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# Sediment Transport in the Piura River. Piura, Peru. 

An application with the HEC-RAS model

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Division of Water Resources Engineering
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Research for this thesis involved being on- site in Piura. Funding for the trip was provided by The Crafoord Foundation. My time in Piura allowed me to gain insight on the topic, which would hardly have been possible from a distance.


#### Abstract

This thesis is an investigation of sediment transport in the Piura River in Peru. Years of anthropogenic changes have redirected the river towards the south and have caused changes in the flow regimen (Alvarado Ancieta \& Ettmer, 2007). It is suspected that the redirection towards the south has resulted in deposition of sediments which has reduced the capacity of the river in the lower reaches (Reyes Salazar, 2024b). As such, the aim of this thesis is to compare hydraulic conditions of the current route through Virrila in the south and the original course towards Sechura to assess sediment transport capacity.

HEC-RAS version 6.3.1 is used for hydraulic modelling. The model is calibrated using flow data from the 2017 flood event. Flows of $1,000 \mathrm{~m}^{3} / \mathrm{s}$ and $3,200 \mathrm{~m}^{3} / \mathrm{s}$ are computed in steady state for velocity, depth, and shear stress profiles as output. Excess shear is used to assess sediment transport capacity. It was found that the general trend for Virrila is deposition of sediment in the lower reach, while the course towards Sechura produces considerably larger excess shear values, thereby indicating increased capacity for sediment transport.


## Glossary

ANA: Autoridad Nacional del Agua- National Authority for Water
ARCC: Autoridad para la Reconstrucción con Cambios- Authority for Reconstruction with Change

ENFEN: Estudio nacional del Fenómeno El Niño- National Study of El Niño Phenomenon

ENSO: El Niño southern oscillation
ICEN: Índice Costero el Niño- El Niño coastal index
ONI: Oceanic Niño index
PNUD study: study by the National University of Piura and University of Piura (2001)

SENAMHI: Servicio Nacional de Meteorología e Hidrología- National Meteorology and Hydrology Service

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## 1. Introduction

### 1.1. General Background

The northern coast of Peru is regularly subject to flooding. The Piura region is consistently among the regions most affected by the floods on a national level together with Tumbes, Lambayeque, and La Libertad (Organización Panamericana de la Salud, 2018) shown in Figure 1. The region covers an area of around $40,580 \mathrm{~km}^{2}$ (UDEP \& UNP, 2001e) and has a population of approximately 1.8 million (ARCC, 2022). The consequences of flooding in the region include loss of agricultural land, destruction of infrastructure, sanitation line failure, and an increase of diseases (SENAMHI, 2014).


Figure 1a- Overview map of Peru (The Only Guide Peru, n.d.). Ib- Map of Piura Region (Google Maps, 2023)

The capital city of the region, also called Piura, has a population of approximately 517,300 (Atarama \& Rashid, 2019) and was founded in 1532 by the Spanish settlers (MGI, n.d.). Piura is partially located in the alluvial fan of the Piura River (Fernández et al., 2021). The river is non-perennial and originates in the highlands to the east at an altitude of $3,370 \mathrm{~m}$ above sea level (Alvarado Ancieta, 2023) where it then flows west towards the Pacific Ocean and passes through a series of lagoons in the Sechura Desert before exiting into the Virrilá estuary (ARCC, 2022). An overview map of the lower basin is shown in Figure 2.


Figure 2- Overview map of lower basin (Google Earth, 2024)

The semi-arid climate in the lower basin changes during El Niño episodes, during which there is frequent and heavy rainfall for a period of 4-5 months (UDEP \& UNP, 2001e). Flood events in Piura are generally linked to the increased precipitation experienced during El Niño. The current river system is incapable of controlling the large quantities of water that develop during this event (UDEP \& UNP, 2001e). In the past, flooding occurred due to a lack of urban drainage together with erosion of the riverbed which resulted in the collapse of river infrastructure (Reyes Salazar, 2024b). However, the most recent flood event of 2016-2017 revealed an ongoing process of aggradation which is reducing the capacity of the river (Reyes Salazar, 2024a). Figure 3 shows data collected by the University of Piura for the bottom level and water level for three different years at three different flows. The bottom level in the urban reach has increased 6 m since 1998. Consequently, the capacity of the river has decreased.


Figure 3- Section Cáceres Bridge showing bottom level development. (adapted from Reyes Salazar, 2024b)

The lower basin has been flooded 6 times in the last 45 years (Castillo Zavaleta, et al., 2019). This has encouraged many investigations and studies concerning vulnerability, urban drainage, El Niño, alleviation plans, and sediment transport. As mentioned, the river flows through a series of lagoons
before exiting into the ocean. For that reason, it should be noted that the Ramón Lagoon has been steadily filling up with sediments. An estimated 30\% of the original capacity remains (Reyes Salazar, 2024b), the sedimentation can be observed as shown in Figure 4.


Figure 4- Sedimentation of the Ramón Lagoon (Google Earth, 2024)
Historically, the Piura River flowed through Sechura in the west (Zegarra Dávila, 2024). However, the growing importance of agriculture in the region led to the construction of dikes which resulted in the redirection towards the south (Alvarado Ancieta, 2023). The current route through Virrilá has a smaller slope and is longer than through Sechura (Reyes Salazar, 2024b). As a result, there is less energy in the river and sedimentation occurs (Reyes Salazar, 2024b). Experts suspect that deposition of sediment along some stretches of the river will continue to reduce the capacity and will render the lower basin vulnerable to flooding even during small flows (Reyes Salazar, 2024b).

### 1.2. Aim

The aim of this thesis is to compare sediment transport in the Piura River for two different river courses. The routes are shown in Figure 5, these are the existing course through the Virrilá Estuary, and a proposed course through Sechura. The software HEC-RAS is used to create a river model for the two routes. Velocity and depth output is used to compare transport capacity.


Figure 5- Piura River routes through Virrilá and Sechura (Google Earth, 2024).

The proposed route towards Sechura follows parts of the ancient river course, and is expected to have a larger transport capacity, thereby reducing sedimentation (Zegarra Dávila, 2024). Various proposals are being developed to decrease the vulnerability in the lower basin, therefore describing sediment transport through river modelling can be a useful tool to reinforce suggestions and for planning of future interventions.

### 1.3. Methodology

This thesis is a collaboration with the University of Piura. Research involved being on site in Piura for improved data acquisition. Information for this thesis was sourced using Google Scholar, company sites, personal communication with experts in the Piura region, field visits, and was provided by the University of Piura. Relevant material included articles, project documents, reports both published and unpublished, and topographic data.

### 1.4. Thesis Outline

This report consists of 10 chapters. Chapter 1, "Introduction" provides background information, the scope of the study, and objectives. The following chapter "Piura River Catchment" provides information on the study area in terms of flood events, hydrological conditions, and morphology of the river. Chapter 3 "Previous Studies" describes studies and projects related to this research. Chapter 4 "Field Visit" describes the excursion to the Piura River and major observations. Next, Chapter 5 "Theoretical Background" describes theory and the use of HEC-RAS. Continuing in Chapter 6 "Data Employed", data acquired for sediment transport capacity analysis and HEC-RAS modelling is described. Chapter 7 "HEC-RAS Model Set Up and Output Processing" describes input, assumptions, and parameters used in HEC-RAS. The following chapter, "Results" describes calibration, HEC-RAS, and sediment transport capacity results. Next, in Chapter 9 "Discussion and Recommendations" provides insight on the implication of the results, gives suggestions for additional sediment investigation and recommendations for further studies. Finally, Chapter 10 "Conclusion" re-visits the aim of the study in relation to the results. A bibliography and appendices are found at the end of the report.

## 2. Piura River Catchment

This chapter provides information on the Piura River basin in terms of land use, historical flood events, hydrology, and river morphology.

### 2.1. Introduction to the Catchment

The Piura River basin has an area of $12,216 \mathrm{~km}^{2}$ (Moreano Segovia \& Yauri Quispe, 2008.) which accounts for $31 \%$ of the Piura region (CIIFEN, 2020). The river originates in the eastern highlands and runs west where it enters a series of lagoons before indirectly exiting into the Pacific Ocean through the Virrilá estuary, it has a length of 326km (Moreano Segovia \& Yauri Quispe, 2008). The catchment is defined by three zones, namely the upper basin, midbasin, and lower basin as shown in Figure 6.


Figure 6- Piura River catchment zones (adapted from Moreano Segovia \& Yauri Quispe, 2008)

These are characterized as follows:

- Upper basin- The section of the basin at an altitude between 350$3,650 \mathrm{~m}$ asl with defined valleys and steep slopes (Moreano Segovia \& Yauri Quispe, 2008). The climate is subhumid, and precipitation ranges from 500-1,000mm yearly (Moreano Segovia \& Yauri Quispe, 2008). Shown in purple in Figure 6.
- Mid basin- This section ranges from $50-350 \mathrm{~m}$ asl, it is a hilly landscape with yearly precipitation ranges from $100-500 \mathrm{~mm}$ (Moreano Segovia \& Yauri Quispe, 2008). Shown in orange in Figure 6.
- Lower basin- The lower basin is very flat land with altitudes ranging from 0-50m asl (Moreano Segovia \& Yauri Quispe, 2008). It is an arid climate with yearly precipitation at less than 100mm (Moreano Segovia \& Yauri Quispe, 2008). Shown in green in Figure 6.

The Piura River is non-perennial. In the highlands the bedrock is closer to the ground and there is more precipitation, as such the river is permanent in the upper basin (Atarama \& Rashid, 2019). In the mid basin, the river feeds irrigation canals and recharges an aquifer since the bedrock is deeper (Atarama \& Rashid, 2019). Thus, the river regularly runs dry from August to December, especially in years with less rainfall (Atarama \& Rashid, 2019). In the lower basin the Piura River is fed by the Chira River to the north at a point 30 km upstream of the city of Piura, it provides the lower basin with $60 \mathrm{~m}^{3} / \mathrm{s}$ of water (PECHP, 2017).

Historical records depict the exit of the Piura River at Sechura, where it discharged into the Pacific Ocean through 2-3 branches (Guererro et al., 2015) as shown by Figure 7 which is a map from 1779 (Zegarra Dávila, 2023).


Figure 7- Map showing the Piura River from 1779 (adapted from Zegarra Dávila, 2023)
During large rain events, a series of lagoons formed further south, namely the Napique/Ramón Lagoon and La Niña Lagoon. When exceeded the capacity through Sechura, the river would first overflow into the Napique/Ramón Lagoons, then form La Niña Lagoon, and discharge into Virrilá (Zegarra Dávila, 2023). The last overflow route was further south through Reventazón as shown in Figure 8 (Zegarra Dávila, 2023).


Figure 8- Overflow routes of the Piura River (Google Earth, 2024)

Agriculture plays an important role in the Piura region; it is considered the main reason for regional employment (MGI, n.d.) and makes up 5.7\% of the national GVA (CIIFEN, 2020). Consequently, the river has been subject to land use changes due to agricultural activity and human settlements. (Alvarado Ancieta, 2023). At the beginning of the $20^{\text {th }}$ century, the rapid growth of the cotton industry led to the construction of dikes to protect crops in the floodplain during increased flow (Zegarra Dávila, 2023). This was the beginning of a gradual redirection of the river towards the lagoon system in the south through the construction of dikes. Large flows would regularly breach the dikes, consequently these were reinforced in the 1970's and became permanent structures (Azurin Gonzáles, 2010). The location of the dikes is shown in Figure 9.


Figure 9- Dikes in lower basin (Google Earth, 2024)

### 2.2. Recent Flood Events

As mentioned, flooding in the Piura region is a recurring phenomenon. Flooding in the years 1983, 1998 and 2017 are regularly referred to as the most disastrous flood events the region has experienced. The most significant damages are identified as the destruction of roads, sewer systems, as well as homes and establishments, which caused a disruption in the provision of potable water and other vital goods and services in the region (UDEP \& UNP, 2001a).

The flow in 1983 peaked at $3,200 \mathrm{~m}^{3} / \mathrm{s}$ (Farías Zegada, 2008). This caused overflow in some spans as well as significant erosion along the course of the river (Azurin Gonzáles, 2010). Erosion was notable at various riverine structures and along the banks in the urban zone which were not concreted at the time (Alvarado Ancieta \& Ettmer, 2007). The flood resulted in the collapse of approximately 15 km of the protective dikes on the west side of the river as well as the collapse of the spillway at the Los Ejidos Dam due to local erosion at the base of both structures (Alvarado Ancieta \& Ettmer, 2007) (Reyes Salazar, 2024b). The dike breach on the west side of the river resulted in the river retaking the old course towards Sechura, where it destroyed fields and settlements along the way (Alvarado Ancieta \& Ettmer, 2007). The flooding due to the breaches as well as inadequate drainage in the urban areas resulted in losses of around $\mathrm{S} / 12,300$ million, which accounted for $11.6 \%$ of the country`s GDP (SENAMHI, 2014).

In 1986, following the flood event, dikes were reinforced with concrete along the urban area and with stones and breakwaters along the base (Alvarado Ancieta \& Ettmer, 2007). However, no works were conducted to protect against general erosion of the riverbed (Azurin Gonzáles, 2010).

In 1998 the flow in the river reached a peak at $3,900 \mathrm{~m}^{3} / \mathrm{s}$ (Farías Zegada, 2008). The event had similar impacts as that of 1983 with losses of approximately $\mathrm{S} / 13$, 100 million, or $6.2 \%$ of the GDP (SENAMHI, 2014). During this event, the increased transport of sediments resulted in erosion at the foot of the dikes and breached the left dike, further the bridges San Miguel, Bolognesi, and Independencia collapsed as shown in Figure 10 (Alvarado Ancieta \& Ettmer, 2007). The San Miguel and Bolognesi are a part of the urban reach while the Independencia bridge is further downstream of the city Piura as shown in Figure 11.


Figure 10a-collapsed San Miguel and Bolognesi Bridges 1998. 10b-collapsed Independencia bridge 1998. (Alvarado Ancieta \& Ettmer, 2007)


Figure 11- Location of 1998 collapsed bridges (Google Earth, 2024)

The 2017 flood peaked at $3,468 \mathrm{~m}^{3} / \mathrm{s}$ (PECHP, 2017). Although this flow is similar to that of the previous events, the river overflowed in unexpected spans in the urban area (Reyes Salazar, 2024b). In previous years the large flow caused erosion which increased the capacity of the river (Azurin Gonzáles, 2010). However, in 2017 the erosion did not occur thereby causing overflow (Reyes Salazar, 2024b). According to the regional director of health in Piura, Dr Hernán, the region had expected strong rains and low impact floods but did not expect the river to overflow (Organización Panamericana de la Salud, 2018). The flood damaged infrastructure such as health and education establishments, roads, and sewers (Organización Panamericana de la Salud, 2018). It also destroyed thousands of homes, as well as agricultural lands, leaving the Piura region in a state of emergency (Organización Panamericana de la Salud, 2018). In addition to the damage to infrastructure, there was a sudden spread of diseases due to inadequate sanitation as well as a dramatic increase in Dengue, which became a further obstacle for the region (Organización Panamericana de la Salud, 2018).

### 2.3. Hydrological Conditions and the Role of EI Niño

Northwestern Peru is a coastal desert with a dry and tropical climate featuring rainy seasons from December to March (Di Liberto, 2017). The average temperature in the day is $33-35^{\circ} \mathrm{C}$ and $18-23^{\circ} \mathrm{C}$ during nighttime (Moreano Segovia \& Yauri Quispe, 2008.). The climate along the Peruvian coast is strongly linked to oceanic conditions in the Pacific Ocean (SENAMHI, 2014). Usually, cold water from the Antarctic region is brought up to the South American coast by the Humboldlt Current while trade winds blowing towards the west push warm water away from the coast and towards Asia (Welt Hunger Hilfe, n.d.). Some years trade winds weaken, and this allows kelvin waves carrying warm water to move back into the eastern pacific (NOAA, 2023c)
(SENAMHI, 2014). This increase in temperature in the eastern pacific is known as El Niño and is the warm phase of the El Niño Southern Oscillation (ENSO) pattern which recurs every two to seven years (NOAA, 2023c). ENSO is monitored in the pacific in two zones as shown in Figure 12.


