-0
- Tsuyuzaki, S., Kushida, K., and Kodama, Y. 2009. Recovery of surface albedo and plant cover after wildfire in a Picea mariana forest in interior Alaska. *Climatic Change* 93: 517–525. https://doi.org/10.1007/s10584-008-9505-y
- USGS. n.d.a. Landsat 4. Retrieved 6 February, 2021, from https://www.usgs.gov/core-science-systems/nli/landsat/landsat-4?qtscience\_support\_page\_related\_con=0#qt-science\_support\_page\_related\_con
- USGS. n.d.b. NDVI, the Foundation for Remote Sensing Phenology Retrieved 28 April, 2021, from https://www.usgs.gov/core-sciencesystems/eros/phenology/science/ndvi-foundation-remote-sensing-phenology?qtscience\_center\_objects=0#qt-science\_center\_objects
- Xiong, X., Chiang, K., Sun, J., Barnes, W. L., Guenther, B., and Salomonson, V. V. 2008. NASA EOS Terra and Aqua MODIS on-orbit performance. *Space Research* 43: 413–422. https://doi.org/10.1016/j.asr.2008.04.008

### Appendix I

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T1: Details on the study sites for the frost drought events. The sites are defined by Bjerke et al. (2017), Parmentier et al. (2018), and Stavanger Aftenblad (2013).

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T2: Sites studied for the fire events. The sites are based on data from MSB (2021), Hamarøy municipality (2020), and DSB (2014).

### Appendix II

### Appendix III



F1. Example of cloud pixels (white pixels) that have not been detected by the CFMask algorithm (Foga et al 2017). The image is a mosaic of 10 Landsat 7 images from 1st of July to 31 of August 2001.