Figure 12-ENSO monitoring zones (NOAA, 2023b)

Zone Niño 3.4 is used to monitor the Oceanic Niño Index (ONI). This index is measured as an average of temperature anomalies during three consecutive months in the region Niño 3.4. Cold and warm episodes are indicated when a threshold of $\pm 0.5^{\circ} \mathrm{C}$ is exceeded (NOAA, 2023b). When the threshold is exceeded for a minimum of five consecutive months, an El Niño (warm) or La Niña (cold) episode is defined (NOAA, 2023b). Further, the National Study of El Niño (ENFEN), a multisectoral committee in charge of studying El Niño in Peru, has developed an index to describe the variation in temperature in Niño region $1+2$ called ICEN (SENAMHI, 2014). The index is calculated in the same way as the ONI and indicates the occurrence of an El Niño event on the coast of Peru, known as El Niño Costero. The ONI index is shown for 19502023 in Figure 13 together with threshold values for El Niño classification. Figure 14 shows the El Niño Costero (ICEN) occurrence since 1950.


Figure 13- ONI values for Niño Zone 3.4 for 1950-2023. Temperature deviation for averages of 3 consecutive months. Thresholds at neutral -0.5 to +0.5 , light +0.5 to +1.0 , moderate +0.9 to +1.4 , and strong $>+1.4$ (SENAMHI, 2014). (NOAA, 2023a)

—ICEN
Figure 14- ICEN values for Niño Zone 1+2 for 1950-2023. Temperature deviation for averages of 3 consecutive months. Thresholds at neutral -1.0 to +0.4 , light +0.4 to +1.0 , moderate +1.0 to 1.7, strong +1.7 to +3.0 , and very strong $>+3.0$. (Instituto Geofisico del Perú, 2023)

Figure 15 shows that the events of 1983 and 1998 were characterized as 'strong' El Niño events on the ONI scale, the 2017 event was below the El Niño thresholds. However, the ICEN value indicating temperature anomalies in El Niño zone $1+2$ coast show that there was a temperature increase locally along the coast, which created the conditions for the increased precipitation and flow experienced in 2017 although Zone 3.4 was in La Niña (cold) conditions.


Figure 15a- ONI values for El Niño Events 1982/83, 1997/98, and 2016/17. Temperature deviation for averages of 3 consecutive months (NOAA, 2023a). 15b-ICEN values for El
Niño Events 1982/83, 1997/98, and 2016/17. Temperature deviation for averages of 3 consecutive months (Instituto Geofisico del Perú, 2023).

In the Piura region El Niño affects the pluviometry and hydrological regimen of the river basin (Moreano Segovia \& Yauri Quispe, 2008.). The arid climate is due to the cool coastal air above the pacific which provides a protective barrier against offshore storms (Di Liberto, 2017). However, during El Niño episodes, the coastal air heats up and allows for an increase in precipitation along the coastal region (Di Liberto, 2017). In general, the average yearly precipitation is around $50-100 \mathrm{~mm}$ (Azurin Gonzáles, 2010) (Moreano Segovia \& Yauri Quispe, 2008.). In years with a 'light' El Niño episode, average yearly precipitation is around 135 mm while years with a 'strong' event feature yearly
precipitation of around $3,000 \mathrm{~mm}$ (Azurin Gonzáles, 2010), this is an increase in yearly precipitation by a factor of $30-60$. The relationship between precipitation and ONI/ICEN is shown in Figure 16 for stations in the upper and lower Piura River basin for the years 1972-2013. The data ranges to 2013 as this is what was available on the website of the National Meteorological and Hydrological Service (SENAMHI) for chosen stations.


Figure 16- ICEN and ONI values with daily accumulated precipitation values for the upper and lower basin from 1972-2013

### 2.4. Piura River Morphology

The morphology of a river is governed by shape and hydrology (Guererro et al., 2015). In the case of the Piura River, the shape has been influenced by anthropogenic interventions, and hydrology is largely influenced by ENSO (Guererro et al., 2015).

In the highlands, the river has a large slope of around $15 \%$, after the mountains the slope lowers to around $0.35 \%$ and further decreases to $0.08-0.13 \%$ in the mid basin (Guererro et al., 2015). In the lower basin, the flat terrain has a slope of around $0.03 \%$ (Guererro et al., 2015), which creates conditions for the
formation of an alluvial fan. Research by the University of Piura indicates that flooding of the river used to bring nutrient rich sediment to the floodplains which allowed for vegetative growth (Guererro et al., 2015). However, now the floodplains are restricted to the area within the dikes, this has been made evident by the difference in vegetative coverage inside and outside the dike area as shown in Figure 17.


Figure 17- Vegetation difference inside and outside the dikes (Guererro et al., 2015)
As concluded by the Class-Salzgitter study in 2001, this vegetative growth creates resistance to flow, thereby causing an increase in the level within the enclosed zone (Class-Salzgitter, 2001 cited in Córdova Elera, 2020). In combination with the restriction of the flood plain created by the protective dikes, there is a more concentrated accumulation of sediments which creates overflow as shown in Figure 18 (Castillo Zavaleta, et al., 2019).


Figure 18- Accumulation of sediments inside the floodplain restricted by dikes (adapted from Castillo Zavaleta, et al., 2019)

The cross section of the Piura River varies throughout the lower basin. In the urban reach from Los Ejidos to the Bolognesi bridge river is channeled and has a width of approximately $130-300 \mathrm{~m}$. From the end of the urban area to downstream of Drain 13.08, the river is restricted by protective dikes with a floodplain of a width of $500-800 \mathrm{~m}$ and a main channel width of $100-200 \mathrm{~m}$ (Alvarado Ancieta \& Ettmer, 2007). After this the flood plain widens to approximately 4 km , in this section it is bordered by dikes on the west side and the Panamerican Highway on the east.

According to Alvarado and Ettmer (2007), the sinuosity of the river should be greater than what it currently is considering the slope and the annual average flow, they conclude that the ability of the river to meander is constricted by the dikes (Alvarado Ancieta \& Ettmer, 2007) as can also be seen from satellite images where the river frequently meets the dikes as shown in Figure 19. In the years 1982 and 1997, the dikes were breached at these points (Alvarado Ancieta \& Ettmer, 2007).


Figure 19- Dike breach points in 1982 and 1997 (adapted from Alvarado Ancieta \& Ettmer, 2007)

The morphology of the river is affected by flow through sediment transport. Typically flood events cause general erosion of approximately 6 m (Reyes Salazar, 2024b). However, in 2017 the flow had insufficient energy to cause transport of sediments. This decrease in energy is a combination of increased vegetation, breakwater effects by the dikes, and the redirection of the river towards the lagoons (Reyes Salazar, 2024b). The original course of the river had a length of approximately 70 km starting at Los Ejidos Dam towards Sechura. However, after the extension towards the south, the river has a length of 118 km . This corresponds to an increase of approximately 50 km of length across flat terrain. As a result, the river is incapable of controlling the large quantities of water which flow during El Niño event in the same manner as in previous years (UDEP \& UNP, 2001e).

## 3. Previous Studies

### 3.1. Studies Concerning the Piura River Flow Regimen

As mentioned, the Piura region is among the most vulnerable during heavy rainfall seasons (ARCC, 2017). As such, this section describes various studies and projects which have been conducted and/or are ongoing related to the flooding of the Piura River.

Among the first and most prominent studies concerning is the "Definitive Study for Reconstruction and Rehabilitation of Defenses in the Lower Piura Basin and Mathematical Model for Prediction of Floods in Real Time in the Piura Basin. Piura, Peru." developed by the Class-Salzgitter consortium in 2001, as well as the "Study for the Integral Plan of the Piura River" developed by the National University of Piura together with the University of Piura also in 2001.

The Class- Salzgitter study focused on the reconstruction and rehabilitation of riverine defenses in the lower Piura basin. The study includes design flows suggested for riverine defenses in different zones along the river, description of sediments in the riverbed, and investigates sediment transport (ClassSalzgitter, 2001 cited in Córdova Elera, 2020). The study shows there are sections in the river experiencing erosion while others tend to aggradation (Class-Salzgitter, 2001 cited in Córdova Elera, 2020). The floodplains in the lower basin were identified as irregular areas with vegetative cover of crops and brush, causing a resistance to flow which has resulted in an increase in the level of the floodplain (Class-Salzgitter, 2001 cited in Córdova Elera, 2020).

The study by the National University of Piura and University of Piura (PNUD study) is an analysis of the parameters which influence the frequency and impact of floods and proposes control measures to reduce the negative impact
of floods and increase riverine security with a focus on the urban Piura and Castilla areas (UDEP \& UNP, 2001e). Part of the study includes hydrological investigation for return periods of precipitation and floods, a sediment study, an evaluation of the state of riverine protection in the study year, a numerical and physical model.

Various other studies exist related to the Piura River. These include investigations of flow conditions and precipitation patterns for classification of return periods of flows through the river, vulnerability studies, and hydraulic modeling of the river.

### 3.2. Projects in the Piura River Basin

One of the most important projects in the Piura River basin is the Chira-Piura Project. It was established in 1981 as a self-governed entity in charge of carrying out a project with the aim of storing and regulating hydraulic resources in the Chira River (UDEP \& UNP, 2001a), which is located to the north of the Piura River. The objectives were to improve and stabilize agriculture in the lower basin, this was accomplished through the construction of reservoirs, irrigation channels, and drainage channels (UDEP \& UNP, 2001a). In the Piura basin, the Chira-Piura Project included the construction of Los Ejidos dam, the Biaggio Arbulu channel, reinforcement of dikes, and drainage channels (UDEP \& UNP, 2001a).

In more recent years following the 2017 flood, works for reconstruction and rehabilitation began for both riverine structures and cities which suffered damage due to the flood. Many projects were executed by the Authority for Reconstruction with Change (ARCC). The objective was to lead an integral plan for rehabilitation for the 13 regions in Peru which were most impacted by the coastal El Niño phenomenon of 2017 (ARCC, 2017). Interventions by the

ARCC can be grouped into three two types, the first being rehabilitation of infrastructure such as roads, sidewalks, sewers, and health and educational establishments (ARCC, 2017). The second is related to preventative measures such as clearing riverbeds, constructing riverine defense systems, urban drainage, and conducting studies in critical basins (ARCC, 2017). The investment necessary for interventions was an estimated $\mathrm{S} / 26,655$ million where $77 \%$ was dedicated to the first project type, and $21 \%$ to the second (ARCC, 2017). In the Piura region, works directly related to the river included clearing the riverbed and reinforcement of river defenses. These interventions accelerated in preparation for the potential El Niño event which was likely to occur in the summer of 2023/24.

Similar to the works conducted by the ARCC, the National Water Authority (ANA) also planned interventions for the potential El Niño event of 2023/24 for seven vulnerable regions in Peru with an investment of approximately S/975 million (Autoridad Nacional del Agua, 2023). In Piura the authority identified 94 critical points which needed intervention and undertook works to clear the riverbed of debris and sediment (Autoridad Nacional del Agua, 2023).

Further works are being undertaken to redirect the river towards the south to exit through Reventazón. This is a part of the ARCC's "Integrated Master Plan for Flood Control and Sediment Transport Management in the Piura River Basin", where a relief channel will be built between Chutuque and Pampa Las Salinas as shown in Figure 20 to improve flow conditions in the lower basin (ARCC, 2022).


Figure 20- Location of Chutuque channel (Google Earth, 2024)
There are also alternative proposals for river redirection and reservoirs being studied. Among these are various proposals suggested by Eduardo Zegarra. In one proposal Zegarra suggests the construction of a reservoir to the east of the city of Piura to redirect water around the city and minimize strain on the Los Ejidos Dam as shown in Figure 21.


Figure 21- Coscomba Reservoir map (Zegarra Dávila, 2023)

Further, Zegarra presented various potential routes during the Conference for Hydraulic Structures during Flood Events organized by the Colegio de Ingenieros del Peru (Chapter of Civil Engineering of Peru) in 2023. The routes are shown in Figure 22. For Option A2 (yellow) through Sechura, a pilot channel has already been designed to compare to the planned route through Reventazón. Zegarra estimates that the costs for the exit through Sechura are approximately $60 \%$ of the investment needed for the exit through Reventazón (Zegarra Dávila, 2023).


Figure 22- Potential routes for Piura River discharge (Zegarra Dávila, 2023)
In this thesis, option A2 (yellow) through Sechura is compared to the current exit through Virrilá. This proposed route was chosen after a virtual meeting with Zegarra during the time in Piura. This proposal already has design parameters since the aforementioned pilot channel was used for cost estimation. As such, a comparison of this route to the existing route using river modelling is a further step to determine the effectiveness of the proposal. The upcoming Chapter 4 describes a field visit following the proposed river course through Sechura.

## 4. Field visit

In order to create a better understanding of the study area and the Piura River basin, field visits were conducted. Appendix A provides a larger collection of images and information collected during the excursions. This section provides insight into information gathered which is relevant to sediment transport and the alternative river course.

The first excursion entailed a trip following the urban reach of the river from the Los Ejidos dam to the Bolognesi bridge as shown in Figure 23. The field visit to the urban river reach was a part of the University of Piura`s master's program for Water Resources. The excursion was conducted together with participants from the master's program and was led by Jorge Reyes Salazar, lecturer in hydraulics and sediment transport.


Figure 23a- Excursion to urban river reach. 23b- Overview Map. (Google Earth, 2024)

The excursion began at Los Ejidos Dam located approximately 3km upstream of the city center. The dam is shown in Figure 24 where a weir structure on the left and the gated spillway on the right are visible. The flow through the river on this day was low and consisted primarily of water supplied by the Chira River to the north.


Figure 24- Los Ejidos dam spillway
Sediment in this area was found to be mostly sand with clayey-sand deposits closer to the riverbank. At the dam works have been conducted to remove approximately 3 m of accumulated sediment as shown in Figure 25.


Figure 25-Sediments at Los Ejidos dam. 25a-View form bottom. 25b-View form top

The excursion continued through the city where stops were made at different bridges. As mentioned, the river is channeled in the city and the riverbanks are made of concrete. Aggradation is visible along the riverbanks as shown in Figure 26. Further, the areas where sediment has been deposited promote vegetative growth.


Figure 26- Vegetative growth at riverbanks. 26a-View downstream of Cáceres. 26bBelow Cáceres Bridge. 26c- View upstream of Bolognesi. 26d- Below Bolognesi Bridge

The second visit covered the rural reach of the river starting at the Bolognesi Bridge and following the west (right) dike up to the lagoon system, then following the alternative route towards Sechura as shown in Figure 27. It should be noted that the span along the west dike is not a part of the ancient route but a part of the proposed route. This span provided insight on vegetation cover, irrigation channels, and riverine structures in the lower basin.


Figure 27a-Excursion to the route through Sechura. 27b- Overview map. (Google Earth, 2024)

The land use starting at the Bolognesi bridge and reaching up to the lagoon system consists primarily of farmland with or without crops and brush as shown in Figure 28. There are fields both inside and outside the floodplain.


Figure 28a-Field without crops. 28b- Dense brush

General aggradation is visible since the level of the floodplain was often observed to be above the level outside of the dikes. However, in some spans erosion is visible. This can be seen in Figure 29, taken close to Drain 13.08


Figure 29- Erosion close to Drain 13.08

The current river course towards Virrilá was followed just upstream of the lagoon system. At this point, a new route towards Sechura was taken to assess land coverage conditions along the ancient river course. The Piura River used to flow west towards Sechura. In this area, the ancient river course is still active and is fed by drainage channels. This river is referred to as Río Loco and is shown in Figure 30.


Figure 30- Rio Loco (Google Earth, 2024)

Following the dike on the right side of the river, The Ramón/Napique Lagoon is bordered by the dikes as shown in Figure 31. The vegetation differs upstream and downstream of the dike, where the surroundings upstream have more greenery than downstream.


Figure 31a- Ñapique Lagoon dike (Google Earth, 2024). 31b- Upstream of dike.
31c- Downstream of dike

Following this dike, the terrain has less vegetation and becomes more of a desert landscape as shown in Figure 32.


Figure 32- Vegetation between Napique Lagoon dike and Río Loco

Río Loco is shown in Figure 33. It passes through Sechura and exits into the Pacific in 3 branches.


Figure 33- Río Loco

## 5. Theoretical Background

This section describes theory used to compare sediment transport capacity through the current river course (Virrilá), and the alternative course (Sechura). Piura River. Further, an introduction to the software HEC-RAS provides insight into data needed for modelling.

### 5.1. Sediment Transport

A river is in equilibrium when the amount of sediment being transported in is also transported out (Lane, 1954). According to Lane (1954), for a river to be in complete equilibrium is a rare condition and may only exist momentarily. This is because discharge through a river is not generally constant, and the ratio of high to low flow quantities is very large (Lane, 1954). However, for engineering purposes, equilibrium state is often assumed for a small section of a river although slow aggradation or sedimentation processes may be ongoing (Lane, 1954).

Particles begin to move along the bed when the existing fluid force is larger than the resisting force (Larson, 2023). The critical shear stress is also the resisting force. In this case, Soulsby is used for non-dimensional critical shear stress for non-cohesive sediment using Equations 1 and 2.

$$
\begin{gather*}
\theta_{c r}=\frac{0.3}{1+1.2 D_{*}}+0.055\left(1-\exp \left(-0.02 D_{*}\right)\right)  \tag{1}\\
\text { (2) } \quad D_{*}=\left(\frac{g(s-1)}{v^{2}}\right)^{1 / 3} * d_{50} \tag{2}
\end{gather*}
$$

The critical shields number is given by Equation 1. This is used to find the critical shear stress is $\tau_{\mathrm{cr}}$ in $\mathrm{N} / \mathrm{m}^{2}$ as shown in Equation 3.

$$
\text { (3) } \quad \theta_{c r}=\frac{\tau_{c r}}{\rho g(s-1) d_{50}}
$$

The shear stress created by flowing water $\left(\tau_{\mathrm{b}}\right)$ is calculated using a drag law based on mean velocity (Larson, 2023) as shown in Equation 4.
(4) $\tau_{b}=\rho C_{D} U^{2}$

Transported sediment is classified into bed load, suspended load, and wash load (Larson, 2023). These are defined as follows:

Bed load: Particles are in contact with the riverbed and are transported along the bottom. The shear stress at the bed governs the transport of sediments (Larson, 2023). This can be calculated through numerous equations. In this thesis, Meyer-Peter and Müller is considered.

Suspended load: Particles are carried in the water column through turbulence in the water. Sediment concentration profiles and current velocity are used to describe the load (Larson, 2023).

Wash load: The wash load consists of very fine particles which are evenly distributed through the water column. These are not influenced by the flow properties (Larson, 2023)

### 5.2. The HEC-RAS Model

HEC-RAS is a river analysis tool developed by the Hydrologic Engineering Center of the U.S. Army Corps of Engineers. This free software allows for one- and two-dimensional flow analysis, steady and unsteady flow calculations, as well as sediment transport and water quality analysis (HEC, 2023). For this thesis, a one-dimensional flow model is set up using steady flow conditions.

Basic profile computations in HEC-RAS are conducted with the standard step method where the Energy Equation is used as shown in Equation 5 to calculate
water surface profiles from cross section to cross section as shown in Figure 34 (HEC, 2024a).
(5) $Z_{2}+Y_{2}+\frac{\alpha_{2} V_{2}^{2}}{2 g}=Z_{1}+Y_{1}+\frac{\alpha_{1} V_{1}^{2}}{2 g}+h_{e}$


Figure 34- Water surface profile calculation in HEC-RAS (HEC, 2024a)

Further, conveyance, kinetic energy, friction loss, and contraction/expansion losses are computed using various equations and input in form of cross section geometry and Mannings numbers (n) (HEC, 2024a). The program uses an iterative process to calculate water surface using predefined error tolerances (HEC, 2024a). Herein the Mannings $n$ values has a considerable impact on the computed water surface (HEC, 2024a). Mannings $n$ is a is an indication of the roughness of a terrain. Smooth surfaces have a low Manning's n, while rough surfaces have a high n value.

The following assumptions are used in one-dimensional steady flow analysis according to the HEC-RAS Hydraulic Reference Manual (2024):

- Flow is steady- energy equation does not include time dependent terms.
- Flow is gradually varied- hydrostatic pressure distribution is assumed.
- Flow is one dimensional- velocity components are assumed to be in the same direction as flow under the assumption that total energy head is the same across a cross section.
- Channels have slopes less than 1:10 which implies that vertical pressure head is approximately equal to the depth of the water

Steady flow computations require boundary conditions, flow regime, and discharge information (HEC, 2024a). Flow regime can be defined as subcritical, supercritical, or mixed flow (HEC, 2024a). Boundary conditions can be specified as known water surface elevations, critical depth, normal depth based on the energy slope, or a rating curve where elevation-flow data is entered (HEC, 2024a). Finally, at least one flow value is needed upstream in the river reach (HEC, 2024a).

## 6. Data Employed

In this section data relevant to the HEC-RAS model set up and sediment transport capacity calculation is described. This includes available flow data, sediment information, observations for Mannings roughness coefficients, topographic data, sea level and a description of the proposal for the course through Sechura provided by Eduardo Zegarra.

### 6.1. Flow data

As mentioned, the Piura River is non-perennial. The flow through the lower basin is heavily influenced by the presence of El Niño due to the increased precipitation during this weather event. Data for maximum flow through the river has been available since 1926 and is shown in Figure 35. The trendline indicates that the flow through the river is increasing over time.


Figure 35- Maximum flows through the Piura River from 1926-2017 (Farias Zegada, 2008)

Data shows that following 1983, maximum flows were larger and more frequent. Statistical analysis of the flows has been conducted in previous studies using different models such as Gamma, Exponential, Person III,

Gumbel, and others. Farías (2018) uses the double Gumbel model to estimate flows. Table 1 shows the maximum flow for different return periods in this double Gumbel investigation. It should be noted that varying values can be found in different literature.

Table 1- Qmax return periods

| $\operatorname{Tr}$ (years) | Max. flow $\left(\mathbf{m}^{\mathbf{3}} / \mathbf{s}\right)$ |
| :---: | :---: |
| 2 | 432 |
| 5 | 2,081 |
| 10 | 2,954 |
| 25 | 4,014 |
| 50 | 4,788 |
| 100 | 5,552 |

Further, the Piura-Chira Project provides daily flow data since 2015 for five stations along the Piura River (PECHP, 2017). The stations provide flow data for daily average, maximum flow, and precipitation. Of these, the station at the Sánchez Cerro bridge provides flow information for the lower basin. The flood event in March 2017 is shown in Figure 36. The hydrograph shows a peak flow of $3,468 \mathrm{~m}^{3} / \mathrm{s}$ on March $28^{\text {th }}$.


Figure 36- Flow measurements at Sánchez Cerro Bridge for March 2017

### 6.2. Sea Level

Hourly sea level data from Talara- a port to the north of Piura- was sourced from the University of Hawaii Sea Level Center. Data is available from 31.05.2015-30.04.2023, with some missing data, as visible in Figure 37. The mean sea level is found to be at approximately 1.5 m with deviations of $\pm 1.2 \mathrm{~m}$, resulting in a maximum sea level of 2.7 m and a minimum of 0.2 m .


Figure 37-Sea Level at Talara, Peru from 31.05.2015-30.04.2023

### 6.3. Sediment data

The study primarily referred to is the study by the National University of Piura and the University of Piura (PNUD study) from 2001 for sediment data. This expansive study covers information relevant for modeling and design of riverine structures. The study gives insight into the geology of the area and provides results for geotechnical evaluations of several samples taken along the river.

The most prominent tertiary formation in the study area is the Zapallal formation (UDEP \& UNP, 2001e). In the year this report was written, bedrock (Zapallal formation) was encountered during sediment sampling at Los Ejidos dam, Cáceres Bridge, and Sánchez Cerro bridge (UDEP \& UNP, 2001e). The PNUD study mentions that this exposure of bedrock is largely attributed to anthropogenic changes which have influenced the flow regime and caused an increase in sediment transport. Bedrock depth is found to range between 0.514.5m (UDEP \& UNP, 2001e).

Quaternary sediments are identified as alluvial, fluvial, and eolian sediments (UDEP \& UNP, 2001e). Alluvial sediments are found to be coarse sand with fine sand on the top, followed by layers of sand and clay (UDEP \& UNP, 2001e). Fluvial sediments are silty-sand sediments and are used for agriculture (UDEP \& UNP, 2001e). Following these layers are the eolian sediments which have been transported by wind and form small dunes in some areas (UDEP \& UNP, 2001e). The study provides results for soil tests such as grain curves, determination of specific weight, water content, shear strength, and Atterberg limits for samples taken in the riverbed starting at Los Ejidos and ending at the Ramón Lagoon.

The following results points are highlighted from the PNUD study:

- Most samples ranged from silty-sand to fine-sand, however a clay sample was also found in the Ramón Lagoon sample.
- Median sediment grain size (d50) ranges from 0.1-0.3mm
- The specific weight of sediments ranges from $2.55-2.8 \mathrm{t} / \mathrm{m}^{3}$.

Further, the PNUD report references another study from 1967 in which information for sediment concentration as a rating curve for suspended sediment in tonnes/month versus monthly discharge (Int. Engineering Company \& Olazabal y Leon S.A, 1967).

### 6.4. Manning

Based on observations from the field visit described in Chapter 4, a possible range of values is determined for roughness. Table 2 shows potential Mannings numbers along the river ranging from 0.014 (concrete) to 0.16 (dense brush) based on value suggestions provided by the HEC-RAS Hydraulic Reference Manual (2024).

Table 2- Land cover roughness, Mannings numbers based on observation

| Element | Description | Range | Example |
| :---: | :---: | :---: | :---: |
| Lined channel | Unfinished concrete | 0.014-0.02 |  |
| Riverbed | Sandy riverbed, short grass, few weeds | $\begin{gathered} 0.022- \\ 0.033 \end{gathered}$ |  |
| Floodplain | Cultivated areas, crops | $\begin{gathered} 0.025- \\ 0.045 \end{gathered}$ |  |
|  | Pasture, no brush, short grass | $\begin{gathered} 0.025- \\ 0.035 \end{gathered}$ |  |
|  | Scattered brush, heavy weeds | 0.035-0.07 |  |
|  | Dense brush | 0.07-0.16 |  |

### 6.5. Topography

Elevation data is necessary for the area stretching from Los Ejidos Dam to Virrilá and including Sechura as shown in Figure 38. This is the lower basin and has an area of approximately $4,000 \mathrm{~km}^{2}$. The river spans 117 km from Los Ejidos to the Virrilá Estuary, and 72 km from Los Ejidos to Sechura.


Figure 38- Study area (Google Earth, 2024)
Topographic data is used for geometry creation in HEC-RAS. For this data from a 2015 LIDAR scan by the national water authority (Autoridad Nacional del Agua) was used. The scan covers a length of 50 km starting just upstream of Los Ejidos Dam and reaching to the Napique/Ramón Lagoon. The width varying from 1.9 km to 7.2 km to include the floodplains and has an area of approximately $165 \mathrm{~km}^{2}$ (Peña Valdivia, 2021) the scanned areas are shown in Figure 39. Data from the scan has been preprocessed by students from the University of Piura in previous master thesis studies to create a raster file. This file was used as a GeoTIFF format as input in HEC-RAS, the result is shown in Figure 39.


Figure 39a-LiDAR scans by ANA in 2015 (Peña Valdivia, 2021). 39b-GeoTIFF file opened in RAS-Mapper

The study area, however, is larger than what is provided in the LIDAR scan. As such, digital elevation model (DEM) data was obtained through GEO GPS Peru (n.d.), where it was available in GeoTiff format with a resolution of 30 m taken in 2008. The file is shown in Figure 40.


Figure 40- ASTER GDEM Piura region

### 6.6. River Course through Sechura

The proposed river course through Sechura has a length of approximately 42 km while the course through Virrilá has a length of approximately 87 km from the split point. The new course follows the right dike towards the Ñapique/Ramón Lagoons, then passes through the dike as shown in Figure 41 and rejoins Río Loco after 24 km of the split point.


Figure 41- Course through Sechura (Google Earth, 2024)
For the redirection towards Sechura, a pilot channel is to be constructed with a depth of approximately 4 m and capacity for $375 \mathrm{~m}^{3} / \mathrm{s}$ (Zegarra Dávila, 2023). The pilot channel is proposed for the reach from the split point up to the connection to Río Loco. Río Loco is estimated to have a flow capacity of $3,200 \mathrm{~m}^{3} / \mathrm{s}$ considering that this was the flow of the flood in 1983, and the river contained this flow following the breach of the dikes (Zegarra Dávila, 2023). The small flow capacity is attributed to the idea that this course will have sufficient transport capacity to flush out sediment and thereby begin to widen
overtime without further construction (Zegarra Dávila, 2023). The general section of the proposed pilot channel is shown in Figure 42.


Figure 42- General cross section of the pilot channel through Sechura (Zegarra Dávila, 2023)

Table 3 shows a more detailed description of the pilot channel based on the typical section shown in Figure 43. Station 0+000 is the split point and $24+000$ is the connection to Río Loco. In this, the slope of the pilot channel varies depending on the existing terrain.

Table 3- Parameters for the pilot channel

| From <br> (stat.) | To <br> (stat.) | $\mathbf{L}$ <br> $(\mathbf{m})$ | $\mathbf{B}$ <br> $(\mathbf{m})$ | $\mathbf{Z}$ <br> $(\mathbf{m})$ | $\mathbf{Q}$ <br> $\left(\mathbf{m}^{3} / \mathbf{s}\right)$ | V <br> $(\mathbf{m} / \mathbf{s})$ | D <br> $(\mathbf{m})$ | $\mathbf{S}$ <br> $(\mathbf{m} / \mathbf{m})$ |
| :--- | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| $0+000$ | $6+000$ | 6000 | 50 | 2 | 300 | 1.19 | VAR | 0.0002 |
| $6+000$ | $18+400$ | 12400 | 50 | 2 | 300 | 1.5 | VAR | 0.0004 |
| $18+400$ | $22+000$ | 3600 | 120 | 2 | 300 | 1.59 | VAR | 0.0006 |
| $22+000$ | $24+000$ | 2000 | 30 | 2 | 300 | 1.31 | VAR | 0.0005 |



Figure 43- Parameters for the pilot channel (Zegarra Dávila, 2023)

## 7. HEC-RAS Model Set Up and Output Processing

HEC-RAS Version 6.3 .1 was used for the hydraulic model. This chapter describes the setup, parameters, and assumptions used for assessment of the hydraulic conditions in the two river courses. The model is set up as onedimensional.

### 7.1. Concept Model

A HEC-RAS model is set up considering elements indicated in Figure 44. Boundary conditions are set in the upstream and downstream sections of the river. These define the behaviour of water entering and exiting the system. The flow through the model is partially limited by dike structures on the left and right as well as the Panamerican Highway which acts as a boundary. The dikes are described as partial boundaries because although they do control flow, there are conditions under which the dikes will be breached. The cross sections are drawn perpendicular to the riverbed and span across the dikes to include the floodplain. Since flood analysis is not the focus of this study, the cross sections will be drawn just beyond the dikes, or just across the riverbed/lagoon in areas where dikes are non-existent.


Figure 44- Concept model

In order to run a flow model in HEC-RAS two files are needed; these are a geometry file and a flow file which are combined in a plan file for computation as shown in the scheme in Figure 45.


Figure 45- HEC-RAS file setup

### 7.2. Geometry File

Geometry files for one-dimensional models in HEC-RAS can be set up through manual input of cross section data or through RAS-Mapper. RASMapper is a tool with geospatial capabilities to visualize geometric data and simulation results (HEC, 2024b).

A combination of the two files described in Chapter 6.5 is used to create cross sections in RAS-Mapper. General handling of the files and notable observations are described below:

- While the LIDAR file shows relevant features more clearly, the ASTER file covers a larger area.
- The LIDAR file requires some manipulation for error removal and simplification.
- Although the riverbed is clear in the LIDAR file, the topography data shows elevation differences in large distance steps of approximately 1 m . As such the cross sections appear step-like. Nevertheless, considering that the general shape of the sections is clear, it has been deemed appropriate for general hydraulic analysis.
- Features in the ASTER file, such as the riverbed and dikes are not so clear in the ASTER file. However, the file is deemed sufficiently accurate to estimate the slope and approximate shape of the terrain.
- Cross sections created using the ASTER were partially manipulated to create an approximation of the riverbed for both river courses.
- It was found that the elevations of the files are similar in the transition from one to another.

The river centerline has the same shape for both courses up to just downstream of Drain 13.08, after this the geometry differs for both courses. Cross sections are spaced at approximately 500 m up to Drain 13.08 , then spacing increases
to approximately $1,000 \mathrm{~m}$ up to the end of the LIDAR topography file. Finally, spacing of cross sections created using the ASTER file are spaced at varying distances. Figure 46 shows the geometry file created using RAS-MAPPER.

Detailed spacing and cross section data is provided in Appendix B.


Figure 46- LIDAR topography and ASTER Topography opened in RAS-Mapper

HEC-RAS generates river stationing based on the distance from the most downstream point of the river centerline. As such, the stationing for the courses differs. While Virrilá reaches from station $1+568$ downstream to $117+307$ upstream, the geometry for the river through Sechura has a downstream river station of $1+870$ and upstream station of $71+617$.

## Upstream of Drain 13.08

This section of the geometry spanning from Los Ejidos to Drain 13.08 is the same for both river courses. The section spans from 117+307 to $87+776$ and from $71+617$ to $42+086$ for Virrilá and Sechura respectively. Cross sections comprise of bank stations, levees and ineffective areas. The riverbanks are approximately placed following the main riverbed. Levee structures are used along the top of the channel in the urban zone as shown in Figure 47 and on top of the dikes in the rural areas as shown in Figure 48.


Figure 47- Reach Virrilá: cross section 117+307


Figure 48- Reach Virrilá: cross Section 103+416
In sections where water is likely to overflow the levees or where flooding is expected to happen, ineffective areas were used outside of the dikes as shown in Figure 49. This ensures that water outside of the floodplain is not contributing to results.


Figure 49- Reach Virrilá: cross section $88+621$

## Reach to Virrilá

The existing reach to Virrilá follows the same set up as previously described reaching up to the Napique/Ramón Lagoon. Levees are used to mark high ground or existing dikes and riverbanks are defined following the main river course. After this the ASTER topography is used and cross sections are drawn to span the width of the riverbed or lagoon. As shown in Figure 50, some features such as high ground, dunes, and the Panamerican highway are visible in the ASTER data. However, the riverbed and estuary are not clear.


Figure 50a- Virrilá Estuary satelite imagery (Google Earth, 2024). 50b-area of Virrilá Estuary in ASTER data.

The Virrilá estuary is influenced by tidal activity up to approximately 20 km upstream and is estimated to have an average depth of 2 m (Céspedes, 2005). Since detailed bathymetric data was not available for the estuary, the bottom level in the last cross section $1+568$ was calculated to be at -1 m based on an assumed slope of $0.2 \%$ along the estuary. The topography of cross sections along the estuary was modified using this slope of $0.2 \%$. Similarly, due to the uncertainty in accuracy of the topography in the lagoon area, elevations were adjusted between riverbanks to create a gradual slope towards the estuary as shown in Figure 51 and Figure 52 for a schematic model.


Figure 51- Reach Virrilá: Profile before adjustment


Figure 52- Reach Virrilá: Profile after adjustment

## Reach to Sechura

The reach towards Sechura is modelled based on the alternative course description in Chapter 6.6 as well as some assumptions. These are as follows:

- The topography through which the channel passes is adapted to the slope defined in Table 3.
- The slope of the cross sections at Río Loco are modelled with an assumed slope of $0.3 \%$. to reach a bottom level of around 0 . A shallower water level than that at Virrilá is assumed due to observations during the field visit.
- The pilot channel capacity is assumed to increase from $300 \mathrm{~m}^{3} / \mathrm{s}$ to $3,200 \mathrm{~m}^{3} / \mathrm{s}$ to match the capacity of Río Loco.
- It is assumed that a flow of $3,200 \mathrm{~m}^{3} / \mathrm{s}$ will be contained by the channel, as such levees are placed on the top of the inserted channel.
- It is assumed that the Napique/Ramón Lagoon will still fill with water in this river course.
- The area outside of the existing dikes is modelled as ineffective zones.

Figure 53 shows a sample section of the channel along the right dike for topography data from the LIDAR scan. The depth of the channel has been assumed to be 4 m and the riverbanks have a slope of approximately $1: 2$. Similar sections are placed on the ASTER topography as shown in Figure 54.


Figure 53- Reach Sechura: cross Section 40+847


Figure 54- Reach Sechura: cross section 18+624

## Manning

Mannings numbers are set for the riverbed and overbanks. As described in Chapter 6.4 the terrain is defined by varying roughness coefficients. Figure 55 indicates approximately where different roughness coefficients are used. The $n$-value is adjusted during calibration as described in the following section considering the following ranges:

- Riverbed: 0.022-0.033
- Concrete: 0.014-0.02
- Crops/Scattered Brush: 0.025-0.07
- Dense Brush: 0.07-0.016
- Pasture: 0.025-0.035


Figure 55-Roughness coefficient regions

### 7.3. Flow and Plans

Steady flow analysis is used for comparison of the river courses. Steady flow data in HEC-RAS requires upstream and downstream boundary conditions. A sensitivity analysis is conducted to investigate the influence of downstream boundary conditions on velocity. For this, the conditions described in Table 4 are used on a flow of $\mathrm{Q}=1,000 \mathrm{~m}^{3} / \mathrm{s}$ through Virrilá.

Table 4- Boundary condition in HEC-RAS for sensitivity analysis

| Boundary Condition | Value | Description <br> Normal depth 0.0003 m/m |
| :--- | :--- | :--- |
| Depth is calculated using the <br> Manning's Equation based on the <br> friction slope (approximately <br> equal to the surface slope) (HEC, <br> 2024a) |  |  |
| Known water surface | 2.7 m | The maximum sea level from <br> Chapter 6.2. |
| Known water surface | 1.5 m | The mean sea level from Chapter <br> 6.2 |
| Known water surface | 0.2 m | The minimum sea level from <br> Chapter 6.2 |
| Normal depth + known <br> water surface | $1.0003 \mathrm{~m} / \mathrm{m}$ <br> The friction slope is set as a <br> boundary condition with an <br> additional known internal water <br> surface of 1.5m set at the <br> downstream-most cross section |  |

As shown in Figure 56, it was found that the boundary condition does not have an impact on upstream stations. The effect of the boundary condition is noted up to approximately 15 km from the end of the reach. As such, the known water surface of +1.5 m is chosen for simulations as this is the average sea level in the region.


Figure 56- Reach Virrilá: Sensitivity analysis for boundary conditions at $Q=1,000 \mathrm{~m}^{3} / \mathrm{s}$

Flows for assessment of hydraulic conditions were chosen as $\mathrm{Q}=3,200 \mathrm{~m}^{3} / \mathrm{s}$ and $\mathrm{Q}=1,000 \mathrm{~m}^{3} / \mathrm{s}$ and are used as upstream boundary conditions. Since the capacity of Río Loco is estimated to be $\mathrm{Q}=3,200 \mathrm{~m}^{3} / \mathrm{s}$, this flow is identified as a maximum flow for the alternative course. Larger flows are not expected because it is assumed that flows above $3,200 \mathrm{~m}^{3} / \mathrm{s}$ will be redirected towards La Niña Lagoon, this spillover concept is not within the scope of the model. Further, a flow of $1,000 \mathrm{~m}^{3} / \mathrm{s}$ is used for comparison as it is assumed that this is likely to be a recurring maximum flow according to the maximum flows shown in Figure 35 in Chapter 6. Both flow scenarios are computed assuming sub critical conditions due to the small slope along the course of the river.

### 7.4. Calibration

Before computation of the above-mentioned flow, the model is calibrated using known water surface elevations for a given flow. In this model, the Virrilá reach is calibrated using data from the 2017 flood event. Although the LIDAR topography is from 2015, it is assumed that calibration using 2017 data is appropriate since there were no major floods between the years 20152017 thereby no major changes in the topography are expected. A peak flow of $3,468 \mathrm{~m}^{3} / \mathrm{s}$ was measured in March 2017 (PECHP, 2017), the water surface elevation was measured for several bridges as shown in Table 5.

Table 5- Maximum water surface for flooding of 2017 (Maza-Sócola, 2019)

| Location | Station and downstream <br> distance (Virrilá Geometry) |  | Water surface <br> elevation (m asl.) |
| :---: | :---: | :---: | :---: |
| Cáceres Bridge | $114+881$ | 128 | 32.30 |
| Sánchez Cerro <br> Bridge | $113+897$ | 392 | $30.74^{(*)}$ |
| Bolognesi Bridge | $112+930$ | 230 | 29.70 |
| Grau Bridge | $108+711$ | 553 | $28.21^{(*)}$ |
| Independencia <br> Bridge | $97+650$ | 612 | 24.55 |

(*) estimated values (Reyes Salazar cited in Maza-Sócola, 2019)

However, additional information is needed to describe the flow regimen further downstream. Since exact water surface elevations are not known, satellite imagery was used to assess the extent of flooding. NASA's EODIS Worldview provides historic satellite images as shown in Figure 57 for the $5^{\text {th }}$ of April 2017. Although the peak flow occurred on March $28^{\text {th }}$, previous dates showed significant cloud coverage and could not be used for assessment. The resolution of the image is not high; however, the size of the lagoons can be estimated, and the image gives insight into whether flooding occurred in certain regions.


Figure 57- Extent of flooding in 2017. Satellite imagery from April 5th, 2017 (adapted from Nasa Worldview, 2017)

Mannings $n$-values are adjusted based on observed water surface elevations and flood extent using a flow of $3,468 \mathrm{~m}^{3} / \mathrm{s}$. The results of calibration are shown in Chapter 8.

### 7.5. Output Tables

The aim of the simulations is hydraulic comparison as well as an investigation of sediment transport capacity. HEC-RAS output is used for the whole cross section as opposed to analyzing main channel and overbanks separately, this is deemed sufficient for analysis of the general flow regime. Hydraulic depth $(\mathrm{m})$, total velocity $(\mathrm{m} / \mathrm{s})$, and total shear $\left(\mathrm{N} / \mathrm{m}^{2}\right)$ are used as output from HECRAS.

## 8. Results

This chapter shows results for calibration, velocity and depth profiles, as well as sediment transport capacity.

### 8.1. Calibration

Mannings n-values were adjusted according to the ranges described in Table 2 in Chapter 6. The final n-values are listed in Appendix B. Calibration results of the measured data from 2017 are shown in Table 6, the error found for each station was deemed acceptable.

Table 6- Calibration results for known water surface

| Location | Obs. Water <br> Surface | Sim. Water <br> Surface | Error (\%) |
| :---: | :---: | :---: | :---: |
| Cáceres Bridge | 32.30 | 32.39 | 0.3 |
| Sánchez Cerro <br> Bridge | $30.74^{*}$ | 30.83 | 0.3 |
| Bolognesi Bridge | 29.70 | 29.49 | 0.7 |
| Grau Bridge | $28.21^{*}$ | 28.43 | 0.8 |
| Independencia <br> Bridge | 24.55 | 24.68 | 0.5 |

(*) estimated values (Reyes Salazar cited in Maza-Sócola, 2019)

As mentioned, no precise measurements were available further downstream. However, satellite imagery showing the extent of flooding in 2017 was used for approximation. Namely, the Mannings numbers were adjusted in some areas to allow flooding to occur, or to keep the water level below the levee structures modelled in HEC-RAS. The calibrated model is shown in Figure 58, where the points on the upstream portion indicate the observed water surfaces listed above.


Figure 58- Reach Virrilá: water surface profile for calibration $Q=3468 \mathrm{~m}^{3} / \mathrm{s}$

### 8.2. HEC-RAS Output: Water Surface Elevation Profiles

Virrilá: Flow $\mathrm{Q}=1,000 \mathrm{~m}^{3} / \mathrm{s}$
Figure 59 shows some increases in depth as between station $25+639$ to $31+691$ and stations $62+057$ to $65+336$, which is expected as these are the areas between the lagoons where the river is narrower than the lagoon.


Figure 59- Reach Virrilá: water surface profile for $Q=1,000 \mathrm{~m}^{3} / \mathrm{s}$

## Virrilá: Flow $\mathrm{Q}=3,200 \mathrm{~m}^{3} / \mathrm{s}$

Figure 60 shows the increase in water depth during the larger flow event. Flooding can be seen along some stretches of the river.


Figure 60 - Reach Virrilá: water surface profile $Q=3,200 \mathrm{~m}^{3} / \mathrm{s}$

Sechura: Flow $\mathrm{Q}=1,000 \mathrm{~m}^{3} / \mathrm{s}$
The profile in Figure 61 shows the flow through the course towards Sechura.
The levees indicated in the lower reach here are not dike structures but are used for modelling purposes. The flow shows a steep drop at around station $40+847$, this may be due to the transition from the original river course to the new channel towards Sechura.


Figure 61-Reach Sechura: water surface profile for $Q=1,000 \mathrm{~m}^{3} / \mathrm{s}$

Sechura: Flow Q $=3,200 \mathrm{~m}^{3} / \mathrm{s}$
Figure 62 shows the surface profile for a flow of $3,200 \mathrm{~m}^{3} / \mathrm{s}$. An increase in depth is seen at station $21+369$ as this is the exit of the Ñapique/Ramón Lagoon, since the river is narrower than the lagoon, there is an increase in depth.


Figure 62- Reach Sechura: water surface profile for $Q=3,200 \mathrm{~m}^{3} / \mathrm{s}$

### 8.3. HEC-RAS Output: Velocity and Depth Comparison

Figure 63 and 64 show the velocity and depth variation for $\mathrm{Q}=1,000 \mathrm{~m}^{3} / \mathrm{s}$ and $\mathrm{Q}=3,200 \mathrm{~m}^{3} / \mathrm{s}$ respectively for the course along Virrilá. Velocity and hydraulic depth are average values across each cross section.


Figure 63- Reach Virrilá: velocity and depth output for $Q=1,000 \mathrm{~m}^{3} / \mathrm{s}$

The course through Virrilá shows velocities reaching from $0.12-2.44 \mathrm{~m} / \mathrm{s}$ for a flow of $1,000 \mathrm{~m}^{3} / \mathrm{s}$. Hydraulic depth and velocity is higher upstream due to the channeling of the river and both decrease towards the lagoons, which is expected due to the expansion of the flood plain.


Figure 64- Reach Virrilá: velocity and depth output for $Q=3,200 \mathrm{~m}^{3} / \mathrm{s}$

For a flow of $3,200 \mathrm{~m}^{3} / \mathrm{s}$, a range of $0.16-4.46 \mathrm{~m} / \mathrm{s}$ develops through the river. The same general trend exists where there are larger velocities in the upstream region and very low velocities in the lagoons. However, for this flow, the velocity is approximately twice as large in the peaks. The hydraulic depth relates to the velocity in that when velocity increases, depth decreases and vice versa. This is particularly relevant for narrower reaches such as the urban reach from $117+307-111+954$ but is less notable in the lagoons where larger floodplains exist.

Figure 65 and Figure 66 show the velocity and depth variation for $\mathrm{Q}=1,000 \mathrm{~m}^{3} / \mathrm{s}$ and $\mathrm{Q}=3,200 \mathrm{~m}^{3} / \mathrm{s}$ respectively for the course along Virrilá. Velocity and hydraulic depth are average values across each cross section.


Figure 65- Reach Sechura: velocity and depth for $Q=1,000 \mathrm{~m}^{3} / \mathrm{s}$


Figure 66- Reach Sechura: velocity and depth for $Q=3,200 \mathrm{~m}^{3} / \mathrm{s}$

The course through Sechura shows velocities reaching from $0.46-4.23 \mathrm{~m} / \mathrm{s}$ for a flow of $1,000 \mathrm{~m}^{3} / \mathrm{s}$ and a range of $0.38-4.46 \mathrm{~m} / \mathrm{s}$ for a flow of $3,200 \mathrm{~m}^{3} / \mathrm{s}$. Velocity shows a spike in Figure 65, this is due a transition between cross sections where in one cross section the floodplain is flooded and in the next it is contained by the riverbed. This peak is larger for $\mathrm{Q}=1,000 \mathrm{~m}^{3} / \mathrm{s}$ than that of $\mathrm{Q}=3,200 \mathrm{~m}^{3} / \mathrm{s}$ because the larger flow causes the floodplain to be flooded in both cross sections. However, disregarding the peak, the velocity range for $\mathrm{Q}=1,000 \mathrm{~m}^{3} / \mathrm{s}$ would be $0.46-2.44 \mathrm{~m} / \mathrm{s}$. This is a similar max velocity to that which develops in the reach through Virrilá, however, the minimum velocity is slightly higher.

The upstream velocity is similar in both courses for both flows. Further downstream the course towards Sechura develops higher velocities than the course through Virrilá as shown in Figure 67 and Figure 68 for comparison of the velocity development.


Figure 67- Velocity through Virrilá and Sechura at $Q=1,00 \mathrm{~m}^{3} / \mathrm{s}$


Figure 68-Velocity through Virrilá and Sechura at $Q=3,200 \mathrm{~m}^{3} / \mathrm{s}$

In the next section, velocity and hydraulic depth are used to estimate sediment transport capacity throughout the course of the river.

### 8.4. Sediment Transport Capacity

HEC-RAS output is applied to the formulas described in Chapter 6 for noncohesive sediment along the riverbed. As mentioned, shear stress is calculated using a drag law based on mean velocity which accounts for roughness conditions along the riverbed. Shear stress in each cross section is compared to the calculated critical shear stress of $\boldsymbol{\tau}_{\mathbf{c r}}=\mathbf{0 . 1 4 4} \mathbf{N} / \mathbf{m}^{\mathbf{2}}$ for a sediment of mean grain size $\mathrm{d}_{50}=0.2 \mathrm{~mm}$ and water of an assumed temperature of $25^{\circ} \mathrm{C}$.

HEC-RAS provides a shear stress value for each cross section based on the hydraulic radius and slope of the cross section. Figure 69 and Figure 70 compare the shear force calculated by the program and shear force calculated using the drag law for a flow of $\mathrm{Q}=3,200 \mathrm{~m}^{3} / \mathrm{s}$ for both river courses.


Figure 69- Shear output comparison for Virrilá $Q=3,200 \mathrm{~m}^{3} / \mathrm{s}$


Figure 70- Shear output comparison for Sechura $Q=3,200 \mathrm{~m}^{3} / \mathrm{s}$

The shear stress computed in HEC-RAS is significantly higher than that calculated using the drag law. Figure 71 and Figure 72 show shear stress results for Virrilá and Sechura for calculations based on the drag law and for HEC-RAS output shear stress in relation to the critical shear stress. Values of $\tau_{\mathrm{b}}<\tau_{\mathrm{cr}}$ indicate that the section is stable. For clear visualization of results, the $y$-axis range is reduced from $0-5$ considering that the critical shear stress is a small number, it should be noted that exact shear stress values are listed in Appendix C.


Figure 71- Drag law shear stress and threshold of motion for $Q=1,000 \mathrm{~m}^{3} / \mathrm{s}$ and $Q=3,200 \mathrm{~m}^{3} / \mathrm{s}$ for Virrilá and Sechura


Figure 72-HEC-RAS shear stress output and threshold of motion for $Q=1,000 \mathrm{~m}^{3} / \mathrm{s}$ and $Q=3,200 \mathrm{~m}^{3} / \mathrm{s}$ for Virrilá and Sechura

The route through Virrilá indicates sections where the calculated shear stress along the bed $\tau_{\mathrm{b}}$ is less than $\tau_{\text {cr. }}$ This implies that the threshold for sediment motion is not exceeded and therefor sediment transport is not expected. The course through Sechura indicates some drops in shear stress but remains above the critical shear stress, therefor indicating that the riverbed is not in a stable state and likely to be eroding. HEC-RAS output does not indicate any sections in stability. The HEC-RAS output is generally larger than that of the drag law calculated shear stress. This is due to the difference in the formulas used for computation where HEC-RAS output is calculated using slope and hydraulic radius (HEC, 2024a) while the drag law is based on bed roughness.

Although the shear stress and critical shear stress comparison gives insight into the sediment transport conditions, the excess shear stress is assumed to be proportional to transport. As such, it can be used to approximate in which sections of the river erosion or deposition is likely to occur. As shown in Figures 73-76, the excess shear is larger along most sections of the river towards Sechura as toward Virrilá. For visual purposes, the $y$-axis is set from 0-60 for the drag law computations and from 0-140 for the HEC-RAS output results, full results are listed in Appendix C.


Figure 73- Excess shear based on drag law for $Q=1,000 \mathrm{~m}^{3} / \mathrm{s}$

——Virrilá Excess Shear
.-...... Virrila Velocity -Sechura Excess Shear -...... Sechura Velocity
Figure 74- Excess shear based on drag law for $Q=3,200 \mathrm{~m}^{3} / \mathrm{s}$


Figure 75-Excess shear based on HEC-RAS output for $Q=1,000 \mathrm{~m}^{3} / \mathrm{s}$

_- Virrilá Excess Shear ........ Virrila Velocity _- Sechura Excess Shear ........ Sechura Velocity
Figure 76-Excess shear based on HEC-RAS output for $Q=3,200 \mathrm{~m}^{3} / \mathrm{s}$

Although the shear stress along the course towards Sechura did not go below the critical shear stress, the gradient of excess shear is indicative of erosion or deposition. Erosion occurs when excess shear increases while deposition ocurrs when it decreases. The figures above show various peaks which are likely due to the computation process in HEC-RAS. For analysis of erosion or deposition, trendlines are used to simplify the peaks as shown in Figure 77 for Virrilá and Sechura at a flow of $\mathrm{Q}=3,200 \mathrm{~m}^{3} / \mathrm{s}$ for the drag law computation and HEC-RAS shear output. The trendlines are approximations purely created for visualization and are not based on numerical values.


Figure 77- Trend for excess shear along Virrilá and Sechura

Erosion likely occurs immediately downstream of Los Ejidos Dam, then the excess shear drops for both scenarios indicating likely deposition of sediments. Following this, the route towards Virrila shows some smaller peaks but has a decreasing trend due to the decreased transport capacity in the lower reach. The route towards Sechura shows more defined peaks in the lower reach and a general increasing trend in excess shear. There is however a drop in excess shear in the lower reach which is a result of the assumption that the Ramón/Napique Lagoon still fills with water. This implies an increased width and a drop in velocity, therefor the excess shear also drops. Comparing both
the HEC-RAS output and the calculated shear stress, the general trend for Virrilá is a decrease in excess shear towards the lower reach while increased larger shear is seen in the course towards Sechura.

As mentioned in Chapter 5, excess shear is related to the Meyer-Peter and Müller bed load transport formula. Bed load transport is calculated for both the shear stress based on the drag law and for the HEC-RAS output shear for a flow of $\mathrm{Q}=3,200 \mathrm{~m}^{3} / \mathrm{s}$. As shown in Figures 78 and 79 , bed load in load per unit width shows small peaks for the transport through Sechura while it stays quite insignificant for transport through Virrilá. The y-axis was adjusted to 00.05 for visualization purposes, full data is provided in Appendix C.


Figure 78- Bed load transport for Virrilá at $Q=3,200 \mathrm{~m}^{3} / \mathrm{s}$


Figure 79- Bed load transport for Sechura at $Q=3,200 \mathrm{~m}^{3} / \mathrm{s}$

## 9. Discussion and Recommendations

The following chapter describes the significance of the hydraulic conditions occurring in the different river courses, provides suggestions improvement of the study, and comments on potential studies related to the issue of sediment transport in the lower basin.

### 9.1. Implication of Results

The HEC-RAS model represents the general flow regimen in the two river courses. While conditions upstream of the split point do not vary for different routes, it is evident that the reaches downstream of the split result in conditions that favor sediment transport in the route towards Sechura.

The course towards Virrilá shows spans totaling to approximately 30 km with an excess shear value of 0 using the drag law calculation, and close to 0 using HEC-RAS output, this is an indication of sedimentation. This is expected due to the small slope and resulting low velocities. The lack of transport capacity results in sedimentation in the lagoons which creates an obstacle for flow. Although the lagoons may have a large capacity to hold sediment before filling up, these lagoons are not being filled evenly (Reyes Salazar, 2024b). Instead, there is sedimentation occurring at the upstream section of the Napique/Ramón Lagoon which acts as a barrier and reduces the velocity of incoming flows (Reyes Salazar, 2024b). As such, the future behavior of the river is increasingly difficult to predict.

The course towards Sechura has larger velocities and larger excess shear values. This is expected since the general slope of this reach is larger than that towards Virrilá. There is, however, potential for sedimentation in the areas where the excess shear decreases. These areas with local sedimentation are in part also due to the Napique/Ramón Lagoon. In the model, it was assumed that
this lagoon fills with water, this however may be different in practice as discussed in the next section and thereby increases the transport capacity in this area.

### 9.2. Additional Sediment Investigations

Additional investigation of sediment transport can be achieved through comparison of different transport formulas considering the large difference in results from the HEC-RAS shear output based on slope and hydraulic radius compared to the calculated shear stress based on the drag law. The actual transport capacity of the river is difficult to pinpoint since sediment transport may be quite unpredictable. These are only two of the many calculation methods available for shear stress, perhaps a range of potential transport capacity values can be determined using more formulas. However, for the purpose for this investigation, the HEC-RAS output and drag law shear stress provided sufficient information to indicate where sedimentation and erosion is likely to occur.

A detailed analysis of suspended sediment transport is also relevant for the Piura River. The river not only interacts with sediment in the riverbed but also with sediment carried from further upstream (Reyes Salazar, 2024b). This is generally clayey-silty sediment which deposits in the lower spans of the river as seen in the field visits. The influence of the sediment load entering the lower basin from further upstream may also have a significant impact on the quantity of deposited sediment.

Further, the 1967 study (Int. Engineering Company \& Olazabal y Leon S.A) mentioned by the PNUD project provides a hydrometer analysis for gradation fine sediments for the Piura River, as well as a rating curve for suspended sediment discharge. This information is useful for sediment modelling in

HEC-RAS, however, suspended sediment quantities in the Piura River are likely to have changed in the since the study due to factors such as deforestation. Deforestation has been an ongoing issue in the upper basin and results in increased runoff going into the river as well as increased sediment transport since there are no roots holding sediment in place (Atarama \& Rashid, 2019).

HEC-RAS features sediment transport modelling in one-dimensional models for a quasi-steady flow. In this, the change in the riverbed is considered in the computations and thereby the behavior of sediment can be analyzed over time for a hydrograph. There are many transport equations and a multitude of boundary conditions to be used in varying combinations for computation. These result in varying outputs for the same cross section and sediment data depending on the transport equation combination used. As such, in order to accurately model the sediment movement along the riverbed calibration parameters are necessary for one or many known flows. Calibration data for sediment transport analysis could be in the form of bathymetric data for different years or before and after a large flow event.

### 9.3. Recommendations

In this model, a simplified version of the river course was created for both the course through Virrilá and Sechura. It was noted that the intervals of 1 m in the topography made a difference in whether or not a particular cross section was flooded or not. However, since the areas outside of the dam were marked as ineffective areas, this did not influence the results. Nevertheless, further investigation could include a flood analysis of the two courses using return periods. In this case, a more accurate topography would be necessary.

The Piura River is a complex system with an abundance of potential solutions to reduce vulnerability of surrounding populations. As suggested by Eduardo Zegarra (Zegarra Dávila, 2023), there are multiple routes which can create an increase in energy in the river and allow for adequate sediment transport. The route discussed in this study is one of the several proposals, however, the implications of the other proposals can also be assessed in comparison to the current route through Virrilá in terms of sediment transport capacity.

As mentioned, the Ñapique/Ramón Lagoon was modelled to allow for flooding in the course towards Sechura. However, in order to prevent flows larger than $3,200 \mathrm{~m}^{3} / \mathrm{s}$ from entering Río Loco, a spillway structure can be constructed to control the flow. In this way, flows above $3,200 \mathrm{~m}^{3} / \mathrm{s}$ can be redirected to flood the lagoons in case of extreme events. This spillway structure would require additional investigation for design.

A further investigation would include the effect of maintenance on the course of the river. As mentioned in Chapter 4, during the field visit many reaches were found to be covered in dense brush. The removal of the brush could potentially reduce the roughness of the terrain and allow for faster velocities to develop along the flood plain. An investigation into the effect of river maintenance can also add to the study of sediment transport in the area.

In this research, observations made during field work for land coverage were used to approximate Mannings roughness values. A further investigation could be a systematic analysis of the vegetation in the study area. This can be done by setting a grid for which vegetation type is determined to create a more accurate description of the land coverage, also considering seasonal changes.

A further ongoing investigation is that of upstream reservoirs to control the flow passing through the city of Piura (Reyes Salazar, 2024c). A difficulty encountered with these suggestions, however, is the lack of an integrated
management system for the Piura River. The river crosses several provinces in the region, these however have their own legislation regarding the river; which is often an obstacle for riverine defense design (Reyes Salazar, 2024c). Consequently, there have been efforts to create an entity in charge of risk management against disasters of hydroclimatic origin on a national level which would allow for adequate riverine management (Reyes Salazar, 2024c), however, this entity has yet to be approved.

## 10. Conclusion

The Piura River regularly poses a threat to Piura and the surrounding agricultural land during El Niño events. Damage or destruction of infrastructure such as roads, communication lines, homes, schools, and health establishments has been a major set pack for the region in the past (Atarama \& Rashid, 2019). The ENSO phenomenon causes a dramatic change in the climatic conditions of the Piura region in the form of storms that are transported inland due to increased temperatures along the coast. These storms cause an increase in flow through the non-permanent river. The various anthropogenic changes to the river have resulted in an unstable flow regimen which has increased the vulnerability of populations in the lower basin.

The river generally caused erosion, which allowed for sufficient capacity to be created as large flows passed through the lower basin. However, in the more recent flood of 2017, there was no erosion (Reyes Salazar, 2024b). This resulted in less capacity of the river along some stretches and thereby created an overflow. The extension of the river towards Virrilá in the south instead of through Sechura has moved the river along a route with a smaller slope, thereby causing a reduction in energy (Reyes Salazar, 2024b). Hydraulic conditions were compared for the ancient course towards Sechura and the current course towards Virrilá.

HEC-RAS was used for modeling hydraulic conditions along the different river courses. For this, topography files were used to create a one-dimensional model of the river. The model was calibrated according to data from the 2017 flood event with a flow of $\mathrm{Q}=3,468 \mathrm{~m}^{3} / \mathrm{s}$ and observed water surface elevations for the upper reach, as well as satellite imagery to identify the extent of flooding for an approximate calibration in the lower reach. The river courses were then compared at flows of $\mathrm{Q}=1,000 \mathrm{~m}^{3} / \mathrm{s}$ and $\mathrm{Q}=3,200 \mathrm{~m}^{3} / \mathrm{s}$ in steady state.

The results provided the general water surface, velocity, and shear stress profiles along both courses. The critical shear stress for sediments of a mean grain size of $\mathrm{d}_{50}=0.2 \mathrm{~mm}$ at a water temperature of $25^{\circ} \mathrm{C}$ was found to be $\tau_{\mathrm{cr}}=$ $0.144 \mathrm{~N} / \mathrm{m}^{2}$. This was compared to HEC-RAS shear output and shear stress based on the drag law to calculate excess shear as $\left(\tau_{\mathrm{b}}-\tau_{\mathrm{cr}}\right)^{3 / 2}$. Excess shear was used to indicate whether aggradation or erosion happens along each profile depending on the general trend of the curves. Despite fluctuations in the output, it was found that the course towards Virrilá shows a general trend of sedimentation in the lower reach while the excess shear along Sechura is considerably higher.

The planned redirection of the river towards Reventazón is suspected to aggravate the issue of sediment deposition along the course of the river (Reyes Salazar, 2024c). Not only is the reach longer than the current reach through Sechura, but it also crosses into a different watershed as shown in Figure 80. Experts in the area also argue that this redirection is also more expensive than other alternatives (Zegarra Dávila, 2023). As such, studying alternative proposals is crucial to better understand what the options are for the Piura River.


Figure 80- Map showing Piura River basin and Cascajal River basin in Piura Region and Lambayeque (adapted from Zegarra Dávila, 2023)

The non-permanent nature of the Piura River makes for difficult prediction of flow regimen. Redirecting the river towards the original course may allow for a more predictable flow regimen due to the increased velocities and sediment transport capacity. The flow regimen would likely return to general erosion along the riverbed. Damage due to erosion can however be prevented through consideration of the potential velocities during design of riverine structures and through reconstruction of currently damaged structures. To conclude, the Piura region will continue to be affected by the recurring ENSO Phenomenon. As such, it is imperative that the region continues to investigate the flow regimen of the Piura River to better prepare for dramatic increases in flow and strive to increase resilience.

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## Appendix A

## Field visits

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## Introduction to Field Visits

Field visits were conducted to gain perspective on the study area. Insight was gained on sediment types, vegetative coverage, riverine structures, as well as the state of protective structures. This appendix presents images taken during the excursions and summarizes comments and discussions held during the visits. Figure 1 shows the route the urban area covered, and Figure 2 shows the route through the rural area.


Figure 1- Overview map rural area (Google Earth, 2024)


Figure 2- Overview map rural area (Google Earth, 2024)

## Urban Area

Field visit: 07.02.2024

The first visit was conducted to the urban reach of the river. The field visit to the urban river reach was a part of the University of Piura`s master's program for Water Resources. The visit was led by Jorge Reyes Salazar, lecturer in hydraulics and sediment transport.

## Los Ejidos Dam

The visit began at Los Ejidos Dam. Figure 3 shows the gated spillway on the right and the weir structure (which collapsed in the 80s) on the left. The Los Ejidos Dam is situated approximately 3 km from Piura city center.


Figure 3- Los Ejidos Dam


Figure 4- Ramp into the floodplain at Los Ejidos

The floodplain just downstream of the dam shows some of the riverbed clearance works which have been conducted. Sediment accumulated at approximately 3 m in height as shown in Figures 3 and 4. It has been/is in the process of being removed.


Figure 5- Sediment build up at Los Ejidos view from bottom


Figure 6- Sediment build up at Los Ejidos view from top

Sediment in the floodplain in this area was found to be mostly sandy, and also clayey/silty-sand closer to the riverbed as shown in the figures below.


Figure 7- Sandy sediment at Los Ejidos


Figure 8- Clayey sediment at Los Ejidos

The dam is fed by a channel from the Chira River which provides a steady flow of $60 \mathrm{~m} 3 / \mathrm{s}$. On the day of the excursion, a low flow was observed, which may have been exclusively from the Chira River. The gauge at Los Ejidos indicating potential water levels which can be reached is shown in Figure 9.


Figure 9- Gauge at Los Ejidos

Starting at Los Ejidos, the river through the urban area is channeled by concrete banks as shown in Figure 10. The dry side of the bank is constructed of compacted sediment. However, local erosion can be seen which exists in different scales and often causes holes large enough to cause instability of the banks, as shown in Figure 11.


Figure 10- Concrete riverbank


Figure 11- Hole on 'dry side' of riverbank

## Cáceres Bridge

The channel section at the Cáceres bridge is approximately 140 m .


Figure 12- Cáceres Bridge

Some sedimentation and vegetative growth is visible at the left bank of the Cáceres bridge as shown below in the view downstream of the bridge in Figure 13.


Figure 13- View downstream of Cáceres Bridge

## Sánchez Cerro Bridge

The channel section at the Sánchez Cerro Bridge is approximately 130 m .


Figure 14- Sánchez Cerro Bridge

As was seen at the Cáceres Bridge, there is significant vegetative growth along the right bank as shown in Figure 15. A gauge was also found at this point of the river.


Figure 15- Below Sánchez Cerro Bridge

## Bolognesi Bridge

The channel section at the Bolognesi Bridge is approximately 150 m .


Figure 17- Bolognesi Bridge

The Bolognesi Bridge marks the end of the Urban zone of the river. Similar to the other bridges, there is aggradation occurring close to the concrete banks with some vegetative growth as shown below in the upstream view of the bridge in Figure 18.


Figure 18- View upstream of Bolognesi Bridge

The gauge on the right bank is partially covered by sediments as shown in
Figure 19.


Figure 19- Gauge at Bolognesi Bridge


Figure 20- Below Bolognesi Bridge

## Rural Area \& Exit through Sechura

Field Visit: 28.02.2024

Biaggio Arbulú Channel
The Biaggio Arbulú channel flows along the right side of the river in the rural area, it provides a steady supply of water for irrigation in the lower basin.


Figure 21- Biaggio Arbulú Channel


Figure 22- Biaggio Arbulú Channel


Figure 23- Biaggio Arbulú Channel

## Independencia

The Indpenendencia bridge has a width of approximately 360 m .


Figure 24- Independencia Bridge

On the bridge, there is a lot of sandy sediment which has been deposited by the wind. Following the riverbanks comes there is vegetation as shown in

Figure 25.


Figure 25-Eolian Sediment on Bridge


Figure 26- View upstream of Bolognesi


Figure 27- Downstream of Bolognesi


Figure 28- Old station point

## Between Independencia and Drain 13.08

Figures 29 and 30 show farmland inside the flood plain. There are small irrigation channels along the dikes.


Figure 29- Farmland


Figure 30- Farmland

The drain 13.08 drains water from Castilla. The sediment around the river in this section is very clayey as shown in Figure 32.


Figure 31- Drain 13.08 on right side of river


Figure 32- Clayey sediment at Drain 13.08

The drain crosses the Piura River in a pipeline. The location of the pipeline is visible across the Piura River. This may be due to erosion in this section of the river. As shown in Figure 34 many rows of the river defense have been revealed due to transport of sediment. This may put the pipeline in risk of failure.


Figure 33- Drain 13.08 crossing through Piura River


Figure 34- Erosion close to Drain 13.08

## Upstream of Ramon Lagoon

This section is just upstream of the lagoon system.


Figure 35- Just upstream of Ramon Lagoon
The banks are very sandy. Some clayey sediments are also present again. The deposition in this area is evident as seen by this dried sediment with a thickness of approximately 15 cm as shown in Figure 37 .


Figure 36- Sandy riverbank


Figure 37- Clayey sediment deposit

The vegetation up to this point just inside of the dikes is very dense in some stretches. Following this section the pilot channel changes course towards Sechura.


Figure 38-Dense vegetation

## Old dikes along the lagoons

The lagoon system is protected by dikes. These feature a very steep slope of 1:1 and are made of large stones. However, this may be deemed unnecessary for the current system since the flow in this section is less aggressive.


Figure 39- Old dikes along Lagoon

## Old River Course

The location where the river used to pass towards Sechura is shown in Figures 40 and 41 where Figure 40 is inside of the dike and Figure 41 is outside towards Sechura. The two sides are separated by the dike structure shown in Figure 42 and could be easily reconnected.


Figure 40- Upstream of dike, in Lagoon


Figure 41-Downstream of dike, outside of the Lagoon


Figure 42- Dike bordering Lagoon

## Terrain from Lagoon to Río Loco

The terrain to cross from the lagoon system to the old river course is desert-
like with sparse vegetation and some salt plains.


Figure 45- Salt plain

## Río Loco

The old river course is called Río Loco and is currently fed by agricultural discharge and by Drain 13.08. This section of the river is influenced by tidal activity as visible in Figure 47 where around 20 cm of sediment is moist.


Figure 46- Río Loco


Figure 47- Tidal influence at Río Loco

## Sechura

Río Loco passes through Sechura. The Sechura bridge is shown in Figure 48.


Figure 48- Sechura Bridge


Figure 49- Downstream view at Sechura Bridge


Figure 50- Upstream view at Sechura Bridge

## Chulliyache

Río Loco currently discharges into the Pacific Ocean through three branches.
The pilot channel is designed to exit through Chulliyache.


Figure 51- River exit at Chulliyache


Figure 52- Chulliyache sign


Figure 53- Upstream view at exit

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## Appendix B

Geometry Data for Virrilá and Sechura

## Contents

Geometry data for Virrilá ............................................. 1
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Table 1- Geometry data for Virrilá

| Station | Reach Length | Manning |  |  | Bottom Level |
| :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | LOB | Channel | ROB |  |
| 117307 | 717.1 | 0.018 | 0.028 | 0.018 | 24 |
| 116590 | 615.3 | 0.018 | 0.028 | 0.018 | 23 |
| 115975 | 593 | 0.018 | 0.028 | 0.018 | 23 |
| 115382 | 501.1 | 0.018 | 0.028 | 0.018 | 22 |
| 114881 | 479.5 | 0.018 | 0.028 | 0.018 | 23 |
| 114401 | 504.6 | 0.018 | 0.028 | 0.018 | 22 |
| 113897 | 459.4 | 0.018 | 0.028 | 0.018 | 22 |
| 113437 | 506.8 | 0.02 | 0.03 | 0.02 | 22 |
| 112930 | 504.1 | 0.02 | 0.033 | 0.02 | 22 |
| 112426 | 472.6 | 0.025 | 0.025 | 0.025 | 22 |
| 111954 | 478.5 | 0.025 | 0.025 | 0.025 | 22 |
| 111475 | 573.5 | 0.025 | 0.025 | 0.025 | 21 |
| 110902 | 560.5 | 0.025 | 0.025 | 0.025 | 21 |
| 110341 | 538.3 | 0.025 | 0.025 | 0.025 | 21 |
| 109803 | 524.1 | 0.025 | 0.025 | 0.025 | 21 |
| 109279 | 568.2 | 0.025 | 0.025 | 0.025 | 20 |
| 108711 | 759.3 | 0.025 | 0.025 | 0.025 | 21 |
| 107951 | 584.3 | 0.03 | 0.025 | 0.03 | 20 |
| 107367 | 517 | 0.03 | 0.025 | 0.03 | 19 |
| 106850 | 532.5 | 0.03 | 0.025 | 0.03 | 19 |
| 106318 | 560.9 | 0.03 | 0.025 | 0.03 | 19 |
| 105757 | 452.1 | 0.03 | 0.025 | 0.03 | 18 |
| 105305 | 454.3 | 0.03 | 0.022 | 0.03 | 18 |
| 104850 | 499.6 | 0.03 | 0.022 | 0.03 | 18 |
| 104351 | 457.2 | 0.03 | 0.022 | 0.03 | 18 |
| 103894 | 477.4 | 0.03 | 0.022 | 0.03 | 18 |
| 103416 | 499.2 | 0.03 | 0.022 | 0.03 | 18 |
| 102917 | 520.5 | 0.03 | 0.022 | 0.03 | 18 |
| 102397 | 662.8 | 0.03 | 0.022 | 0.03 | 18 |
| 101734 | 552.2 | 0.035 | 0.025 | 0.035 | 19 |
| 101182 | 640.9 | 0.035 | 0.025 | 0.035 | 18 |
| 100541 | 512.3 | 0.035 | 0.025 | 0.035 | 19 |
| 100028 | 589.7 | 0.035 | 0.025 | 0.035 | 18 |
| 99439 | 642.5 | 0.035 | 0.025 | 0.035 | 18 |
| 98796 | 544.8 | 0.035 | 0.025 | 0.035 | 18 |
| 98251 | 601.1 | 0.05 | 0.025 | 0.05 | 18 |
| 97650 | 731.3 | 0.05 | 0.025 | 0.05 | 17 |
| 96919 | 593.3 | 0.05 | 0.025 | 0.05 | 18 |
| 96326 | 697.7 | 0.05 | 0.025 | 0.05 | 17 |
| 95628 | 697.8 | 0.05 | 0.025 | 0.05 | 17 |
| 94930 | 696.1 | 0.05 | 0.025 | 0.05 | 16 |
| 94234 | 675.4 | 0.035 | 0.025 | 0.035 | 16 |
| 93559 | 729.9 | 0.035 | 0.025 | 0.035 | 16 |
| 92829 | 647.6 | 0.035 | 0.025 | 0.035 | 16 |


| 92181 | 578.8 | 0.035 | 0.025 | 0.035 | 16 |
| :---: | :---: | :---: | :---: | :---: | :---: |
| 91602 | 816.9 | 0.03 | 0.025 | 0.03 | 15 |
| 90785 | 565.5 | 0.03 | 0.025 | 0.03 | 15 |
| 90220 | 809.2 | 0.03 | 0.025 | 0.03 | 15 |
| 89411 | 790.2 | 0.03 | 0.025 | 0.03 | 15 |
| 88621 | 844.2 | 0.03 | 0.025 | 0.03 | 14 |
| 87776 | 707.2 | 0.03 | 0.025 | 0.03 | 14 |
| 87069 | 595.9 | 0.035 | 0.025 | 0.035 | 14 |
| 86473 | 902.7 | 0.035 | 0.025 | 0.035 | 14 |
| 85571 | 808.7 | 0.035 | 0.025 | 0.035 | 14 |
| 84762 | 1008.5 | 0.035 | 0.025 | 0.035 | 13 |
| 83753 | 739.6 | 0.04 | 0.025 | 0.04 | 13 |
| 83014 | 884.6 | 0.04 | 0.025 | 0.04 | 13 |
| 82129 | 1157.9 | 0.06 | 0.025 | 0.06 | 13 |
| 80971 | 1373.1 | 0.06 | 0.025 | 0.06 | 12 |
| 79598 | 939.5 | 0.06 | 0.025 | 0.06 | 12 |
| 78659 | 930.1 | 0.08 | 0.025 | 0.08 | 12 |
| 77729 | 1032 | 0.08 | 0.025 | 0.08 | 11 |
| 76697 | 1008.3 | 0.08 | 0.025 | 0.08 | 11 |
| 75688 | 1011.7 | 0.08 | 0.025 | 0.08 | 11 |
| 74677 | 1169.8 | 0.08 | 0.025 | 0.08 | 9 |
| 73507 | 1411.3 | 0.06 | 0.025 | 0.05 | 10 |
| 72096 | 1495.1 | 0.06 | 0.025 | 0.05 | 10 |
| 70600 | 1601.5 | 0.06 | 0.025 | 0.05 | 9 |
| 68999 | 3663.3 | 0.06 | 0.025 | 0.06 | 8 |
| 65336 | 3278.7 | 0.04 | 0.025 | 0.04 | 7.63 |
| 62057 | 2885.9 | 0.04 | 0.025 | 0.04 | 7.88 |
| 59171 | 4721.9 | 0.04 | 0.025 | 0.04 | 7 |
| 54449 | 4227.9 | 0.04 | 0.025 | 0.04 | 6.42 |
| 50221 | 4817.7 | 0.035 | 0.025 | 0.035 | 5.53 |
| 45403 | 4526.9 | 0.035 | 0.025 | 0.035 | 5.38 |
| 40877 | 4827.1 | 0.035 | 0.025 | 0.035 | 5.18 |
| 36050 | 4658.9 | 0.035 | 0.025 | 0.035 | 4.37 |
| 31391 | 5752 | 0.03 | 0.025 | 0.03 | 4.23 |
| 25639 | 4796.9 | 0.03 | 0.025 | 0.03 | 3.6 |
| 20842 | 4734.8 | 0.03 | 0.025 | 0.03 | 2.7 |
| 16107 | 5235.8 | 0.03 | 0.025 | 0.03 | 1.66 |
| 10871 | 3835.8 | 0.03 | 0.025 | 0.03 | 0.71 |
| 7035 | 2454.9 | 0.03 | 0.025 | 0.03 | -0.13 |
| 4580 | 3011.9 | 0.03 | 0.025 | 0.03 | -0.64 |
| 1568 |  | 0.03 | 0.025 | 0.03 | -1.17 |

Table 2- Geometry data for Sechura

| Station | Reach Length | Manning |  |  | Bottom Level |
| :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | LOB | Channel | ROB |  |
| 71617 | 717.1 | 0.018 | 0.028 | 0.018 | 24 |
| 70900 | 615.3 | 0.018 | 0.028 | 0.018 | 23 |
| 70284 | 593 | 0.018 | 0.028 | 0.018 | 23 |
| 69691 | 501.1 | 0.018 | 0.028 | 0.018 | 22 |
| 69190 | 479.5 | 0.018 | 0.028 | 0.018 | 23 |
| 68711 | 504.6 | 0.018 | 0.028 | 0.018 | 22 |
| 68206 | 459.4 | 0.018 | 0.028 | 0.018 | 22 |
| 67747 | 506.8 | 0.02 | 0.03 | 0.02 | 22 |
| 67240 | 504.1 | 0.02 | 0.033 | 0.02 | 22 |
| 66736 | 472.6 | 0.025 | 0.025 | 0.025 | 22 |
| 66263 | 478.5 | 0.025 | 0.025 | 0.025 | 22 |
| 65785 | 573.5 | 0.025 | 0.025 | 0.025 | 21 |
| 65211 | 560.5 | 0.025 | 0.025 | 0.025 | 21 |
| 64651 | 538.3 | 0.025 | 0.025 | 0.025 | 21 |
| 64112 | 524.1 | 0.025 | 0.025 | 0.025 | 21 |
| 63588 | 568.2 | 0.025 | 0.025 | 0.025 | 20 |
| 63020 | 759.3 | 0.025 | 0.025 | 0.025 | 21 |
| 62261 | 584.3 | 0.03 | 0.025 | 0.03 | 20 |
| 61677 | 517 | 0.03 | 0.025 | 0.03 | 19 |
| 61160 | 532.5 | 0.03 | 0.025 | 0.03 | 19 |
| 60627 | 560.9 | 0.03 | 0.025 | 0.03 | 19 |
| 60066 | 452.1 | 0.03 | 0.025 | 0.03 | 18 |
| 59614 | 454.3 | 0.03 | 0.022 | 0.03 | 18 |
| 59160 | 499.6 | 0.03 | 0.022 | 0.03 | 18 |
| 58660 | 457.2 | 0.03 | 0.022 | 0.03 | 18 |
| 58203 | 477.4 | 0.03 | 0.022 | 0.03 | 18 |
| 57726 | 499.2 | 0.03 | 0.022 | 0.03 | 18 |
| 57226 | 520.5 | 0.03 | 0.022 | 0.03 | 18 |
| 56706 | 662.8 | 0.03 | 0.022 | 0.03 | 18 |
| 56043 | 552.2 | 0.035 | 0.025 | 0.035 | 19 |
| 55491 | 640.9 | 0.035 | 0.025 | 0.035 | 18 |
| 54850 | 512.3 | 0.035 | 0.025 | 0.035 | 19 |
| 54338 | 589.7 | 0.035 | 0.025 | 0.035 | 18 |
| 53748 | 642.5 | 0.035 | 0.025 | 0.035 | 18 |
| 53105 | 544.8 | 0.035 | 0.025 | 0.035 | 18 |
| 52561 | 601.1 | 0.05 | 0.025 | 0.05 | 18 |
| 51960 | 731.3 | 0.05 | 0.025 | 0.05 | 17 |
| 51228 | 593.3 | 0.05 | 0.025 | 0.05 | 18 |
| 50635 | 697.7 | 0.05 | 0.025 | 0.05 | 17 |
| 49937 | 697.8 | 0.05 | 0.025 | 0.05 | 17 |
| 49240 | 696.1 | 0.05 | 0.025 | 0.05 | 16 |
| 48543 | 675.4 | 0.035 | 0.025 | 0.035 | 16 |
| 47868 | 729.9 | 0.035 | 0.025 | 0.035 | 16 |
| 47138 | 647.6 | 0.035 | 0.025 | 0.035 | 16 |


| 46490 | 578.8 | 0.035 | 0.025 | 0.035 | 16 |
| :---: | :---: | :---: | :---: | :---: | :---: |
| 45912 | 816.9 | 0.03 | 0.025 | 0.03 | 15 |
| 45095 | 565.5 | 0.03 | 0.025 | 0.03 | 15 |
| 44529 | 809.2 | 0.03 | 0.025 | 0.03 | 15 |
| 43720 | 790.2 | 0.03 | 0.025 | 0.03 | 15 |
| 42930 | 844.2 | 0.03 | 0.025 | 0.03 | 14 |
| 42086 | 729.9 | 0.03 | 0.025 | 0.03 | 14 |
| 41356 | 508.4 | 0.035 | 0.025 | 0.035 | 14 |
| 40847 | 568.9 | 0.035 | 0.025 | 0.035 | 13.9 |
| 40279 | 646.8 | 0.035 | 0.025 | 0.035 | 13.78 |
| 39632 | 1028.4 | 0.035 | 0.025 | 0.035 | 13.66 |
| 38603 | 818.7 | 0.04 | 0.025 | 0.04 | 13.45 |
| 37785 | 677.6 | 0.04 | 0.025 | 0.04 | 13.29 |
| 37107 | 464.3 | 0.06 | 0.025 | 0.06 | 13.15 |
| 36643 | 439.3 | 0.06 | 0.025 | 0.06 | 13.06 |
| 36204 | 570.5 | 0.06 | 0.025 | 0.06 | 12.97 |
| 35633 | 845.2 | 0.08 | 0.025 | 0.08 | 12.86 |
| 34788 | 956.7 | 0.08 | 0.025 | 0.08 | 12.52 |
| 33831 | 926.5 | 0.08 | 0.025 | 0.08 | 12.1 |
| 32905 | 883.9 | 0.08 | 0.025 | 0.08 | 11.76 |
| 32021 | 1114.8 | 0.08 | 0.025 | 0.08 | 11.41 |
| 30906 | 1061.5 | 0.06 | 0.025 | 0.05 | 10.96 |
| 29844 | 1083.2 | 0.06 | 0.025 | 0.05 | 10.54 |
| 28761 | 1350.5 | 0.06 | 0.025 | 0.05 | 10 |
| 27411 | 1441.8 | 0.06 | 0.025 | 0.06 | 9.57 |
| 25969 | 2458.7 | 0.03 | 0.025 | 0.03 | 8.5 |
| 23510 | 2141.1 | 0.03 | 0.025 | 0.03 | 7.51 |
| 21369 | 1555.9 | 0.03 | 0.025 | 0.03 | 6.23 |
| 19813 | 1188.9 | 0.025 | 0.025 | 0.025 | 5.4 |
| 18624 | 2493.6 | 0.025 | 0.025 | 0.025 | 4.8 |
| 16131 | 1852.2 | 0.025 | 0.025 | 0.025 | 4.1 |
| 14279 | 1896.7 | 0.025 | 0.025 | 0.025 | 3.5 |
| 12382 | 1559.5 | 0.025 | 0.025 | 0.025 | 2.9 |
| 10822 | 1908.2 | 0.025 | 0.025 | 0.025 | 2.5 |
| 8914 | 1858.6 | 0.025 | 0.025 | 0.025 | 1.9 |
| 7056 | 2501.6 | 0.025 | 0.025 | 0.025 | 1.3 |
| 4554 | 2683.7 | 0.025 | 0.025 | 0.025 | 0.6 |
| 1870 |  | 0.025 | 0.025 | 0.025 | 0.08 |
|  |  |  |  |  |  |

Table 3- Reference point cross sections

| Refence Point | Route Virrilá | Route Sechura |
| :---: | :---: | :---: |
| Cáceres Bridge | 114881 | 61190 |
| Sánchez Cerro Bridge | 113897 | 68711 |
| Bolognesi Bridge | 112930 | 67240 |
| Grau Bridge | 108711 | 63020 |
| Independencia Bridge | 97650 | 51960 |
| Drain 13.08 | 89411 | 43720 |
| Start Ñapique/Ramon | 68999 | 25969 |
| End Napique/Ramon | 62057 | 23510 |
| Start La Niña | 54449 | - |
| End La Niña | 31391 | - |
| Virrilá Estuary | 25639 | - |

Note: listed cross sections are closest upstream section to the reference point and are not exactly at the reference point.

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## Appendix C

Output and Calculated Results

## Contents

Critical Shear Stress ................................................... 1
Drag Law Shear Stress ........................................... 1
Output and Calculation....................................... 2

## Critical Shear Stress

The parameters in Table 1 are assumed to calculate critical shear stress using Equations 1-3
(1) $\quad \theta_{c r}=\frac{0.3}{1+1.2 D_{*}}+0.055\left(1-\exp \left(-0.02 D_{*}\right)\right)$
(2) $D_{*}=\left(\frac{g(s-1)}{v^{2}}\right)^{1 / 3} * d_{50}$
(3)

$$
\theta_{c r}=\frac{\tau_{c r}}{\rho g(s-1) d_{50}}
$$

Table 1-Critical shear stress parameters

| Parameter | Value | Unit | Note |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: |
| Sediment |  |  |  |  |  |
| $\rho \mathrm{s}$ | 2600 | $\mathrm{~kg} / \mathrm{m}^{3}$ | UDEP \& UNP 2001c |  |  |
| d 50 | 0.0002 | m | UDEP \& UNP 2001c |  |  |
| Water |  |  |  |  |  |
| $\rho \mathrm{w}$ | 997 | $\mathrm{Kg} / \mathrm{m}^{3}$ | at $25^{\circ} \mathrm{C}$ |  |  |
| 0 | $0.893^{*} 10^{-6}$ | $\mathrm{~m}^{2} / \mathrm{s}$ | at $25^{\circ} \mathrm{C}$ |  |  |
| s | General |  |  |  | $\rho \mathrm{l} / \rho \mathrm{w}$ |
|  |  |  |  |  |  |
| K | 0.4 | - | Karmann |  |  |
| g | 9.81 | $\mathrm{~m} / \mathrm{s}^{2}$ | Gravity |  |  |

Critical shear stress is found to be $\boldsymbol{\tau}_{\mathbf{c r}}=\mathbf{0 . 1 4 4} \mathbf{N} / \mathbf{m}^{\mathbf{2}}$

## Drag Law Shear Stress

Shear stress is calculated using Equations 4-7 assuming a rough bed condition.
(4) $\tau_{b}=\rho C_{D} U^{2}$
(5) $\quad C_{D}=\frac{\mathrm{k}}{1+\ln \left(\frac{Z_{0}}{h}\right)}{ }^{2}$
(6) $z_{0}=\frac{k_{s}}{30}$
(7) $k_{s}=2.5 * d_{50}$

The following tables give output result and calculation results for $\mathrm{Q}=1,000 \mathrm{~m}^{3} / \mathrm{s}$ and $\mathrm{Q}=3,200 \mathrm{~m}^{3} / \mathrm{s}$ for Virrilá and Sechura.
Output and Calculation

| General Output |  |  |  | rag law Calculation |  |  |  |  | Bed Load Calculation |  |  | HEC Output |  |  | Bed Load Calculation |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Station | $\underset{\substack{\text { QTotal } \\ \text { (m3/s) }}}{\text { ata }}$ | ${ }_{\text {Hyd depth }}^{(m)}$ | ${ }_{\substack{\text { Vel tot } \\ \text { (m/s) }}}^{\text {a }}$ | 20/h | ${ }^{\circ}$ |  | rb>cter | $\left.\right\|^{(t-\operatorname{trct})^{3 / 2}}$ | ${ }^{\text {ob }}$ | $8^{*}+0$ b- | $\underbrace{\text { (kfmms }}_{\text {asb }}$ |  | rb>cter | (rb-ctr) ${ }^{\text {3/ }}$ | ${ }^{\text {өb }}$ | $8^{8^{*}+0 b^{2}}$ | $\underbrace{}_{\substack{\text { asb } \\ \text { (ksms) }}}$ |
| 1568 | 1000 | ${ }_{2} 2.26$ | ${ }_{0} 0.83$ | 7.37E.06 | 0.00137 | 0.94 | erosion | 0.71 | 0.226 | ${ }_{0} 0.857$ | ${ }^{\text {9.636-06 }}$ | ${ }_{3} .21$ | erosion | 5.37 | 1.707 | 17.845 | 2.00E-04 |
| 4580 | 1000 | 2.01 | 0.58 | 8.299-06 | 0.00140 | 0.47 | erosion | 0.19 | 0.059 | 0.114 | 1.28E-06 | 1.63 | erosion | 1.81 | 0.576 | 3.499 | 3.93E-05 |
| 7035 | 1000 | 1.82 | ${ }_{0} 0.93$ | ${ }_{9} 9.16 E-06$ | 0.00142 | 1.23 | erosion | 1.13 | ${ }_{0}^{0.359}$ | 1.720 | ${ }_{1}^{1.935-05}$ | ${ }_{4.36}$ | erosion | ${ }_{8.66}$ | ${ }^{2.753}$ | 36.537 | 4.10-04 |
| 10871 | 1000 | 1.61 | 0.67 | 1.04E-05 | 0.00146 | 0.65 | erosion | 0.36 | 0.15 | 0.313 | 3.522-06 | 2.32 | erosion | 3.21 | 1.021 | 8.251 | 9.27-05 |
| 16107 | 1000 | 1.39 | 0.42 | 1.20E-05 | 0.00150 | 0.26 | erosion | 0.04 | 0.013 | 0.012 | 1.366.07 | 0.95 | erosion | 0.72 | 0.230 | 0.884 | 9.93 E.06 |
| 20882 | 1000 | 1.04 | 0.86 | 1.60E-05 | 0.00159 | 1.17 | erosion | 1.04 | 0.331 | 1.521 | 1.711-05 | 4.5 | erosion | 9.09 | 2.891 | 39.324 | 4.42-04 |
| 25639 | 1000 | 1.07 | 1.19 | 1.568 .05 | 0.00158 | 2.23 | erosion | 3.01 | 0.957 | 7.885 | 8.411-05 | 6.44 | erosion | 15.80 | 5.023 | 90.071 | 1.01-.03 |
| 31391 | 1000 | 1.25 | 0.34 | ${ }^{1.3535-05}$ | 0.00153 | 0.18 | erosion | 0.01 | ${ }_{0}^{0.002}$ | 0.001 | 7.311-09 | 0.64 | erosion | ${ }_{0} .35$ | 0.111 | 0.297 | 3.33E.06 |
| 36050 | 1000 | 1.53 | 0.14 | 1.09E.05 | 0.00147 | 0.03 | stable | 0.00 | 0.000 | 0.000 | 0.OOE+OO | 0.13 | stable | 0.00 | 0.000 | 0.000 | $0.00+00$ |
| 4087 | 1000 | 1.26 | 0.15 | ${ }^{1.325-05}$ | 0.00153 | 0.03 | stable | 0.00 | 0.000 | 0.000 | 0.00E+00 | 0.19 | erosion | 0.01 | 0.003 | 0.001 | 1.60E.08 |
| 45403 | 1000 | 1.25 | 0.2 | 1.33E-05 | 0.00153 | 0.06 | stable | 0.00 | 0.000 | 0.000 | 0.OOE+00 | 0.31 | erosion | 0.07 | 0.022 | 0.025 | 2.855-07 |
| 50221 | 1000 | 1.17 | 0.12 | 1.42E.05 | 0.00155 | 0.02 | stable | 0.00 | 0.000 | 0.000 | 0.OOE+00 | 0.11 | stable | 0.00 | 0.000 | 0.000 | $0.00+00$ |
| 54449 | 1000 | 0.83 | 0.29 | 2.011-05 | 0.00166 | 0.14 | stable | 0.00 | 0.000 | 0.000 | 0.OOE+OO | 0.87 | erosion | 0.62 | 0.197 | 0.698 | 7.855.06 |
| 59771 | 1000 | 1.15 | 0.79 | ${ }^{1.455-05}$ | ${ }^{0.00156}$ | 0.97 | erosion | 0.75 | ${ }^{0.238}$ | 0.228 | 1.04E-05 | 3.43 | erosion | 5.96 | 1.894 | ${ }^{20.856}$ | 2.34 E .04 |
| 62057 | 1000 | 1.39 | 0.96 | 1.20E-05 | 0.00150 | 1.38 | erosion | 1.37 | 0.436 | 2.330 | 2.588.05 | 8.39 | erosion | 23.68 | 7.529 | 165.88 | 1.86E-03 |
| ${ }^{65336}$ | 1000 | ${ }^{1.56}$ | 0.14 | 1.076-05 | ${ }^{0.000147}$ | 0.03 | stable | ${ }^{0.00}$ | 0.000 | ${ }^{0.000}$ | ${ }^{\text {a }}$ | 0.19 | erosion | ${ }^{0.01}$ | ${ }^{0.003}$ | ${ }^{0.001}$ | 1.1.00.08 |
| ${ }^{69999}$ | 1000 | ${ }^{1.01}$ | 0.21 | ${ }^{1.655-05}$ | ${ }^{0.00160}$ | 0.07 | stable | ${ }^{0.00}$ | ${ }^{0.0000}$ | 0.000 | ${ }^{0.008 F+00}$ | 1.19 | erosion | ${ }_{1}^{1.07}$ | 0.330 | ${ }^{1.588}$ | ${ }^{1.7885-05}$ |
| 70600 | 1000 | 0.64 | 0.29 | 2.60E.05 | 0.00175 | 0.15 | erosion | 0.00 | 0.000 | 0.000 | 3.986-11 | 1.63 | erosion | 1.81 | 0.576 | 3.499 | 3.93E.05 |
| 72096 | 1000 | 0.92 | 0.61 | ${ }^{1.818 .05}$ | ${ }^{0.00163}$ | 0.60 | erosion | ${ }^{0.31}$ | 0.099 | 0.250 | 2.800.06 | 8.11 | erosion | 22.48 | 7.149 | 15.222 | 1.72-.03 |
| 73507 | 1000 | 0.5 | 0.39 | 3.33E-05 | 0.00185 | 0.28 | erosion | 0.05 | 0.016 | 0.016 | 1.82-.07 | 4.24 | erosion | 8.29 | 2.636 | 34.239 | 3.85E-04 |
| 74677 | 1000 | 0.7 | 0.32 | ${ }^{2.388 .05}$ | 0.00172 | 0.18 | erosion | 0.01 | 0.002 | 0.001 | 6.92E-09 | 5.76 | erosion | 13.31 | 4.232 | ${ }_{69.647}$ | 7.82-.04 |
| 75688 | 1000 | 1.14 | 0.25 | ${ }^{1.466-05}$ | ${ }^{0.00156}$ | 0.10 | stable | 0.00 | 0.000 | 0.000 | ${ }^{0.008+500}$ | 2.57 | erosion | ${ }^{3.78}$ | ${ }^{1.202}$ | 10.538 | 1.188.04 |
| 76697 | 1000 | ${ }^{1.11}$ | 0.51 | ${ }^{1.5005} \mathbf{- 0 5}$ | ${ }^{0.00157}$ | 0.41 | erosion | 0.13 | 0.043 | 0.071 | 7.955-07 | ${ }^{8.64}$ | erosion | 24.75 | 7.874 | 176.70 | 1.996.03 |
| 7729 | 1000 | 0.89 | 0.43 | 1.87-.05 | 0.00164 | 0.30 | erosion | 0.06 | 0.020 | 0.023 | $2.544-07$ | 5.24 | erosion | 11.50 | 3.658 | ${ }_{55.971}$ | 6.29E.04 |
| 78659 | 1000 | 0.83 | 0.38 | 2.01E-05 | 0.00166 | 0.24 | erosion | 0.03 | 0.009 | 0.007 | 8.14E-08 | 3.46 | erosion | 6.04 | 1.920 | 21.287 | 2.39E-04 |
| 79598 | 1000 | 0.83 | 0.35 | 2.01E-05 | 0.00166 | 0.20 | erosion | 0.01 | 0.005 | 0.002 | 2.78E-08 | 3.12 | erosion | 5.13 | 1.633 | 16.688 | 1.87E-04 |
| 80971 | 1000 | ${ }^{1.08}$ | 0.29 | ${ }^{1.544-05}$ | ${ }^{0.00157}$ | 0.13 | stable | 0.00 | 0.000 | 0.000 | ${ }^{0.0005+00}$ | 1.99 | erosion | ${ }^{2} .51$ | ${ }^{0.798}$ | 5.699 |  |
| ¢ $\begin{gathered}82129 \\ 83014\end{gathered}$ | 1000 1000 | 0.56 0.83 | ${ }_{0}^{0.47} 0$ | ${ }^{2} 2.985-050$ | 0.00180 0.00166 | 0.40 0.25 | ${ }_{\substack{\text { erosion } \\ \text { erosion }}}$ | 0.13 0.04 0 | 0.041 0.011 | 0.065 0.010 | 7.335-07 $1.085-07$ | 3.9 1.63 | ${ }_{\substack{\text { erosion } \\ \text { erosion }}}$ | 7.28 1.81 | 2.315 0.576 | 28.174 3.499 |  |
| ${ }_{83753}$ | 1000 | 1.21 | ${ }_{0}^{0.73}$ | ${ }_{1}^{2} .388 .05$ | ${ }_{0}^{0.00154}$ | ${ }_{0.82}$ | erosion | ${ }_{0.55}$ | ${ }_{0}^{0.176}$ | 0.591 | ${ }_{6}^{1.64 E-06}$ | ${ }_{4.27}$ | erosion | ${ }_{8.38}$ | ${ }_{2} .665$ | 34.806 | ${ }_{\text {3.91E-04 }}$ |
| 84762 | 1000 | ${ }^{0} 7$ | 1.09 | ${ }^{2.388 .05}$ | 0.00172 | 2.04 | erosion | 2.61 | ${ }^{0.828}$ | 6.032 | 6.78E-05 | 4.47 | erosion | 9.00 | 2.861 | 38.717 | 4.355-04 |
| 85571 | 1000 | 1.25 | 0.71 | ${ }^{1.335-05}$ | 0.00153 | 0.71 | erosion | 0.49 | 0.157 | 0.499 | 5.600.06 | 3.05 | erosion | 4.95 | 1.575 | 15.818 | 1.78E-04 |
| ${ }^{86773}$ | 1000 | 1.44 | 1.07 | 1.16E.05 | ${ }^{0.000149}$ | 1.70 | ${ }^{\text {erosion }}$ | 1.94 <br> 56 | ${ }^{0.617}$ | ${ }^{3.877}$ | ${ }^{4.355-05}$ | ${ }_{6}^{6.64}$ | erosion | ${ }_{1652}^{16.56}$ | ${ }^{5.265}$ | ${ }_{96,636}$ | 1.096-03 |
| ${ }_{8}^{877099}$ | 1000 | ${ }^{0.9}$ | 1.42 | ${ }^{1.855-05}$ | ${ }^{0.00163}$ | ${ }^{3.28}$ | ${ }^{\text {erosion }}$ | ${ }_{5}^{5.56}$ | 1.769 | (18.829 | 2.12-.04 | ${ }^{6.63}$ | ${ }^{\text {erosion }}$ | ${ }_{16.52}^{16.5}$ | ${ }^{5.1525}$ | ${ }^{99.302}$ | 1.08E.03 |
| 87776 | 1000 | 1.2 | 1.03 | ${ }^{1.395-05}$ | 0.00154 | 1.63 | erosion | 1.81 | 0.57 | 3.507 | 3.94E-05 | ${ }^{3.5}$ | erosion | 6.15 | 1.955 | 21.869 | 2.46E-04 |
| 88621 | 1000 | 1.98 | 1.25 | 8.42E-06 | 0.00140 | 2.18 | erosion | 2.91 | 0.926 | 7.129 | 8.011 .05 | 5.87 | erosion | 13.70 | 4.357 | 72.754 | 8.17E.04 |
| ${ }^{89411}$ | 1000 | 1.34 | 1.17 | 1.24E-05 | 0.00151 | 2.06 | erosion | 2.65 | 0.844 | 6.200 |  | ${ }^{4.3}$ | erosion | 8.47 | 2.694 | 35.378 | 3.97E.04 |
| 90220 | 1000 | 1.72 | 0.9 | 9.99E-06 | 0.00144 | 1.16 | erosion | 1.03 | 0.327 | 1.94 | 1.68-.05 | ${ }^{3.34}$ | erosion | 5.71 | 1.817 | 19.593 | 2.208-04 |


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Table 3- Output and results for Sechura

| General Output |  |  |  | Drag Law Calculation |  |  |  |  | Bed Load Calculation |  |  | HEC Output |  |  | Bed Load Calculation |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Station | Q Total | Hyd Depth <br> (m) | Vel tot (m/s) | 20/h | Cd | $\begin{gathered} \mathrm{Tb} \\ (\mathrm{~N} / \mathrm{m} 2) \end{gathered}$ | rb><tcr | (rb--tr) ${ }^{3 / 2}$ | ${ }^{\text {өb }}$ | $\begin{aligned} & 8^{* *}(0 \mathrm{Ob}-2 \\ & \left.\mathrm{O}_{1}\right)^{3} / 2 \end{aligned}$ | $\begin{gathered} \mathrm{csb}^{(\mathrm{kg} / \mathrm{ms})} \end{gathered}$ | Shear Tot (N/m2) | tb><cr | $(\mathrm{tb}-\mathrm{trr})^{3} 3 / 2$ | өb | $\begin{gathered} 8^{\star}(\Theta \mathrm{b}- \\ \theta \mathrm{cr})^{\wedge} 3 / 2 \end{gathered}$ | $\begin{gathered} \text { qsb } \\ (\mathrm{kg} / \mathrm{ms}) \end{gathered}$ |
| 1870 | 1000 | 0.92 | 0.58 | 1.81159E-05 | 0.00163 | 0.55 | erosion | 0.25 | 0.081 | 0.184 | 2.07E-06 | 1.86 | erosion | 2.25 | 0.715 | 4.836 | 5.438-05 |
| 4554 | 1000 | 1.3 | 1.96 | 1.28205E-05 | 0.00152 | 5.82 | erosion | 13.51 | 4.296 | 71.231 | 8.00E-04 | 21.55 | erosion | 99.04 | 31.491 | 1413.709 | 1.59E-02 |
| 7056 | 1000 | 2.89 | 1.46 | 5.76701E-06 | 0.00131 | 2.78 | erosion | 4.28 | 1.360 | 12.682 | 1.42E-04 | 9.21 | erosion | 27.30 | 8.680 | 204.578 | 2.30E-03 |
| 8914 | 1000 | 2.81 | 1.42 | 5.9312-06 | 0.00131 | 2.64 | erosion | 3.95 | 1.255 | 11.248 | 1.26E-04 | 8.63 | erosion | 24.72 | 7.860 | 176.302 | 1.988-03 |
| 10822 | 1000 | 2.82 | 1.43 | 5.91017E-06 | 0.00131 | 2.68 | erosion | 4.03 | 1.282 | 11.612 | 1.30E-04 | 8.66 | erosion | 24.85 | 7.902 | 177.708 | 2.00E-03 |
| 12382 | 1000 | 2.86 | 1.34 | 5.82751E-06 | 0.00131 | 2.34 | erosion | 3.27 | 1.038 | 8.462 | $9.51 \mathrm{E}-05$ | 7.57 | erosion | 20.24 | 6.435 | 130.581 | 1.47E-03 |
| 14279 | 1000 | 1.69 | 1.34 | 9.86193E-06 | 0.00144 | 2.58 | erosion | 3.81 | 1.213 | 10.684 | 1.20E-04 | 5.05 | erosion | 10.87 | 3.455 | 51.385 | 5.77E-04 |
| 16131 | 1000 | 1.62 | 0.98 | 1.02881E-05 | 0.00146 | 1.39 | erosion | 1.40 | 0.444 | 2.370 | 2.66E-05 | 3.06 | erosion | 4.98 | 1.583 | 15.941 | 1.79E-04 |
| 18624 | 1000 | 2.48 | 0.66 | 6.72043E-06 | 0.00134 | 0.58 | erosion | 0.29 | 0.093 | 0.226 | 2.54E-06 | 1.96 | erosion | 2.45 | 0.778 | 5.493 | $6.17 \mathrm{E}-05$ |
| 19813 | 1000 | 2.02 | 0.82 | 8.25083E-06 | 0.00140 | 0.94 | erosion | 0.71 | 0.224 | 0.849 | 9.54E-06 | 3.24 | erosion | 5.45 | 1.732 | 18.241 | 2.05E-04 |
| 21369 | 1000 | 1.57 | 1.24 | 1.06157E-05 | 0.00146 | 2.24 | erosion | 3.05 | 0.968 | 7.623 | 8.56E-05 | 8.09 | erosion | 22.40 | 7.122 | 152.059 | 1.71E-03 |
| 23510 | 1000 | 1.5 | 1.3 | 1.11111--05 | 0.00148 | 2.49 | erosion | 3.59 | 1.142 | 9.762 | 1.10E-04 | 9 | erosion | 26.36 | 8.380 | 194.070 | 2.188-03 |
| 25969 | 1000 | 1.09 | 0.6 | 1.52905E-05 | 0.00157 | 0.56 | erosion | 0.27 | 0.087 | 0.204 | 2.29E-06 | 2.39 | erosion | 3.37 | 1.070 | 8.860 | 9.95E-05 |
| 27411 | 1000 | 0.52 | 0.47 | 3.20513E-05 | 0.00183 | 0.40 | erosion | 0.13 | 0.042 | 0.069 | 7.75E-07 | 4.14 | erosion | 7.99 | 2.540 | 32.387 | 3.64E-04 |
| 28761 | 1000 | 0.4 | 0.46 | 4.16667E-05 | 0.00194 | 0.41 | erosion | 0.14 | 0.043 | 0.072 | 8.13E-07 | 1.13 | erosion | 0.98 | 0.311 | 1.390 | $1.56 \mathrm{E}-05$ |
| 29844 | 1000 | 0.87 | 1.51 | 1.91571E-05 | 0.00164 | 3.74 | erosion | 6.82 | 2.168 | 25.531 | 2.87E-04 | 12.78 | erosion | 44.92 | 14.282 | 431.804 | 4.855-03 |
| 30906 | 1000 | 1.57 | 1.18 | 1.06157E-05 | 0.00146 | 2.03 | erosion | 2.60 | 0.826 | 6.000 | 6.74E-05 | 7.31 | erosion | 19.18 | 6.100 | 120.519 | 1.355-03 |
| 32021 | 1000 | 1.62 | 1.14 | 1.02881E-05 | 0.00146 | 1.89 | erosion | 2.30 | 0.731 | 5.002 | 5.62E-05 | 6.73 | erosion | 16.90 | 5.374 | 99.675 | 1.12E-03 |
| 32905 | 1000 | 1.64 | 1.12 | 1.01626E-05 | 0.00145 | 1.82 | erosion | 2.16 | 0.688 | 4.562 | 5.12E-05 | 6.54 | erosion | 16.18 | 5.144 | 93.321 | 1.05E-03 |
| 33831 | 1000 | 1.67 | 1.11 | 9.98004E-06 | 0.00145 | 1.78 | erosion | 2.09 | 0.664 | 4.329 | 4.86E-05 | 6.32 | erosion | 15.35 | 4.880 | 86.254 | 9.69E-04 |
| 34788 | 1000 | 1.63 | 1.13 | 1.02249E-05 | 0.00145 | 1.85 | erosion | 2.23 | 0.709 | 4.778 | 5.37E-05 | 6.66 | erosion | 16.63 | 5.289 | 97.307 | 1.09E-03 |
| 35633 | 1000 | 1.6 | 0.91 | 1.04167E-05 | 0.00146 | 1.20 | erosion | 1.09 | 0.347 | 1.638 | 1.84E-05 | 4.38 | erosion | 8.72 | 2.772 | 36.929 | 4.155-04 |
| 36204 | 1000 | 1.65 | 0.89 | 1.0101E-05 | 0.00145 | 1.15 | erosion | 1.00 | 0.319 | 1.440 | 1.62E-05 | 4.11 | erosion | 7.90 | 2.512 | 31.843 | 3.58E-04 |
| 36643 | 1000 | 1.67 | 0.88 | 9.98004E-06 | 0.00145 | 1.12 | erosion | 0.96 | 0.305 | 1.351 | 1.52E-05 | 4.02 | erosion | 7.63 | 2.427 | 30.240 | 3.40E-04 |
| 37107 | 1000 | 1.69 | 0.87 | 9.86193E-06 | 0.00144 | 1.09 | erosion | 0.92 | 0.292 | 1.266 | 1.42E-05 | 3.86 | erosion | 7.16 | 2.278 | 27.504 | 3.09E-04 |
| 37785 | 1000 | 1.71 | 0.86 | 9.74659E-06 | 0.0014 | 1.06 | erosion | 0.88 | 0.280 | 1.185 | 1.33E-05 | 3.8 | erosion | 6.99 | 2.223 | 26.515 | 2.988-04 |
| 38603 | 1000 | 1.73 | 0.85 | 9.63391E-06 | 0.00144 | 1.04 | erosion | 0.84 | 0.268 | 1.108 | 1.25E-05 | 3.68 | erosion | 6.65 | 2.114 | 24.597 | $2.76 E-04$ |
| 39632 | 1000 | 1.74 | 0.84 | 9.57854E-06 | 0.00144 | 1.01 | erosion | 0.81 | 0.256 | 1.039 | 1.17E-05 | 3.62 | erosion | 6.48 | 2.061 | 23.668 | $2.66 \mathrm{E}-04$ |
| 40279 | 1000 | 1.74 | 0.83 | 9.57854E-06 | 0.00144 | 0.99 | erosion | 0.77 | 0.246 | 0.975 | 1.10E-05 | 3.5 | erosion | 6.15 | 1.955 | 21.869 | 2.46E-04 |
| 40847 | 1000 | 1.76 | 0.84 | 9.4697E-06 | 0.00143 | 1.01 | erosion | 0.80 | 0.255 | 1.033 | 1.16E-05 | 3.59 | erosion | 6.40 | 2.034 | 23.211 | 2.61 -04 |
| 41356 | 1000 | 1.76 | 0.83 | 9.4697E-06 | 0.00143 | 0.98 | erosion | 0.77 | 0.245 | 0.970 | 1.09E-05 | 3.54 | erosion | 6.26 | 1.990 | 22.460 | 2.52E-04 |
| 42086 | 1000 | 1.82 | 4.23 | 9.15751E-06 | 0.00142 | 25.40 | erosion | 126.92 | 40.354 | 2050.754 | 2.30E-02 | 89.99 | erosion | 851.63 | 270.781 | 35646.481 | 4.00E-01 |
| 42930 | 1000 | 1.69 | 1.48 | 9.86193E-06 | 0.00144 | 3.15 | erosion | 5.22 | 1.660 | 17.109 | 1.92E-04 | 7.78 | erosion | 21.10 | 6.710 | 139.036 | 1.566-03 |
| 43720 | 1000 | 1.79 | 1.37 | 9.31099E-06 | 0.00143 | 2.67 | erosion | 4.02 | 1.279 | 11.567 | 1.30E-04 | 7.24 | erosion | 18.90 | 6.011 | 117.886 | 1.32E-03 |
| 44529 | 1000 | 1.58 | 0.98 | 1.05485E-05 | 0.00146 | 1.40 | erosion | 1.41 | 0.448 | 2.398 | 2.69E-05 | 3.77 | erosion | 6.91 | 2.196 | 26.028 | 2.92E-04 |
| 45095 | 1000 | 1.4 | 1.69 | 1.19048E-05 | 0.00150 | 4.26 | erosion | 8.36 | 2.658 | 34.664 | 3.89E-04 | 7.2 | erosion | 18.74 | 5.960 | 116.396 | 1.31E-03 |
| 45912 | 1000 | 1.74 | 0.89 | 9.57854E-06 | 0.00144 | 1.13 | erosion | 0.99 | 0.313 | 1.403 | 1.58E-05 | 3.43 | erosion | 5.96 | 1.894 | 20.856 | 2.34E-04 |
| 46490 | 1000 | 1.34 | 1.05 | 1.24378E-05 | 0.00151 | 1.66 | erosion | 1.87 | 0.593 | 3.656 | 4.111-05 | 5.81 | erosion | 13.49 | 4.289 | 71.050 | 7.988-04 |


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